Comparison of scanning aerosol LIDAR and *in-situ* measurements of aerosol physical properties and boundary layer heights

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Abstract. The spatial-temporal distribution of aerosol particles in the atmosphere has a great impact on radiative transfer, clouds, and air quality. Modern remote sensing methods as well as airborne *in-situ* measurements by unmanned aerial vehicles (UAV) or balloons are suitable tools to improve our understanding of the role of aerosol particles in the atmosphere. To validate the measurement capabilities of three relatively new measurement systems and to bridge the gaps that are often encountered between remote sensing and *in-situ* observation as well as to investigate aerosol particles in and above the boundary layer, we conducted two measurement campaigns and collected a comprehensive dataset employing a scanning aerosol LIDAR, a balloon-borne radiosonde with the Compact Optical Backscatter Aerosol Detector (COBALD), an optical particle counter (OPC) on an UAV, as well as a comprehensive set of ground-based instruments. The extinction coefficients calculated from near-ground-level aerosol size distributions measured *in-situ* are well correlated with those retrieved from LIDAR measurements with a slope of 1.037 ± 0.015 and a Pearson correlation coefficient of 0.878, respectively. Vertical profiles measured by an OPC-N3 on an UAV show similar vertical particle distributions and boundary layer heights as LIDAR measurements. However, the sensor, OPC-N3, shows a larger variability in aerosol backscattering coefficient measurements with a Pearson correlation coefficient of only 0.241. In contrast, the COBALD data from a balloon flight are well correlated with LIDAR-derived backscatter data from the near ground level up to the stratosphere with a slope of 1.063 ± 0.016 and a Pearson correlation coefficient of 0.925, respectively. This consistency between LIDAR and COBALD data reflects a good data quality of both methods and proves that LIDAR can provide reliable and spatial distributions of aerosol particles with high spatial and temporal resolutions. This study shows that the scanning LIDAR has the capability to retrieve backscattering coefficients near ground level (from 25 m to 50 m above ground level) when it conducts horizontal measurement which isn't possible for vertically pointing LIDAR. These near-ground-level retrievals compare well with ground-level *in-situ* measurements. In addition, *in-situ* measurements on the balloon and UAV validated scanning LIDAR retrievals within and above the boundary layer. The scanning aerosol LIDAR allows us to measure aerosol particle distributions and profiles from the ground level to the stratosphere with an accuracy equal or better than *in-situ* measurements and with a similar spatial resolution.

1 Introduction

The large varieties of aerosol spatial-temporal distributions in the atmosphere cause large uncertainties in radiative forcing
globally (Ramanathan et al., 2001) and these uncertainties have a great impact on climate change simulations (Stocker, 2014).
The distributions and evolution of aerosol are related to the emission of aerosols (Grythe et al., 2014; Tegen and Schepanski,
2018; Hamilton et al., 2022) and their loss pathway (Poreh and Cermak, 1964; Cheng et al., 2011; Xiang et al., 2019; Xue
et al., 2022). In addition, another important factor affecting radiative forcing is aerosol optical properties (e.g. single scattering
albedo (SSA), LIDAR ratio, scattering and absorption coefficients) (Alam et al., 2011; Romshoo et al., 2021), which also have
large varieties for different types of aerosols (Lesins et al., 2002; Floutsi et al., 2022).

Many methods have been used to measure the spatial-temporal distribution of aerosol optical parameters regionally and glob-
ally. One of the most successful instruments for this purpose is the Moderate Resolution Imaging Spectroradiometer (MODIS)
on Terra and Aqua satellites (Filonchyk and Hurynovich, 2020; Qin et al., 2021). MODIS can provide column-integrated op-
tical parameters like aerosol optical depth (AOD), Ångström exponent (AE), and single scattering albedo (SSA) to study the
optical properties of mineral dust (Kaufman et al., 2005; Ginoux et al., 2012), urban aerosol (More et al., 2013; Munchak et al.,
2013), forest fire smoke (MAE, 2009; Huesca et al., 2009) etc. Another successful satellite mission is the Cloud-Aerosol Li-
dar and Infrared Pathfinder Satellite Observations (CALIPSO). CALIPSO combines an active LIDAR instrument with passive
infrared and visible images to probe the vertical structure and properties of thin clouds and aerosols over the globe (Winker
et al., 2009; Wang et al., 2021; Salehi et al., 2021; Tesche et al., 2014). China launched its first space-borne aerosol-cloud high-
spectral-resolution lidar (ACHSRL) on April 16, 2022, which is capable for high accuracy profiling of aerosols and clouds
around the globe (Ke et al., 2022). Also, the Earth Cloud, Aerosol and Radiation Explorer (EarthCARE) is a satellite mission
implemented by the European Space Agency (ESA), in cooperation with the Japan Aerospace Exploration Agency (JAXA), to
measure global profiles of aerosols, clouds and precipitation properties together with radiative fluxes and derived heating rates,
due for launch in May 2024 (Wehr et al., 2023).

In addition to these satellite missions, ground-based remote sensing methods are used to investigate aerosol optical properties
(Adam et al., 2020; Mylonaki et al., 2021). Over the last decades, many ground-based observation networks were established
to investigate aerosol properties regionally and globally. For example, the AERONET (AErosol RObotic NETwork) project is a
federation of ground-based remote sensing aerosol networks that provides globally distributed observations of spectral aerosol
optical depth (AOD), inversion products, and precipitation water in diverse aerosol regimes (Holben et al., 1998; Prasad and
Singh, 2007; Mielonen et al., 2009). The Micro-Pulse LIDAR Network (MPLNET) is a federated network of Micro-Pulse
LIDAR (MPL) systems designed to measure aerosol and cloud vertical structure, and boundary layer heights (Welton et al.,
2006; Lolli et al., 2018). The European Aerosol Research Lidar Network (EARLINET) is a multi-wavelength LIDAR network
designed to create a quantitative, comprehensive, and statistically significant database for the horizontal, vertical, and temporal
distribution of aerosols on a continental scale (Pappalardo et al., 2014a; Marinou et al., 2017).

Aerosol elastic scattering lidar is widely used in lidar observation networks as it can provide detailed information with high
spatial and temporal resolution. However, retrieving backscattering coefficients from this kind of lidar data requires assump-

tions of lidar ratios and reference values (Fernald, 1984b; Klett, 1985b). One of the successfully used technology to overcome this problem is the Raman lidar (Wandinger, 2005; Groß et al., 2015; Baars et al., 2016; Hu et al., 2022). Another widely used technology is the high spectral resolution lidar (HSRL) (Liu et al., 1999) which used narrow-band filter (e.g. atom or molecule

60   filter) to separate signals from molecule and particle backscatter (Piironen and Eloranta, 1994). And this HSRL allows us better to investigate aerosol optical properties (Burton et al., 2012, 2014; Groß et al., 2013). Recently, a HSRL that uses an interferometer as filter has been deployed at other wavelengths. The recently launched Doppler Wind Lidar, ALADIN, uses this technology to measure tropospheric wind profiles on a global scale but can also obtain vertical aerosol profiles (Schillinger et al., 2003).

65   *In-situ* measurements can also help us better understand aerosol optical properties. The most common instruments are the nephelometer and aethalometer, which can measure the wavelength-dependent optical parameters like scattering and absorption coefficients of aerosol particles (Anderson et al., 1996; Zieger et al., 2011; Drinovec et al., 2015). The aerosol optical parameters are determined by particle size distribution, particle shape, and complex refractive index (Bohren and Huffman, 2008; Yao et al., 2022). The size distribution can be measured by different kinds of particle sizers like Scanning Mobility

70   Particle Sizer (SMPS), Optical Particle Counter (OPC), and Aerodynamic Particle Sizer (APS). The aerosol complex refractive index is related to the aerosol chemical composition which can be measured by aerosol mass spectrometry as well as the relative humidity (Zieger et al., 2015). For decades, these *in-situ* aerosol characterization instruments not only provided valuable datasets at ground level (Huang et al., 2019; Jiang et al., 2022) but also were deployed on aircraft, balloons, mountains (Zieger et al., 2012), and unmanned aerial vehicles to get vertical profiles of aerosol parameters (Bahreini et al., 2003; Zhen et al.,

75   2018; Brunamonti et al., 2021).

Although many results have reported aerosol measurements by LIDAR (Matthias and Bösenberg, 2002; Pappalardo et al., 2014b; Hofer et al., 2020), there are fewer reports on comparison of *in-situ* measurement with LIDAR measurement to quantify uncertainties of LIDAR retrievals (Xiafukaiti et al., 2020; Düsing et al., 2018). In addition, most vertical pointing LIDAR systems have overlap gap between the detector's field of view and the laser beam from tens to around one thousand meters,

80   which makes it difficult to get valid measurement near the surface (Wandinger and Ansmann, 2002) to compare with ground level *in-situ* measurements. However, scanning LIDAR can conduct horizontal measurements allowing to get vertical profiles of aerosol particles and boundary layer structure near the ground level (Althausen et al., 2000). In addition, scanning aerosol LIDAR can also determine LIDAR ratios to reduce the uncertainties in the LIDAR retrievals (Fernald, 1984a; Zhang et al., 2022).

85   In recent years, vertical profiles of aerosol are also investigated more and more by Unmanned Aerial Vehicles (UAV) and LIDAR. For example, Liu et al. (2020) used the UAVs and LIDAR to study the vertical distribution of $PM_{2.5}$ and interactions with the atmospheric boundary layer during the development of heavy haze pollution. Ferrero et al. (2019) compared the backscattering coefficient retrieved from LIDAR with that calculated from aerosol size distributions measured by OPC on tethered balloons in the Arctic to study the role of aerosol chemistry and dust composition in a closure experiment. Zhang

90   et al. (2021) compared boundary layer heights retrieved from aerosol LIDAR and tethered balloon measurements in semi-arid regions. Liu et al. (2021) found that wind shear generating turbulence reshaped the vertical profiles of parameters such as

potential temperature (θ) and $PM_{2.5}$ in the nocturnal boundary layer, which was the key factor leading to the development of entrainment at nighttime. Reineman et al. (2016) used ship-launched fixed-Wing UAVs to measure the marine atmospheric boundary layer and ocean surface processes. In addition, the vertical profiles of atmospheric parameters related to aerosol

95   process such as temperature (Zarco-Tejada et al., 2012), relative humidity (Spiess et al., 2007), wind (Spiess et al., 2007) and ozone concentration (Guimarães et al., 2019) are also obtained from UAV flights.

However, to our best knowledge, so far no dedicated comparison of scanning LIDAR measurement with *in-situ* observation has been performed over a wide altitude range and over such a long time period for comparison at ground level (e.g. one month dataset with 10 minute resolution). Also in order to bridge the gaps that are often encountered between remote sensing and

100   *in-situ* observation, we compared datasets on aerosol spatial-temporal distributions and evolution combining remote sensing and *in-situ* measurements. Two field campaigns were conducted employing a scanning aerosol LIDAR, a radiosonde with a backscatter sensor, an OPC on a UAV, and a comprehensive set of ground-level instruments. The first field campaign was conducted in downtown Stuttgart to compare LIDAR retrievals with ground level *in-situ* measurements. The second field campaign was done at the Jülich research center to compare LIDAR retrievals with OPC measurements on a UAV and a COBALD

105   backscatter sensor on a radiosonde. The aim of this work is to compare the different methods in aerosol measurements, to validate scanning LIDAR retrievals, to discuss the uncertainties of the different methods and the boundary layer evolutions from LIDAR and UAV retrievals.

2   Methods

Two field campaigns were conducted in downtown Stuttgart and at Jülich research center to compare scanning aerosol LIDAR

110   measurements with different *in-situ* measurements. The first field campaign was conducted from February 5th to March 5th, 2018 in downtown Stuttgart (9.2024° E 48.7986° N, 247 m above sea level) employing a mobile container and a scanning aerosol LIDAR on the roof of the container. The ground-level *in-situ* measurements deployed in this mobile container provided aerosol particle size distributions, aerosol chemical composition, and meteorological information (Huang et al., 2019). The second field campaign was conducted from July 5th to 12th, 2018 at Jülich research center (6.4131° E, 50.9084° N, 110

115   m above sea level) employing a scanning aerosol LIDAR, a COBALD sensor hosted by a Vaisala RS41-SGP radiosonde, and an OPC on UAV. The scanning LIDAR called KASCAL used in these two field campaigns was developed by Raymetrics (LR111-ESS-D200, Raymetrics Inc.). A UAV (eBee, senseFly) carrying one OPC (OPC-N3, Alphasense Inc.), weather sensors and Global Positioning System (GPS) sensors provided altitude-dependent particle size distribution and also meteorological information above the Jülich research center. In addition, atmospheric parameters like pressure, temperature, relative humidity

120   and wind information from the ground to 30 km above Jülich research center were gathered by a GPS-equipped radiosonde onboard a balloon that carried COBALD to measure altitude-dependent *in-situ* backscattering coefficients at two wavelengths (455 nm & 940 nm) (Brunamonti et al., 2021). The measurements during this work indicated that the basckatter was dominated by smaller particles with low depolarisation ratios so that it seemed justified to use a spherical model to represent these aerosol

particles (Khlebtsov et al., 2005; Moroz, 2009; Wang et al., 2023). Hence, a Mie code (Leinonen, 2016) was used to calculate extinction coefficients and backscattering coefficients from aerosol size distributions for comparison with the LIDAR retrieval.

2.1 Scanning aerosol LIDAR

The 3D scanning LIDAR (KASCAL) used in the above two field campaigns has an emission wavelength of 355 nm and is equipped with elastic, depolarization, and vibrational Raman channels, hence allowing to retrieve extinction coefficients, backscattering coefficients, and depolarization ratios. The laser pulse energy and repetition frequency are 32.1 mJ and 20 Hz, respectively. The laser head, 200 mm telescope, and LIDAR signal detection units are mounted on a rotating platform allowing zenith angles from -7° to 90° and azimuth angles from 0° to 360°. This LIDAR works automatically, time-controlled, and continuously via software developed by Raymetrics. Detailed information can be found at https://www.raymetrics.com/product/3d-scanning-LIDAR, last access: 8 March 2021 (Avdikos, 2015; Zhang et al., 2022). During the first field campaign in downtown Stuttgart, the LIDAR conducted zenith scanns with an elevation angle from 90 ° to 5° in steps of 5°. The measurements at 5° were used over a range representative of an altitude of 25-50 m to compare with ground-level *in-situ* measurements (3.7 m above ground level). It is assumed that these values are comparable within the mixing layer. During the second field campaign at Jülich research center, the LIDAR conducted zenith scans during UAV launch and the measurements at all elevation angles were used to get vertical profiles of aerosols from ground level up to the free troposphere to compare with an OPC measurement on the UAV. In addition, the LIDAR also conducted vertical pointing measurements in the night of July 12th, 2018 at Jülich research center to compare the vertical profiles of backscattering coefficients from LIDAR retrievals and COBALD measurement on board of a radiosonde.

For the data analysis and calibration of the LIDAR system, we followed the quality standards of the EARLINET (Freudenthaler, 2016). For data analysis of zenith scans, we determined the vertical backscattering coefficient profiles using the Klett-Fernald method (Fernald, 1984a; Klett, 1985a). And these vertical profiles of aerosol backscattering coefficients was used as the reference values for other observation angles. In addition, the measured temperatures and pressures from UAVs and balloons were used to calculate the molecular backscattering coefficients which can be used in LIDAR retrievals.

The atmospheric boundary layer height can be determined from LIDAR by using the Haar wavelet transform (HWT) method (Baars et al., 2008). Furthermore, the boundary layer height was retrieved from vertical profiles of potential temperature by using the gradient method (Seidel et al., 2010; Li et al., 2021).

2.2 Ground-level *in-situ* measurements in downtown Stuttgart

The ground-level *in-situ* instruments were deployed in a mobile container that was deployed in a parks downtown Stuttgart. Ambient temperature, relative humidity, wind direction, wind speed, global radiation, pressure, and precipitation data were measured by a meteorological sensor (WS700, Lufft GmbH). Trace gases (O_3, CO_2, NO_2, SO_2) were measured with the gas monitors (Environment S.A). Particle number concentrations were recorded with two CPCs (CPC 3774, 3022, TSI Inc,). Particle size distributions were measured with SMPS (DMA: TSI 3080, TSI Inc; CPC: CPC3022, TSI Inc), and OPC (Fidas200, Palas, Inc). The OPC (Fidas200, Palas, Inc.) continuously measured particles in the size range of 0.18 - 18 μm. The OPC

used Lorenz-Mie theory to determine the particle number size distribution and this size distribution can be used to calculate extinction coefficients via a Mie code (Leinonen, 2016). In this experiment, Fidas200 was operated with a flow rate of 5 L/min and with a time resolution of 1 s.
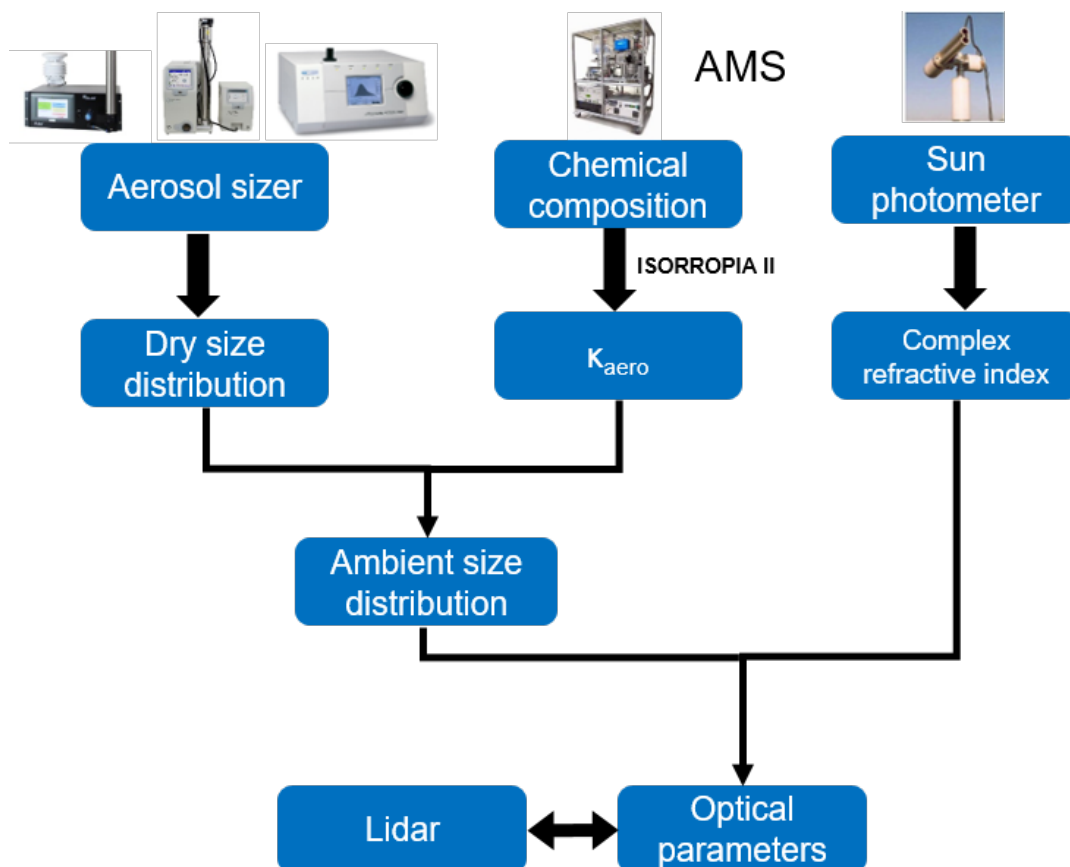
Figure 1. Process flow in deriving aerosol extinction coefficients from Mie calculation and parameters used in Mie calculations. κ_{aero} is the composition dependent hygroscopicity growth factor.

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Figure 1 shows the workflow in deriving the aerosol extinction coefficients from Mie calculations based on *in-situ* aerosol characterisation instruments. The aerosol sizer (e.g. OPC) can provide dry aerosol particle size distribution, which can be converted to the ambient aerosol size distributions by using hygroscopic growth factors (κ) calculated from aerosol chemical composition using the ISORROPIA II thermodynamic equilibrium model (Fountoukis and Nenes, 2007). The aerosol chemical

165 composition was measured by HR-TOF-AMS (Aerodyne Inc.). As most aerosol particles are constrained and well-mixed within the boundary layer, the aerosol complex refractive index remains almost constant (Raut and Chazette, 2008). Although, the sun photometer is integrating over the whole vertical column, the relatively high aerosol concentrations in the boundary layer

dominate (Li et al., 2017). Therefore, it seem justified to use the aerosol complex refractive index derived from a nearby sun photometer (CE-318). Hence, we used the aerosol complex refractive index derived from a nearby sun photometer (CE-318).

170 With ambient aerosol size distribution and complex refractive index, optical parameters (e.g. extinction coefficients) were calculated to compare with LIDAR retrievals.

2.3 UAV and balloon-borne measurements at Research Center Jülich

Data of an OPC (OPC-N3, Alphasense, Inc) on a UAV and a COBALD backscatter sensor (Institute for Atmospheric and Climate Science, ETH Zurich) on a balloon were collected at Jülich research center in July 2018. The UAV used in this field

175 campaign is a fixed-wing drone (eBee, senseFly) which is operated by the Institute of Energy and Climate Research - Troposphere (IEK-8). Its payload is 320 g at a total weight of 750 g with the highest observation altitude of approximately 1200 m above ground level. The ascent and descent velocity of this UAV was around 3.2 m/s. The measurement sensors were mounted inside the UAV. The size distributions were measured in real-time with a time resolution of 1.6 s by OPC-N3. Additionally, atmospheric parameters such as air temperature, air pressure, relative humidity, wind speed, and wind direction were measured

180 with a temporal resolution of 1 s. The UAV was launched 5 times during the morning from 7:00 to 10:00 on July 9th to measure the boundary layer dynamics in the early morning and was launched 7 times from 03:50 to 16:30 on July 12th to measure the boundary layer transition from nocturnal boundary layer to the mixing layer. The detailed UAV fights information can be found in Table 1.

Table 1. Time, altitude, and duration of UAV flights for the experiments on July 9th and July 12th, 2018.

Flight number	Date	Minimum altitude (m a.s.l.)	Maximum altitude (m a.s.l.)	Duration (s)
2018070901	2018.07.09 07:39	90.9	1246.5	709.5
2018070902	2018.07.09 07:48	92.4	1244.8	705.1
2018070903	2018.07.09 08:10	90.9	1243.8	711.7
2018070904	2018.07.09 08:29	89.5	1235.5	691.6
2018070905	2018.07.09 09:34	93.1	1752	1105.5
2018071201	2018.07.12 04:16	91.4	1247.1	701.3
2018071202	2018.07.12 04:31	94.8	1246.1	721.7
2018071203	2018.07.12 07:09	92.7	1246.5	719.6
2018071204	2018.07.12 07:33	93.2	1240.9	717.8
2018071205	2018.07.12 09:44	98.6	1253.7	722.3
2018071206	2018.07.12 14:30	92.8	1248.9	721.3
2018071207	2018.07.12 16:30	92.9	1240.2	716.5

Besides, a radiosonde balloon which was operated by the Institute of Energy and Climate Research - Stratosphere (IEK-7)

185 measured the atmospheric parameters from ground to 25 km altitude. COBALD was part of a CFH / ECC ozone / RS41 payload to provide the backscattering coefficients as well as air temperature, air pressure, relative humidity, wind, and ozone

concentration with the temporal and spatial resolution being 1s and about 5 m vertically.

The COBALD is a lightweight (500 g) aerosol backscatter detector for balloon-borne measurements developed at the Institute for Atmospheric and Climate Science (ETH Zürich), based on the original approach by Rosen and Kjome (1991). Two
190 light-emitting diodes (LEDs) as light sources and a photodiode detector with a FOV of 6° provide high-precision *in-situ* measurements of aerosol backscatter at wavelengths of 455 nm (blue visible) and 940 nm (infrared). COBALD has been originally developed for the observation of high-altitude clouds, such as cirrus (Brabec et al., 2012; Cirisan et al., 2014) and polar stratospheric clouds (Engel et al., 2014), while recently it was proven able to detect and characterize aerosol layers in the upper troposphere–lower stratosphere (Vernier et al., 2015, 2018; Brunamonti et al., 2018, 2021). In this work, we compared
195 COBALD measurements with scanning aerosol LIDAR measurements for validating LIDAR retrievals and investigating the vertical distribution of aerosols. A summary of sensors used on UAV and balloon fights is shown in Table 2.

Table 2. Summary of sensors used on on UAV and balloon fights.

	Measurement	Instrument	Manufacturer	sample flow (lpm)	Time resolution	Mode of operation
UAV						
	Particle size distribution (0.35 - 40 μ m)	OPC-N3	Alphasensor	5.5	1.6 s	24 size bin
	T, RH	ChipCap2 sensor	Telaire		1s	
	Pressure, wind speed & direction	eBee sensors	AgEagle Aerial Systems Inc.		1 s	
	Lat, lon,				1 s	
Balloon						
	Backscattering ratio (455 nm & 940 nm)	COBALD	IAC (ETH, Zürich)		1 s	
	Ozone	Electrochemical concentration cell (ECC)	JOSIE (Smit et al., 2007)		1s	
	Water vapor	Cryogenic frostpoint hygrometer (CFH)	EnSci (Vömel et al., 2007; Vömel et al., 2016)		1s	
	Temperature	Vaisala RS41-SGP	Vaisala		1s	
	Altitude, lat, lon & horizontal wind				1s	

3 Results and Discussion

3.1 Comparison of LIDAR data with ground level *in-situ* measurements
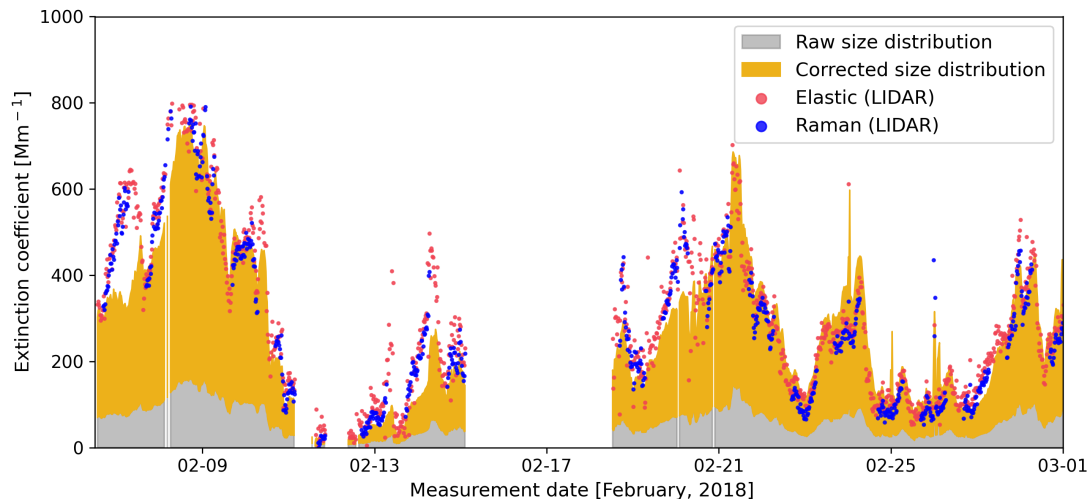
Figure 2. Time series of ground-level extinction coefficients retrieved from LIDAR measurements (both elastic and Raman methods) and from Mie calculations based on OPC raw size distributions as well as size distributions corrected by counting efficiency from February 5th to March 5th, 2018 in downtown Stuttgart.

The comparison of LIDAR retrievals with ground-level aerosol sizer data was conducted during a field campaign from February 5th to March 5th, 2018 in downtown Stuttgart. In this campaign, the aerosol LIDAR did zenith scans with an elevation angle from 90 ° to 5° in steps of 5°. The nearly horizontal measurement at 5° allows to retrieve extinction coefficients near ground level (from 25 m to 50 m above ground level) by using short-range LIDAR data (ranges: 285 m to 570 m) that can be compared with the ground-level in situ measurements (sampled 3.7 m above ground level). The ground-level *in-situ* aerosol sizer, Fidas200, measured the aerosol size distributions which were used to calculate the aerosol extinction coefficients via Mie code. Figure 2 shows the extinction coefficients retrieved from LIDAR measurements and from Mie calculations based on aerosol size distribution (labeled as "Raw size distribution"). The extinction coefficients obtained from LIDAR were both retrieved from the slope and Raman retrieval methods (Seidel et al., 2010; Ansmann et al., 1992). In the slope and Raman retrieval methods, a linear regression was used and the correlation coefficients of linear regressions are 0.99 ± 0.05 and 0.99 ± 0.06 for slope and Raman retrieval methods, respectively. This is also an indication for a rather homogeneous distribution of the aerosol particles within the altitude range from 25 to 50 m corresponding to a range between 285 and 570 m. This figure shows that the extinction coefficients from Mie calculations based on raw OPC size distributions are systematically lower than those from LIDAR retrievals by a factor of 4.70 ± 1.49. The reason for this phenomenon is that the Fidas200 underestimates the particle number by a factor of 2-10 at diameters between 0.25 μm and 0.5 μm when compared with SMPS data as shown in

Figure S1. The left side of Figure S1 shows the number size distribution from Fidas200 and the merged size distribution from

215 SMPS and APS measurements. The right plot of Figure S1 shows the accumulated extinction coefficients calculated from Mie theory based on those two size distributions, which shows the substantial difference by a factor of four. Hence, we conclude that the underestimation of particle numbers from 0.25 μm to 0.5 μm is one of the main reasons for the underestimation of extinction coefficients based on uncorrected OPC data.
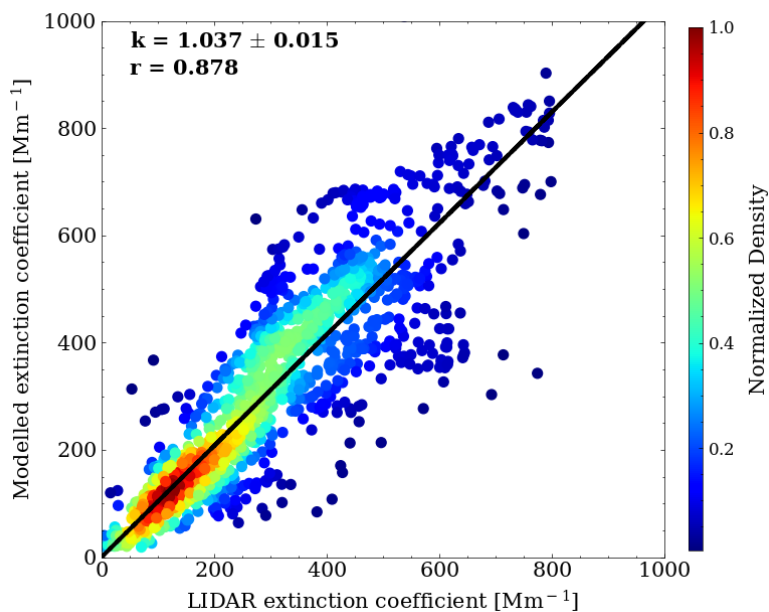
Figure 3. Correlation of extinction coefficients from LIDAR retrieval and Mie calculation from February 5th to March 5th, 2018 in Stuttgart. The relative humidity used in the model is container indoor relative humidity and the black line is the regression fitting curve of them. The red line is the regression fitting curve between the LIDAR-derived extinction coefficients and those from Mie calculation by using ambient relative humidity.

220 Based on systematic laboratory measurements with the differrent particle sizers Fidas200 OPC, SMPS, and APS the FIDAS200 counting efficiency was determined (see Figure S2). This counting efficiency was used to correct all measured size distributions. The corrected size distributions were used to calculate the extinction coefficients via Mie calculation. The time series of the extinction coefficients calculated from the corrected size distribution is shown in Figure 2 (orange area). The calculated extinction coefficients show a reasonable agreement with the LIDAR retrievals.

225 The correlation plot between the extinction coefficient from Fidas 200 and the LIDAR-derived extinction coefficient is shown in Figure 3, which shows a slope and a Pearson correlation coefficient of 1.037 ± 0.015 and 0.878, respectively. As shown in Figure 2, the extinction coefficients retrieved from LIDAR measurement show a similar trend with those calculated based on corrected Fidas 200 size distributions. Please note, that the extinction coefficient based on Fidas 200 data are still a little lower

than those based on lidar measurements. This may be caused by a partial loss of water from the aerosol particles due to higher

230 temperatures inside the container. However, the aerosol particles are not expected to reach equilibrium within the residence time

of 3 seconds in the sampling line inside the warm container. Please note, that there was no dryer in the sampling line. From the

fraction of extinction coefficients shown in Figure 2, we can determine that the main reason for causing extinction coefficient

inconsistency between *in-situ* measurement and LIDAR retrieval is the undercounting by the Fidas 200. The relatively good

agreement of the extinction coefficients after our reasonable corrections reflects the reliability of our methods and the good

235 quality of the LIDAR retrievals.

3.2 Comparison of LIDAR data with *in-situ* measurements on a UAV

The comparison of LIDAR and UAV measurements was conducted for two days, on July 9th and July 12th, 2018 to study the

vertical distribution of aerosols and the boundary layer structure. The sky was almost free of clouds during UAV flights on July

9th while it was affected by clouds within the boundary layer on July 12nd.
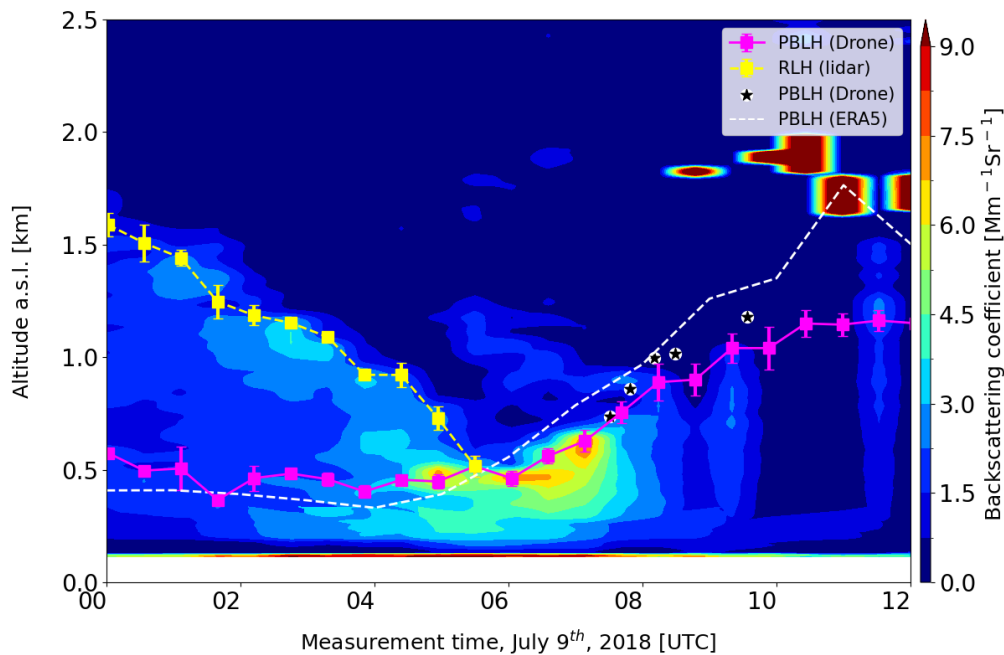
Figure 4. Time series of backscattering coefficients (contour), boundary layer heights (PBLH, pink squares) and residual layer heights (RLH, yellow squares) retrieved from scanning LIDAR, as well as boundary layer heights obtained from UAV measurements (black star with white circle surrounded) and from ERA5 dataset (white dashed line) measured by UAV on July 9th, 2018.

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Figure 4 shows the time series of backscattering coefficients, boundary layer heights (pink squares) and residual layer heights

(yellow squares) retrieved from LIDAR measurement (pink squares) as well as boundary layer heights (a.s.l. - above see level)

obtained from UAV measurements (black star with white circle surrounded) and from ERA5 dataset (white dashed line) on July 9th, 2018. This figure shows that the boundary layer height retrieved from the LIDAR measurement is consistent with the boundary layer height from the UAV measurement which both show an increasing trend of the boundary layer during the morning of this day. In addition, the boundary layer from ERA5 also shows a similar trend as the observations but overestimates boundary layer height, especially during daytime. A possible reason for this overestimation is that the existence of clouds during daytime reduced solar radiation and a low value of solar radiation caused a shallow boundary layer at this time. Figure 4 also shows a stable nocturnal boundary layer and a residual layer during nighttime measured by scanning aerosol LIDAR. The low and stable boundary at night time can suppress the dispersion of aerosol near the surface. Hence, the backscattering coefficients within the boundary layer are maximum (highest aerosol concentration) during the morning rush hour due to the combined effect of the shallow boundary layer and local anthropogenic emissions. After sunrise, the convection became stronger, which caused an increase of the boundary layer height and dilution of aerosols within the boundary layer, so the aerosol concentrations within the boundary layer decreased. Figure S3 shows the time series of range-corrected LIDAR signal and boundary layer heights retrieved from LIDAR as well as boundary layer heights obtained from UAV measurement (black star with white circle surrounded) and from ERA5 dataset (white dash line) on July12^{th}, 2018. The reason for showing range-corrected LIDAR signal instead of backscattering coefficients is that low-level clouds prevented retrieving the backscattering coefficients from range-corrected LIDAR signal by the Klett-Fernald method. This figure also shows consistency in boundary layer heights among LIDAR, UAV, and ERA5. More interestingly, the cloud existed at the top of the boundary layer from 05:00 to 13:00 and the cloud base increased with boundary layer height as captured by the LIDAR measurements. The reason for the cloud existing on the top of the boundary layer is that the relative humidity has a maximum value at the top of the boundary layer in the well-mixing boundary layer and this high relative humidity ambient environment provided a good conditions for cloud formation. Figure S4 shows the correlation of boundary layer heights from LIDAR and radiosonde retrievals for both two days, which show a good correlation with a slope of 1.01 ± 0.24 and a Pearson correlation coefficient of 0.793.

A comparison of the vertical profile of aerosols from LIDAR and UAV measurements was conducted in the following steps. First, we used the temperature and pressure measured by UAV instead of an atmospheric model to calculate molecular backscattering coefficients, and these molecular backscattering profiles were used for LIDAR retrievals. Second, the backscattering coefficients at all observation angles were calculated using the Klett-Fernald method with reference values obtained from vertical profiles of the backscattering coefficients. Finally, Mie theory was used to calculate the aerosol backscattering coefficients based on size distributions measured by the UAV-borne OPC and the complex refractive index from a nearby sun photometer. As there are no dryer before OPC-N3 sampling and no temperature difference between sampling tube and ambient environment, the effect of relative humidity on aerosol sampling was not considered. Figure 5 shows the backscattering coefficients retrieved from LIDAR measurements and from Mie calculations based on size distributions measured by the OPC on the UAV. In this experiment, the LIDAR performed zenith scans using elevation angles from $90°$ to $5°$ with steps of $5°$ during the UAV flights. Consequently, we retrieve the backscattering coefficients for each observation angle and the average of these backscattering coefficients is shown as thick red line to compare with the UAV measurements. The average time is around
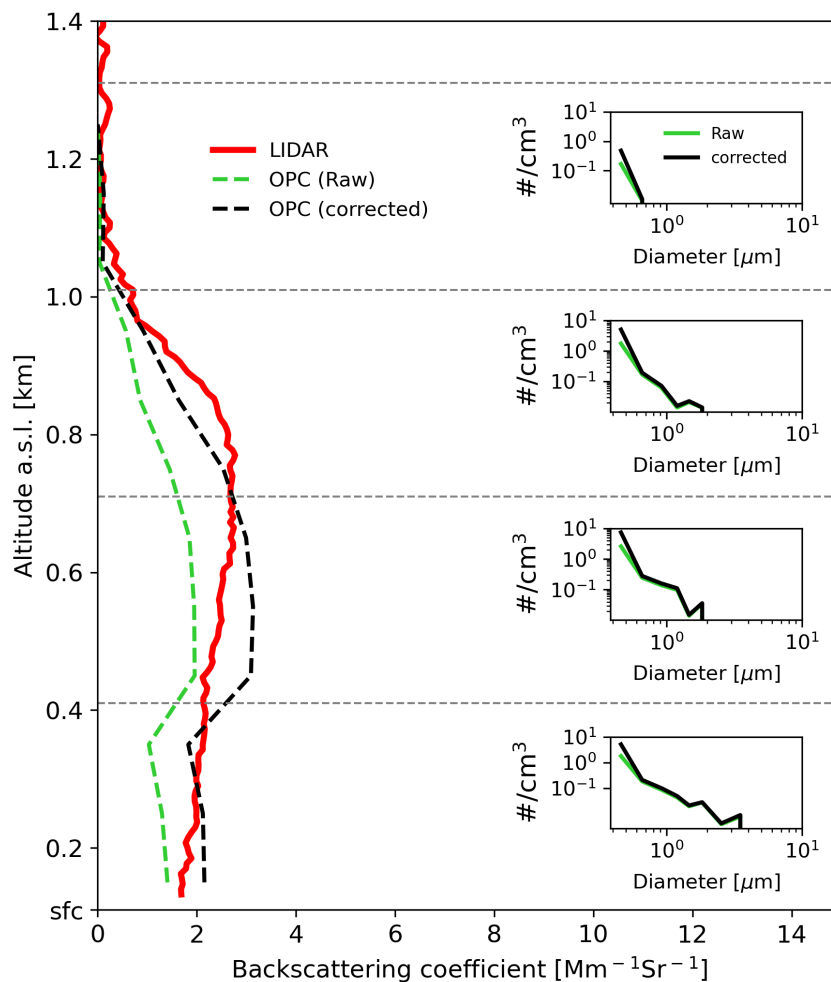
Figure 5. Vertical distribution of backscattering coefficients from LIDAR measurement (solid red line, averaged from 08:14 - 08: 25), as well as backscattering coefficients derived from UAV-based measurements for raw size distributions (dashed green line), and corrected particle size distributions (dashed black line) (inserts on the right) on July 9th, 2018. Note: The 'sfc' on the y-axis indicates ground surface level.

11 minutes for lidar measurement from 08:14 - 08: 25 on July 09, 2018. This figure shows that the vertical distribution of the aerosol particles in the well-mixed boundary layer is reflected well in both LIDAR and OPC measurements. Furthermore, the backscattering coefficients from UAV retrievals (green dashed line in figure 5) show the same aerosol mixing height and the same order of backscattering coefficients as LIDAR retrievals. The smaller backscattering coefficients calculated based on airborne OPC measurements may be related to an undercounting of the smaller particles as we have seen for ground based OPC measurements by the Fidas 200 instrument. The size distributions were corrected by the counting efficiency curve introduced in section 3.1. The backscattering coefficients from corrected size distributions (black dashed line in figure 6) were

285 consistent with the lidar-derived backscattering coefficients. Although Fidas200 is a different OPC sensor as OPC-N3, the same undercounting phenomenon was observed for both sensors. Please note that the particle size is averaged over 300 m and the horizontal dashed lines represent these average altitude ranges. These vertical size distributions show that larger particles were detected only below 300 m above ground level.
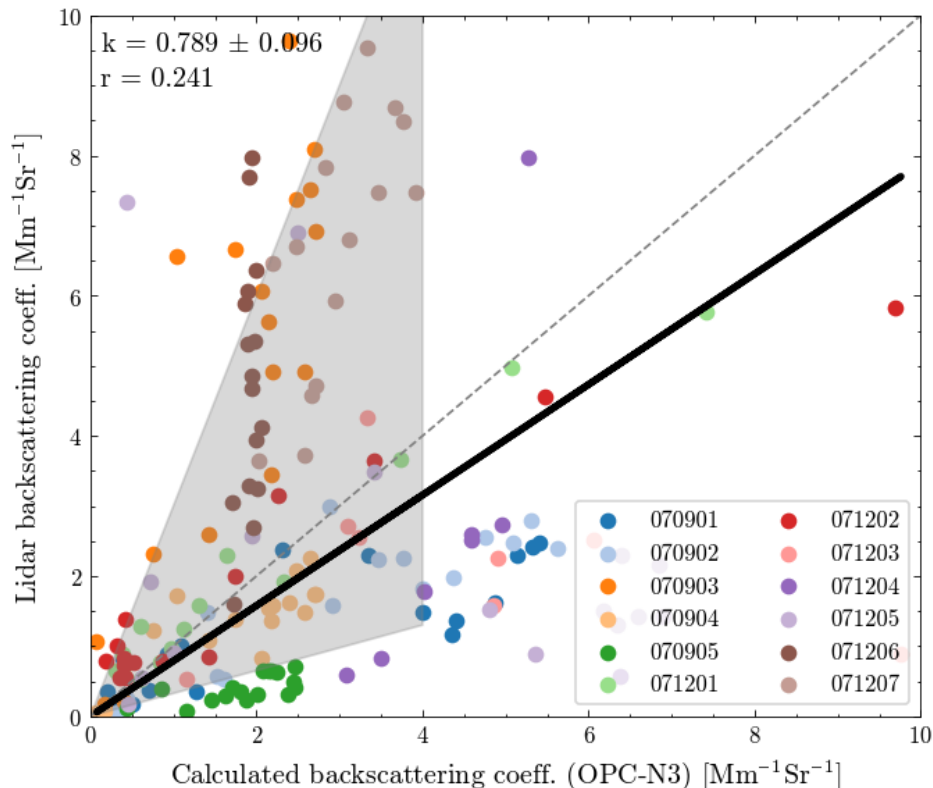
Figure 6. Correlation of backscattering coefficients retrieved from LIDAR measurement and modelled from Mie calculation based on aerosol size distribution measured by OPC-N3 on the UAV for all UAV flights on July 9th and July 12th, 2018. The different scatter point colours indicates different UAV flights. The thick black line is a linear fit to the data and the thin dashed line is the 1:1 line

290 12 UAV flights were conducted on July 9th and July 12th as shown in Table 1 to compare with LIDAR retrievals. Figure 6 shows the correlation of backscattering coefficients retrieved from LIDAR measurement and from Mie calculation based on aerosol size distributions measured by OPC-N3 on the UAV. The data from LIDAR and UAV was averaged into 60 m vertical bins to reduce the noise of the OPC-N3 measurement. The colours of the scatter points indicated different UAV flights. This figure shows that the backscattering coefficients retrieved from LIDAR correlated on average with the backscattering coefficients

295 calculated from the OPC with a slope of 0.789 ± 0.096 and a Pearson correlation coefficient of 0.234. This figure also shows that 75% of data points are within the grey shaded area, which indicates that these data are within a factor of 3. However, in

contrast to the ground level OPC measurements a dedicated correction of the low cost OPC data for potential sampling artefacts or undercounting was not possible. This figure also shows that the UAV measurements reflect the same aerosol mixing process within the boundary layer and the same order of magnitude of the backscattering coefficient. However, the backscatter coefficients retrieved from the UAV-borne OPC in certain UAV flights still show a relatively large deviation from LIDAR retrievals. One reason for this variability is that the UAV cruising speed may affect aerosol sampling by the OPC-N3. The sample were collected perpendicular to the flight direction into the OPC, so we can expect size-dependent discrimination of larger particles. Compared to the Fidas200 OPC as shown in section 3.1, the OPC-N3 data show a significantly higher variability. This means that we must be careful with the quality and the operation of such *in-situ* measurements especially when no reference data like lidar are available.

3.3   Comparison of LIDAR data with *in-situ* measurements onboard a balloon

A balloon which carried the COBALD sensor to measure backscattering coefficients *in-situ* was launched to an altitude of around 30 km on the night of July 12th, 2018 to validate LIDAR retrievals. The LIDAR did vertically pointed measurements with an integration time of 60 s for each profile during the balloon launch. Figure 7a shows the range corrected LIDAR signal for two hours of continuous measurement and the vertical trajectory of the balloon. As shown in this figure, the LIDAR signal did not vary much in the first hour (the period was highlighted in this figure) while showing changes in the second half of the experiment. Hence, we selected the first hour to compare with balloon measurements. Figure 7b shows the horizontal trajectory of the radiosonde with the colour of the plot indicating the radiosonde altitude and the circle indicating the distance from the LIDAR observation station. This figure shows that the horizontal displacement of the radiosonde is about 10 km when the radiosonde reached an altitude of 10 km and this horizontal displacement may cause a difference in backscattering coefficients between LIDAR and COBALD. For the LIDAR analysis in this experiment, the backscattering coefficients were retrieved from elastic and Raman data with the vertical profiles of the molecular backscattering coefficient being calculated from tem-
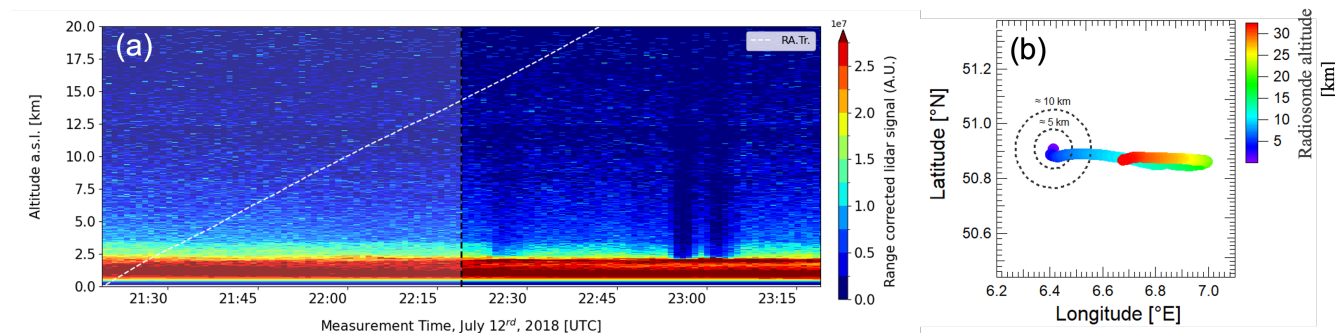
Figure 7. Time series of range corrected LIDAR signal and radiosonde vertical trajectory (white dash line) (a) and horizontal displacement of the balloon during this experiment (b) on July 12th, 2018 at Jülich research center.

perature and pressure measured by the balloon. The COBALD data analysis follows the procedure proposed by Brunamonti et al. (2021). First, a wavelength extrapolation yielded the backscattering coefficient at a wavelength of 355 nm from COBALD

320 measurement. The Ångström exponent (AE) used for this wavelength conversion is measured by COBALD at two wavelengths (455 nm & 940 nm) and extended to the wavelength of 355 nm. Second, as the Field of View (FOV) of LIDAR and COBALD are different (the FOV of COBALD is 6° whereas the FOV of LIDAR is 2.3 mrad), a FOV correction is necessary. The correction factors are calculated based on Mie theory and are shown in Figure 2 in Brunamonti et al. (2021).
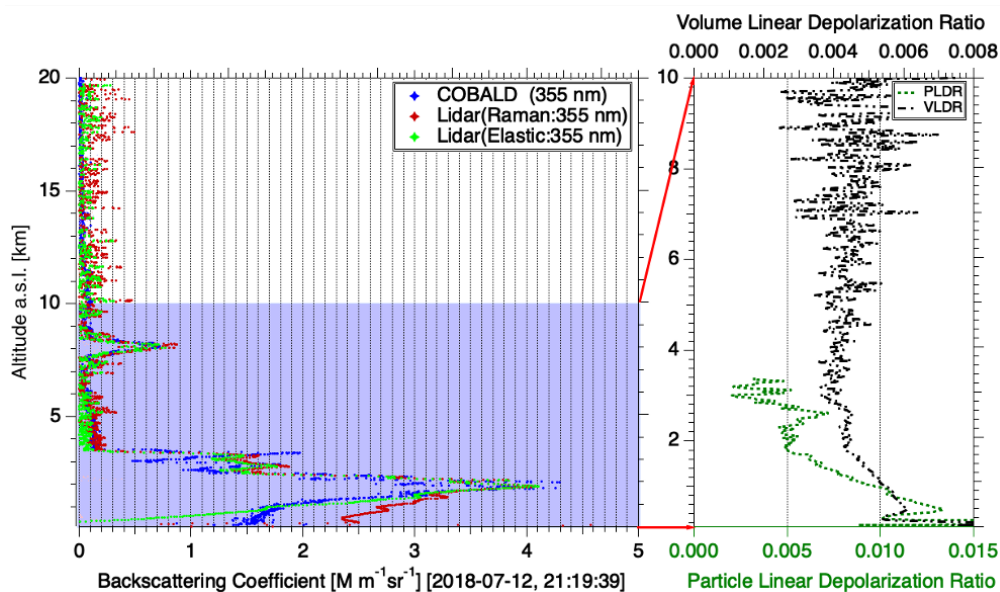
Figure 8. Backscatter coefficients measured by balloon-borne COBALD and LIDAR (left) as well as aerosol volume and particle depolarization ratio measured by LIDAR (right) on the night time of July 12th, 2018 at Jülich research center. (The integration time of the LIDAR data is 1 hour from 21:19 to 22:19.)

Figure 8 shows the backscattering coefficients from COBALD and LIDAR measurement for a LIDAR integration time of 1

325 hour. These two profiles of backscattering coefficients from LIDAR are retrieved from elastic and Raman channel data respectively. The retrieval of backscattering coefficients from elastic channel data remained with larger uncertainty due to the assumption of a LIDAR ratio in the Klett-Fernald method. Hence, it is more meaningful to compare backscattering coefficients from Raman data with those from COBALD measurements. In addition, the volume and particle depolarization ratios measured by LIDAR are shown on the right side of Figure 8. The low depolarization ratios support our assumption that the particles are

330 spherical and that we can use Mie calculations for the FOV correction. This figure shows a good agreement in backscattering coefficients between LIDAR Raman data retrieval and COBALD measurement at an altitude above 2 km. However, there is a significant discrepancy at altitudes below 2 km.

The discrepancy of the backscattering coefficients between LIDAR retrievals and COBALD measurements at lower altitudes is due to the temporal evolution of aerosol particle concentrations in the boundary layer as can be seen from vertical profiles
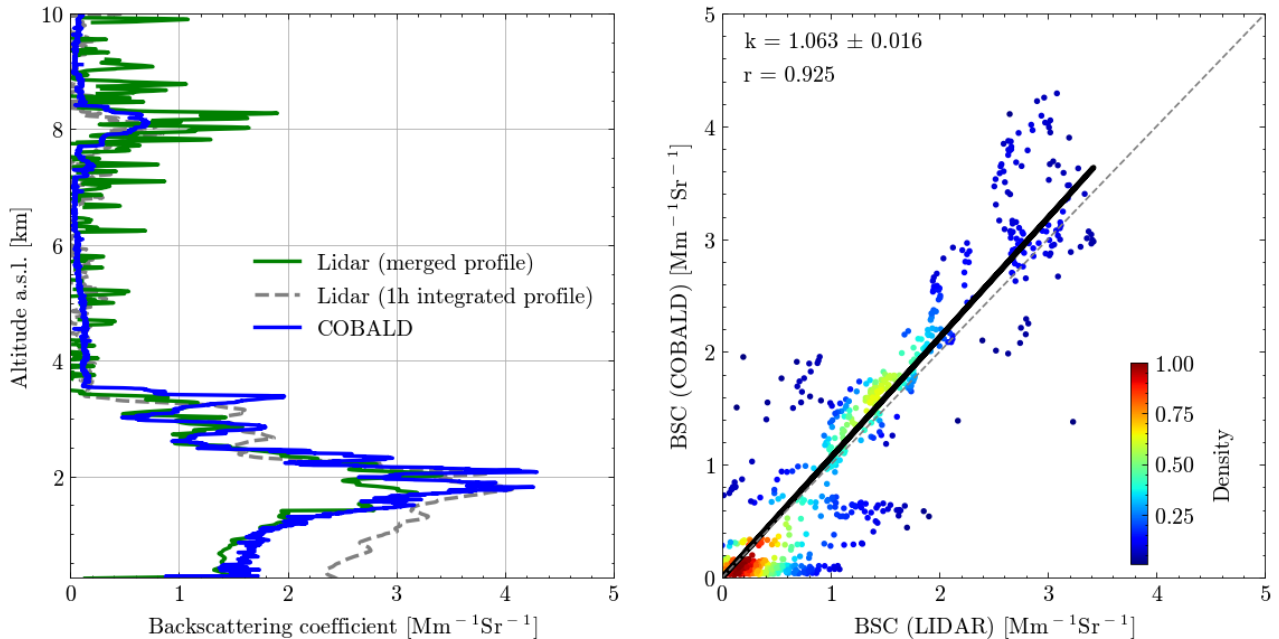
Figure 9. Profiles of backscattering coefficients from LIDAR for integration over 1 hour (grey dash line) and sliding 5-minute merged backscattering coefficients (green line) as well as the vertical profile of *in-situ* backscattering coefficient measured by balloon-borne COBALD (blue line) on July 12th, 2018 at the Jülich research center (left). Correlation between LIDAR merged backscattering coefficients and balloon-borne COBALD backscattering coefficients (right).

335 of backscattering coefficients with high temporal resolution in Figure S5. This figure shows profiles of backscattering coefficients retrieved from LIDAR Raman data with 5 - minute temporal resolution and backscattering coefficients measured by COBALD as well as the vertical balloon trajectory. This figure shows a good agreement in backscattering coefficients between COBALD measurement and LIDAR Raman data retrievals at the altitude of the balloon passing by. The backscattering values at the altitude of the balloon passing by are extracted as shown as the red line in Figure S5 to obtain merged Raman

340 backscattering coefficients. The merged Raman backscattering coefficients and backscattering coefficients from COBALD measurements are shown on the left side of Figure 9, showing very good agreement of backscattering coefficients from LIDAR and COBALD measurements at all altitudes. The correlation between LIDAR merged Raman backscattering coefficients and COBALD backscattering coefficients is shown on the right side of Figure 9, which shows these two backscattering coefficients are well correlated with a slope of 1.063 ± 0.016 and a Pearson correlation coefficient of 0.925. This consistency between

345 LIDAR and COBALD sensor reflects a good data quality of both methods and proves that LIDAR can provide reliable and vertical profiles of aerosol particles with high spatial-temporal resolution.

4 Conclusions

This paper presents results of aerosol spatial-temporal distribution and optical properties measured by a scanning aerosol
LIDAR, a radiosonde with a backscatter sensor, an OPC-N3 on a UAV, and a comprehensive set of ground level *in-situ* mea-
surements. Modern aerosol characterisation methods including remote sensing and *in-situ* methods helped us better understand
the aerosol physical properties and build a bridge between remote sensing and these *in-situ* methods. This paper focuses on the
comparison of aerosol measurement between LIDAR retrievals and *in-situ* measurements at ground level, in the troposphere,
and in the stratosphere, thus validating LIDAR retrievals at all altitude levels.

The comparison of ground-level *in-situ* extinction coefficients with LIDAR-derived ones shows that Fidas200 underestimated
particle number concentration by a factor of 2-10 at the diameter range between 0.25 μm and 0.5 μm, thus causing the total
extinction calculated from this size distribution to be systematically lower than that from LIDAR retrievals by a factor of 4.70
$\pm$ 1.49. The extinction coefficient calculated from the Fidas200 aerosol size distribution corrected by SMPS size distribution
shows good agreement with LIDAR-derived extinction coefficient with a slope of 1.037 $\pm$ 0.015 and a Pearson correction
coefficient of 0.878. The comparison also shows that the undercounting of aerosol particles between 0.25 μm and 0.5 μm is the
main reason for the large discrepancy between LIDAR retrieval and ground-level *in-situ* Fidas200 measurements. In addition, a
comparison between LIDAR and UAV shows good agreement in boundary layer height measurements and both methods show
a similar trend as the ERA5 boundary layer height evolution. The OPC-N3 aboard UAV shows a similar aerosol vertical distri-
bution and comparable backscattering coefficients as LIDAR measurement. However, the backscattering coefficients calculated
from OPC-N3 were unstable and large uncertainties still remained for different flights most likely due to the effect of UAV
cruising on OPC-N3 sampling. Adapting the inlet design of the OPC may improve the data quality for future measurements.
Finally, the backscatter from balloon-borne COBALD measurement shows very good agreement with the backscatter retrieved
from LIDAR measurement if compared with 5-minute resolution LIDAR data with a slope of 1.063 $\pm$ 0.016 and a Pearson
correlation coefficient of 0.925. This consistency between LIDAR and COBALD sensor validated our LIDAR retrievals and
proves that LIDAR can provide reliable and high-resolution vertical profiles of aerosols. And this comparison highlights the
complementary advantages of lidar's continuous measurement capability and COBALD *in-situ* two wavelength data for char-
acterising aerosol particles from near ground level up to the stratosphere. In conclusion, the retrievals from scanning aerosol
LIDAR measurements show good agreement with *in-situ* measurements at all altitude levels and these LIDAR measurements
can also used as reference for other low cost in-situ measurements.

Code availability. The code used to analyse the LIDAR data is property of Raymetrics Inc, but we have shown that it gives the same results
as the code "single calculus chain" (SCC) provided by EARLIENT https://www.earlinet.org/index.php?id=earlinet_homepage, last access:
14 February 2023 and public available. The Mie code used in this paper is available via github repository https://github.com/jleinonen/
pymiecoated, last access: 14 February 2023.

Data availability. The LIDAR raw data and ground *in-situ* data are available via the open access data repository KITopen (link to be added).

380   The UAV data and balloon data are available via the open access data repository Jülich DATA (link to be added).

Author contributions. CR, RT, CW, and HS performed the measurements and analyzed *in-situ* measurement data. HZ analysed the LIDAR remote sensing data. FGW post-processed the COBALD data. HZ wrote the manuscript with support from HS as well as contributions from all co-authors.

Competing interests. The authors declare that they have no conflict of interest.

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