ar-2025-12: "AIDA Arctic transport experiment (part 1): simulation of northward transport and aging effect on fundamental black carbon properties"

Answers of the authors to Reviewer#2

While the reviewer's comments are given in **black bold**, our answers are given below in <u>blue letters</u>. Additionally, we added the changes made in the revised manuscript in <u>blue italic letters</u>.

The manuscript "AIDA Arctic transport experiment (part 1): simulation of northward transport and aging effect on fundamental black carbon properties" is a nicely written report on the ARTEX experiment. The authors used the AIDA simulation chamber to mimic the atmospheric processes that age black carbon particles during transport from the mid-latitudes to the Arctic region. In Part 1, they report changes in the physical properties of the aerosol (e.g. diameter, fractal dimension, coating) and relate them to atmospheric conditions. The manuscript is clear and well written, but some sections need more explanation. These parts are summarised below, with some other technical suggestions.

We would like to thank the referees for their detailed and constructive comments, which helped us to improve our manuscript. Here we provide some major considerations. Several reviewers noted an insufficient discussion on the ambient representativity of transport conditions, soot generation and coating species. This has now been addressed in more detail in both the Methods and Results sections, representing the major modification to the text. Reviewers also identified inconsistencies in the use of acronyms and abbreviations. These have been reviewed and corrected throughout the manuscript text and figures. The reviewer's specific comments are addressed as it follows.

44: "primarily" >> BC is exclusively emitted from combustion sources.

In this case, we intended the BC to be a primary aerosol, we agree the sentence was misleading. Now the text reads:

... "Black carbon (BC) is a primary carbonaceous aerosol emitted by combustion processes" ...

109: "Siberian open fires" >> In fact, many of the BC particles transported to the Arctic come from forest fires. On the other hand, burning biomass releases not only BC but also huge amounts of organic aerosol (OA), which is mostly mixed internally with BC. The OA emissions from the burner used in the experiment are negligible compared to the combustion of biomass. OA ageing and secondary OA contribute to changes in the physical and optical properties of the aerosol mixture. I miss the inclusion of organic components in the experiment. You should mention the importance and role of OA and explain why you focused mainly on nitrate-based coatings.

Siberian biomass burning is an important source of black carbon and organic aerosol during fire season. However, our experiment did not mean to focus on forest fires emissions, but rather on the interaction of elemental carbon rich soot with nitrate species. Nitrate coating has been largely ignored in the past years, and gained more and more importance in view of the increase of nitrate aerosol in the Arctic, which is in contrast to decreasing sulphate and black carbon concentrations (AMAP, 2021). The lack of knowledge on BC-nitrate interaction, was the main reason to design ageing experiments focussing on nitrogen-based species. In this context, ARCTEx did

not aim to reproduce biomass burning events. In fact, the choice of our region of interest (Eurasia), besides being an efficient transport corridor, was justified by relatively homogeneous meteorological conditions compared to other Arctic transport patterns over the Atlantic and Pacific. This homogeneity highly simplified the operation of the AIDA-chamber and increased the representativity of ambient conditions. We do realize, however, that our choice was not properly justified in the main text:

• The text in Section 2.1.1 was modified to better describe the importance of homogeneous meteorological conditions within the region of interest:

... "This sector is associated with intense export of anthropogenic (Backman et al., 2021) and natural (McCarty et al., 2021) black carbon emissions. Although other efficient transport patterns exist, the continental Eurasian sector is characterized by a reduced temperature variability (Przybylak, 2016), ensuring more homogeneous atmospheric conditions compared to the Atlantic and Pacific transport pathways. This uniformity is critical for the design of chamber experiments, as it minimizes external variability and facilitates the representation of the ambient conditions, thereby reducing the operational complexity of the AIDA-chamber." ...

• The text in section 1 "Introduction" was modified to better explain the relevance of nitrate aerosol in the Arctic region:

... "BC variability in the Arctic was often associated with co-emitted sulfate aerosol from anthropogenic sources (e.g. Massling et al., 2015) and organic aerosol from biomass burning events (e.g. Moschos et al., 2022), while its correlation with nitrate was mostly ignored (AMAP, 2021). Similarly, chamber studies focused on the evolution of BC properties as function of internal mixing with sulfate (e.g. Möhler et al., 2005; Khalizov et al., 2009; Henning et al., 2012) and organics (e.g. Lefevre, 2019; Wittbom et al., 2014). As a consequence, the impact of BC-nitrate internal mixing on fundamental and climate relevant properties remained poorly assessed (Yuan et al., 2020). Internal mixing of BC with nitrate species becomes particularly important in the Arctic region, where nitrate aerosol concentration has been increasing since the '80s despite an overall reduction of nitrogen oxides emissions (AMAP, 2021). The same report underlined how few studies had focused on nitrate aerosol in the Arctic, introducing a knowledge gap on the sources of its precursors, its formation mechanisms and its interaction with other atmospheric species such as BC." ...

111: The NO2/BC ratio was fixed according to the CAMS data set. However, the OA/BC ratio was specific to the burner. This means that your experiment represents evolution of the inorganic coating only. This is not a problem, but you should state this fact explicitly. Otherwise, based on the introduction, the reader is expecting a comprehensive simulation of the Arctic transport of BC particles emitted by wildfires.

• To specify that we did not have control on the OA/BC ratio with AIDA, we introduced a dedicated subparagraph at the end of Section "2.1.4 Atmospheric composition" which reads:

... "It must be noted that volatile organic compounds, which are a by-product of combustion, were simultaneously emitted with BC and injected in the AIDA chamber without active control. As a result, the organic aerosol content in AIDA reflects the specific emissions of the burner and not ambient-like conditions. Therefore, although the experiments primarily targeted the evolution of BC mixing with nitrate coatings, the presence of organic vapours may interact or compete with NO_2 during condensation and coating formation, introducing additional complexity to the ageing dynamics." ...

• More details on the production of organic compounds by the mini-CAST burner were added at the end of Section 2.41 "Soot aerosol generation":

... "In addition to BC particles, the mini-CAST burner emits a wide range of volatile and semi-volatile organic compounds (Mamakos et al., 2013; Mason et al., 2020; Daoudi et al., 2023). Although the burner was configured to maximize soot production while minimizing the organic content of the particles (Ess and

Vasilatou, 2019), a certain amount of VOCs was inevitably co-emitted and introduced into the AIDA chamber during injection." ...

183-184: In the results you did not separate BC into hydrophilic and hydrophobic parts. In this respect, this sentence is misleading.

CAMS provides the concentration of the hydrophilic and hydrophobic components of BC. The total BC must be calculated as the sum of the two components. To avoid misunderstanding, the sentence was modified as:

... "While NO_2 is a direct CAMS product, the ARCTEx BC mass concentration was calculated as the sum of the hydrophilic and hydrophobic components of BC provided by CAMS (Li et al., 2024). Hence, in our analysis, BC was treated as a single component without distinguishing the hydrophilic and hydrophobic fractions." ...

274: A plot of the rBC mass distribution and the fit would be useful.

The use of the fitting curve to estimate the rBC mass concentration below and above the SP2 size limits is a common procedure for the SP2 data treatment. However, the reviewer comment rises an interesting point. The mass concentration outside of the size detection range is relatively minor compared to ambient observations, due to the narrow size distribution produced by the cast burner. Overall, the fitting allowed recovering the lost rBC mass below the size detection limit in the first 6 hours of each experiment, when the rBC size distribution peaked near the lower size detection limit. Afterwards, coagulation processes pushed the size distribution within the lower and upper detection limit of the SP2, removing the need for correction. The following statements were added in the main text, while a figure was included in the supplementary.

... "The missing mass was attributed to the lower detection limit of the SP2 and the relatively small size of the injected particles (Figure **Error! Reference source not found.** in Section **Error! Reference source not found.** of the supplementary material). During the first two hours after injection, the undetected mass was estimated to range between 2% and 5%. However, as coagulation-driven growth progressed, the rBC particles fully shifted into the SP2 detection range within approximately six hours, effectively eliminating the initial under-detection issue." ...



Figure S 1 Example of mass size distribution and lognormal fit to estimate the mass concentration of rBC particles below the detection limit of the SP2. Dots: mass size distribution as measured by the SP2, organized in 50 bins covering the 51-370 nm diameter range. Green line: lognormal curve fitted to the measured mass size distribution. Yellow area: estimated mass below the 51 nm, used to correct the rBC mass concentration of the SP2. The example represents the first hours of summer low altitude experiment.

302-303: "Since the concentrations of sulfate, ammonium, and chloride consistently remained below the detection limit" >> yes, because you did not inject.

This makes sense, the sentence was modified as follows:

... "Since sulfate, ammonium, and chloride were not introduced into the AIDA chamber, their concentration consistently remained below the detection limit. Hence, the coating was assumed to consist solely of nitrate and organic components."

304: Source of organic components? The burner only? Why is it representative of forest fires in Siberia? Or for other sources of Arctic BC?

The answer of this comment was split in two parts:

• Answer to "Source of organic components? The burner only?": There were two sources of organic components inside AIDA. The first is primary organic carbon, produced during incomplete combustion and embedded in the soot matrix alongside elemental carbon. The mini-CAST burner was operated to minimize this fraction, which nonetheless remained present in lesser amounts. In our experiments, the organic carbon content at injection time ranged between 10% and 30%. The second source is secondary organic aerosol (SOA), formed through oxidation of volatile organic compounds (VOCs) co-emitted with soot during combustion. While the primary organic fraction was comparable across all experiments, SOA concentrations varied significantly depending on the conditions within AIDA. Details on the primary organic content of fresh soot are provided in Section C of the Supplementary Material, while SOA formation is discussed in Section 3.4 of the Results. To clarify the sources of organic material in AIDA, the text in Section 2.4.3 has been revised as follows:

... "The soot produced by the mini-CAST burner contains minimal organic carbon (Moore et al., 2014), but volatile organic compounds emitted as combustion by-products(Mamakos et al., 2013; Daoudi et al., 2023) may oxidize to form organic coatings (Lim et al., 2019). As a result, organic aerosol was observed during all experiments. In view of these considerations the coating was assumed to consist solely of nitrate and organic components." ...

• Answer to "Why is it representative of forest fires in Siberia? Or for other sources of Arctic BC?": We did not aim to reproduce forest fires emission, as described in the answer to a previous comment.

228-229: Eq2 gives d_f as the slope of the regression if you plot log(m) against $log(k D^d_f)$. In a log-log representation the slope has different meaning. This way the sentence and Fig S1c are ambiguous.

The reviewer is right, the log-log was used as a representation in Figure S1, while the d_f was calculated for each DMA-APM scan using a power law fit as actually described by Equation 2. The text was modified to clarify the message:

... "Following Park et al. (2004), the relationship between particle mass and its diameter is described as a power law, where the exponent represents the fractal dimension (d_f) and k_f a fit constant." ...

353: Equation 4 gives the average material density of a particle with two components. The average material density is the same as the density of the particle if a solid core-shell structure is considered. Here this model cannot be applied since the core is a spongy structure (fractal-like particle), so the "coating" not only covers the surface but also fills the holes and cavities of the core particle. Thus, the coating increases the density through the compactness of the particle even if it has the same material density as the core particle. It would therefore be better to use the term "average material density".

We agree with the consideration of the reviewer's comment, which made us realize that the effective to particle density ratio discussed in the result section was never properly introduced. Thus, we modified the text in Section 2.4.5:

... "We note that this approach assumes a simplified dual-component system with solid and void-free core-shell structure and does not capture the morphology change induced by coating formation. Hence, the calculated ρ_p should be interpreted as an average material density. The ratio ρ_p/ρ_e was used as an indicator of particle compaction. Values near unity suggest a compact particle with internally mixed structure and minimal voids, while lower values indicate a fractal morphology, where internal voids reduce the effective density relative to the material density." ...

380: Eq8 is only valid for spherical solid core-shell structures (for the reason mentioned above). The actual coating thickness can be obtained from the SP2 measurement. It would be interesting to compare the measured and calculated coating thicknesses, which may give an indication of the core particle compactness.

The reviewer is correct. One of the technical objectives of this study was to compare the coating thickness derived from DMA-APM-SP2 measurements (as reported in the manuscript) with that obtained using the leading edge only fit (LEO-fit) method applied to SP2 data. LEO-fit is typically applied within a defined range of rBC core diameters. The specific SP2 instrument used during ARCTEx was previously employed to retrieve coating thicknesses of Arctic rBC particles in the 200-260 nm range (Zanatta et al., 2018). During ARCTEx, rBC core growth occurred via coagulation in all experiments; however, the majority of the rBC number size distribution remained below 200 nm. The fraction of rBC cores exceeding 200 nm was lower than 1% by the end of each experiment, making counting statistics the first limiting factor. A second limitation was related to the encapsulation of rBC cores. While internal mixing led to increased mobility diameters, the optical diameters remained below or close to the detection threshold of the SP2's scattering and position-sensitive detectors. We attempted to apply the LEO-fit method to the WL experiment, when the highest degree of internal mixing was observed, but we were unable to extract a sufficiently clean scattering signal from particles with rBC cores in the 200-250 nm range to apply LEO-fit reliably. Given the high uncertainty associated with coating thickness estimates derived from LEO-fit, we initially hoped that our results could help clarify the robustness of the method. However, we concluded that dedicated experiments are needed to specifically evaluate this technique, based on size selected rBC cores and thicker coatings.

These technical limitations were briefly mentioned at the end of Section 2.4.6 "Volume equivalent diameter and coating thickness", and we have now expanded that discussion to provide more details.

... "It must be noted that the SP2 is also capable of estimating the coating thickness of rBC-containing particles (Laborde et al., 2012) by applying the leading edge only fit (LEO-fit) as proposed by Gao et al. (2007). To apply the method with this specific SP2 model, it is essential to obtain signals from the scattering and position-sensitive detectors with a high signal-to-noise ratio for particles having an rBC-core diameter in the 200-260 nm range (Zanatta et al., 2018). On one side, the number fraction of rBC cores exceeding 200 nm in diameter remained below 1% throughout our experiments, limiting counting statistics. More importantly, even if internal mixing led to increased mobility diameters during the SL and WL experiments, the optical diameters of the rBC-containing particles remained below or close to the detection threshold of the SP2's scattering and position-sensitive detectors. Hence, although the approach proposed here involved the use of several instruments, it represented

the only technical solution to determine the coating thickness for small and thinly coated BC-containing particles." ...

Additionally, a cross-reference to Section 2.4.6 has been added to the SP2 description in Section 2.4.2 "Concentration and size distribution of total and refractory black carbon aerosol".

... "Due to suboptimal combination of low signal-to-noise ratio of the scattering and position sensitive detectors, and the small optical size of the particles present in the AIDA chamber, it was not possible to quantify the coating thickness with SP2 measurements as proposed by Gao et al. (2007). More details are given in Section **Error! R** eference source not found.." ...

Figure 3: The discrepancy between CASM and AIDA RH and NO2/BC values should be discussed. Especially for WL and WH scenarios. How do those differences affect the representativeness of the simulation?

Since this point was raised by multiple reviewers we addressed it in detail. Besides representing a major air transport pattern, the region of interest was chosen also because it shows a limited longitudinal variability in surface temperature, proving relatively stable atmospheric conditions. This was made clear at the end of Section 2.1.1 "Region of interest":

... "Although other efficient transport patters exist, this specific continental area is characterized by a reduced temperature variability (Przybylak, 2016), ensuring more homogeneous atmospheric conditions compared to the Atlantic and pacific transport pathways.

a. To provide a better context to the reanalysis results, a short sub chapter was added in Section 3.1 "Northward transport conditions:

... "These temperature and humidity profiles are consistent with general Arctic conditions, which typically feature stronger latitudinal cooling and more humid conditions near the surface compared to higher altitudes (Przybylak, 2016). The longer atmospheric lifetime of NO₂ during winter (Levy II et al., 1999) may be responsible for the high NO₂/BC in winter then in summer at both low and high altitude scenarios. Although elevated NO₂ relative to BC concentration is a prerequisite for nitrate coating formation, the production pathways in the Arctic are strongly influenced by extreme environmental factors such as low temperatures (Alexander et al.; 2020) and limited sunlight (Schaap et al., 2004), which affect both the efficiency and timing of nitrate formation. As reported by the recent AMAP (2021) report, the limited horizontal and vertical coverage of NO₂ measurements does not allow for further comparison with ambient data." ...

b. Comparability of simulated and measured conditions. Considering that the estimated and observed temperatures were in close agreement we focused on relative humidity and NO₂/BC ratio. The mismatch in RH may have affected gas-phase aqueous-phase chemistry processes controlling the formation of coating-precursors and coating material. The most affected scenario would be WL, when RH values were underestimated by approximately 10-15% during the entire duration of the experiment. The effect of NO₂/BC mismatch would be, however, dominant on the coating formation potential. In this regard, we might have introduced contrasting bias in the low altitude scenarios. We have added a short paragraph at the end Section 3.1 "Northward transport conditions" to explicitly discuss the implications of these discrepancies. The modified text in section 3.1 now reads:

... "The mismatch in RH may have affected gas-phase and aqueous-phase chemistry processes influencing the formation of coating-precursors and coating material. As shown by the high values of the standard deviation of the NO₂/BC ratio measured in the AIDA chamber, the control of NO2 concentration proved to be the most complicated. This technical difficulty led to non-systematic bias during the various experiments. As an example, we introduced contrasting bias in the low altitude scenarios. While the high NO₂/BC in SL might promote a higher degree of internal mixing in AIDA compared to CAMS, the depleted level of NO2 might have hindered coating formation and internal mixing in WL. In summary, atmospheric conditions extracted from ERA-5 and CAMS revealed complex and heterogeneous transport conditions, which were well reproduced, day-by-day, in the AIDA chamber. Nonetheless, discrepancies between reanalysis and measured NO2 levels might accelerate or hinder coating formation in non-systematic way across the experiments." ...

c. Figure 3 was modified to accommodate the comments done by different referees. The modifications include:
i) shading representing the standard deviation of both reanalysis and measurements; ii) different colouring to improve readability; iii) adjustment of axis labels and legend.



Figure 3 Left axis: Latitudinal profiles of temperature and relative humidity extracted from ERA-5 and of NO₂/BC mass ratio extracted from CAMS in the region of interest (40-90°N and 60-140°E). Mean and standard deviation (SD) calculated for equidistant latitude bands 10° wide. Right axis: temporal variability of temperature, relative humidity and NO₂/BC ratio measured in the AIDA chamber. Mean and standard deviation (SD) calculated over 24 hours.

468: Any explanation why the size distributions of the low and high scenarios do differ? And SL and WL? Could the chamber temperature/pressure affect the particle properties after injection?

One parameter for explaining the observed differences in fresh rBC particle size distributions across scenarios could be the duration of injection. During the preparation phase of ARCTEx, we conducted a series of tests in a smaller vessel (aerosol preparation and characterization chamber, APC; Wagner et al., 2024) under ambient conditions and found a high degree of reproducibility in particle diameter during short injections (<1 minute), while longer injection times promoted rapid coagulation and thus larger particles. Considering the long injection time during ARCTEx scenarios (40-50 minutes), minor differences between SH and WH scenarios could be attributed to small variations (minutes) in injection duration. More details on injection duration were added in the method Section 2.4.1 and former Section B of the supplementary material.

Also chamber temperature could play a dominant role. Lower temperatures may enhance the immediate condensation of semi-volatile organic compounds in the first seconds after injection. This rapid condensation could compact the particle structure, potentially leading to a decrease in measured diameter (as seen in Figure S1a). This mechanism may also explain the observed increase in effective density from the warmer experiments

(WS–WL, -9 to -25 °C) to the colder ones (SL–WH, -35 to -52 °C). That said, we do not feel confident elaborating on these mechanisms in the main text, as coagulation processes may counteract condensation-driven shrinkage. A brief note has been added to former Section B of the supplementary material:

... "The low density and diameter values observed in SH and WH suggest that the lower injection temperatures (-35 to -52 °C) compared to the warmer temperatures in SL and WL (25 to -9 °C)may have promoted early condensation of semi-volatile compounds on the soot cores, leading to compaction of the originally ramified structure. However, the fractal dimension and rBC mass fraction did not fully confirm this hypothesis." ...

Figure 4 caption: what do you mean "Concentrations adjusted to the ambient conditions inside the AIDA chamber."

We have now clarified that all concentrations are adjusted to the actual in situ conditions (temperature and pressure) inside the AIDA chamber at the time of measurement, meaning that they are not normalized to standard conditions. This better reflects the physical conditions under which the aerosol processes occurred. The caption of Figure 4 was modified together with the final statement of Section 2.4.2.

- Modified caption: ... "Concentrations normalized to the gas temperature and pressure of the AIDA chamber at the time of measurement." ...
- Modified text: ... "All number and mass concentrations reported hereafter refer to the actual temperature and pressure conditions inside the AIDA chamber (i.e., not normalized to standard conditions) at the time of the measurement." ...

514: Figure 6 is not discussed in the text. 521: Where is the FM_{org} result presented? A relevant figure should be added. 548: Same for FM_{NO3} . 553: same

We thank the reviewer for these observations. Although Figure 6 does show the evolution of the mass fractions of organic and nitrate coatings (FM_{org} and FM_{nit}), it was mentioned only briefly at the beginning of Section 3.4 ("Evolution of BC morphology and mixing state during aging"). This limited referencing, as in the case of Figure 5, aimed to keep the text concise and avoid excessive cross-referencing. However, we recognize that this may have unintentionally made it harder for the reader to connect specific parts of the text with the relevant figures. In response, we revised Section 3.4.1, Section 3.4.2 and Section 3.4.3 to include more frequent references to Figures 5 and 6.

571-572: Do you have any explanation for this? Temperature effect? Slower chemistry?

We hypothesize that temperature plays the dominant role in preventing the formation of thicker coatings. However, several reaction pathways involving ozone, nitrogen oxides and volatile organic compounds might take place and compete.

Considering nitrate coatings formation during day time, the photolysis of ozone is the major production pathway of OH radicals in the atmosphere (Finlayson-Pitts and Pitts, 1986), which in turn oxidizes NO₂ to form HNO₃

- $O_3 + h\nu \rightarrow O(1D) + O_2 \quad (\lambda \le 310 \text{ nm})$
- $0(1D) + H_2O \rightarrow 2OH$
- $NO_2 + OH \rightarrow HNO_3$

During nighttime, the lifetime of NO_3 (formed from the oxidation of NO_2 in presence of ozone) is long enough to trigger the formation of HNO_3 via

- Reaction with volatile organic carbon (Ng et al., 2017): $\circ NO_3(g) + DMS/CH(g) \rightarrow HNO_3(g)$
- Heterogeneous hydrolysis of N₂O₅ (Alexander et al., 2020):
 - $\circ \ NO_3(g) + NO_2(g) \rightleftharpoons N_2O_5(g)$
 - $\circ \quad N_2 O_5(g) + H_2 O(l) \rightarrow H N O_3(l)$

Gas-phase loss of HNO₃ must be considered as well (Dulitz et al., 2018):

- $HNO_3(g) + OH(g) \rightleftharpoons NO_3(g) + H_2O$
- $HNO_3(g) + h\nu \rightarrow OH + NO_2$

According to literature, the photolysis rate of ozone slightly decreases with temperature (Matsumi et al., 2002), potentially decreasing the OH radical formation and nitric acid formation during day time. Cold temperature would however prolong the lifetime of N_2O_5 , promoting the hydrolysis reaction pathway (Alexander et al., 2020), but also increase the reconversion of nitric acid to nitrogen oxides (Dulitz et al., 2018).

Secondary organic aerosol (SOA) coatings may be produced from volatile organic compounds via several pathways (Srivastava et al., 2022). Although cast burners were proven to emit volatile organic compounds, polycyclic aromatic hydrocarbons and alkanes (Mason et al., 2020), their specific reactivity and SOA formation yields were never reported. However, the review of Srivastava et al. (2022) indicated a positive effect of RH on SOA yields, while temperature had a negative effect or little influence on SOA formation due to enhanced partitioning into the condensed phase at lower temperature. At subzero temperature, relevant for the high troposphere, the lifetime of biogenic VOCs in presence of ozone is highly enhanced (Tillmann et al., 2009). Overall, considering the formation of secondary organic aerosol coatings, anthropogenic organic emissions are still poorly constrained (Jathar et al., 2014)

In view of these details, at high temperature and prolonged light conditions (SL), O₃ was consumed quickly at a maximum rate of 10 ppm per hour immediately after injection. However, the hourly decrease of NO₂ remained almost constant, leading to an almost negligible formation of nitrate. In this scenario with warm conditions and low NO₂/BC ratio, SOA formation was clearly dominant. In WL, during irradiation time, the hourly consumption of O₃ was similar to SL, but was followed by a marked decrease of NO₂. This indicates an efficient photolysis of ozone, OH radical formation and conversion to nitric acid and nitrate. In contrast, nighttime formation and deposition of nitrate presumably play a secondary contribution. These two experiments might confirm that the reactivity of VOCs with ozone is extremely efficient above 0 °C and competes with the oxidation of nitrate during day and nighttime. In summarized form, these considerations were provided in the main text in Section 3.4.1 and Section 3.4.2.

During the high-altitude scenarios, we observed an almost constant hourly depletion of NO₂ and O₃, largely independent of light conditions. Notably, their depletion rates closely followed that of CO₂ (used as a dilution tracer), suggesting limited chemical reactivity at temperatures below $-30 \,^{\circ}$ C. This supports the hypothesis that temperature plays a dominant role in suppressing both daytime and nighttime chemistry under these conditions. It must also be considered that the water vapour pressure was much lower in the coldest experiment, potentially limiting OH formation and hydrolysis of N_2O_5 . However, our experimental setup did not permit the identification and quantification of all relevant organic and inorganic precursors, nor the absolute determination of the coating mass. As a result, the temperature effect on the coating formation "yield" remains unquantifiable. Given these limitations, a detailed discussion of specific reaction pathways would be speculative. For this reason, and in line with the scope of this study, we did not elaborate further on this in Sections 3.4.1 and 3.4.2. Nonetheless, we briefly raised our hypothesis and the current limitations at the end of Section 3.4.3 ("Slower aging during high-altitude transport"). We are open to include the extended discussion provided here in the supplementary material, in case considered necessary by the reviewers and editor.

... "This aging reduction, was likely linked to low temperatures, which may slightly reduce the yield of ozone photolysis (Matsumi et al., 2002), shorten the lifetime of HNO₃ (Dulitz et al., 2018), and minimize the oxidation of volatile organic compounds to secondary organic aerosol coatings (Saathoff et al., 2009; Tillmann et al., 2009). This hypothesis is reinforced by the near-constant depletion rates of NO₂ and O₃, irrespective of light conditions,

suggesting limited photochemical and nocturnal reactivity at temperatures below -30 °C. However, due to the lack of comprehensive gas-phase (nitric acid, OH radical, volatile organic compounds) and particle-phase (absolute quantification of nitrate, organic and organo-nitrate compounds) chemical speciation, we cannot quantify the coating yields or attribute them to specific reaction pathway." ...

Figure 5: Note which are the direct measurement results, and which are calculated ones (e.g. particle density)?

All the variables presented in Figure 5 are calculated combining multiple instruments, and they are fully described in the technical section and summarized in the annex table. To avoid repetitions, we chose not to modify the caption of Figure 5.

Figure 5: What are the gaps between the points?

The properties shown in Figure 5 are derived from DMA-APM and DMA-APM-SP2. These measurements were not performed continuously but, usually, before and after the virtual midnight (change of conditions). This is now specified at the end of the introduction of Section 3.4 "Evolution of BC morphology and mixing state during aging":

... "Excluding FM_{nit} and FM_{org}, the other properties described in this section were calculated from the DMA-APM and DMA-APM-SP2 scans. These scans were performed before and after the virtual midnight, resulting in a low temporal resolution." ...

625-626: This finding should be more elaborated. What is the exact relationship you found between coagulation growth and coating?

The slow ageing observed during the prolonged experiments in AIDA allowed noticing several side processes. Among other observations, we noted an abrupt flattening of the exponential growth of rBC cores (Figure S3a) at t24 during WL. At t24, rBC particles increased density, and fractal dimension approached spherical limits (Figure 5). Hence, we hypothesized that increase of fractal dimension, corresponding to a decrease in the active particle cross section, limited the collision efficiency reducing the coagulation rate. To the best of our knowledge, this process was simulated with numerical models but never really observed with in-situ measurements.

• The statement in the main text was simplified as:

... "The concurrent increase in fractal dimension may have reduced active surface area, limiting collisions and slowing coagulation. Despite prior studies (e.g. Schnaiter et al., 2003; Maricq, 2007), coagulation during the transition from external to internal mixing remains poorly characterized." ...

• More details are now given in Section E of the Supplementary material:

... "In WL, the diameter growth turned from an exponential function to a linear function in WL after the second aging step (t24), when fractal dimension and coating mass fraction increased (50°N; **Error! R** *eference source not found.*). Particle restructuring at the onset of coating may have reduced the active surface area, thereby limiting the collision frequency and coagulation rate. This hypothesis aligns with numerical simulations of Naumann (2003), suggesting that coating formation may reduce coagulation rates near the sources or in polluted regions. Although particle coagulation has been extensively studied in other settings (Schnaiter et al., 2003; Maricq, 2007), our finding raises the need of targeted experiments to quantify the change of coagulation rate of soot particles in the transition regime from external to internal mixing in ambient like conditions." ...

661: More explanation and/or reference is needed for Hill equation.

The Hill equation is a sigmoidal curve mostly used in biology/pharmacology to describe saturable processes or temporal appearance of events (Weiss, 1997; Goutelle et al., 2008). Hill equation was chosen to represent a process with well-defined minimum (bare rBC) and maximum (fully coated rBC) Rm_{coat} asymptotes. The upper limit was considered asymptotic due to the dominant role of coating mass over an almost negligible mass of rBC core, within the ARCTEx experimental time. Several tests performed with more generalized Sigmoid functions indicated that the Hill equation was able to capture the sudden increase of Rm_{Coat} , showing smaller residuals, at low elapsed time. The text was modified as:

... "Next, we derived the aging timescale necessary to reach the critical coating mass for each scenario. We applied a Hill equation (Weiss, 1997; Goutelle et al., 2008) to model the non-linear relationship between Rm_{coat} and BC age, forcing the upper (maximum Rm_{coat} observed during WL) and lower (initial Rm_{coat} from each scenario) boundaries along with a 95% confidence interval (**Error! Reference source not found.**a, dashed lines). C ompared to a more generic sigmoid equation, the Hill function allowed capturing the sudden increase of Rm_{coat} , in the early phase of the ageing." ...

Figure S1: Mention the instrument/technique used to obtain the results. For example, plot a) is an SMPS result I guess? Plot c) APM vs. SMPS and so on.

The caption of the figure now includes the name of the instruments.

Figure S1a: Any explanation why the size distributions of the low and high scenarios do differ? And SL and WL? Could the chamber temperature/pressure affect the emission after injection?

This comment was addressed above.

Figure S3: only 3 curves are visible in panel a). WH is missing, I guess.

The reviewer is correct - panel a of Figure S3 appears to show only three visible curves because SH and WH are almost perfectly overlapping. While this overlap may reduce the immediate readability of the figure, it also highlights the excellent repeatability of our experiments under conditions where no significant internal mixing occurred. We increase the resolution of the image to facilitate the reading.

References

Alexander, B., Sherwen, T., Holmes, C. D., Fisher, J. A., Chen, Q., Evans, M. J., and Kasibhatla, P.: Global inorganic nitrate production mechanisms: comparison of a global model with nitrate isotope observations, Atmospheric Chem. Phys., 20, 3859–3877, https://doi.org/10.5194/acp-20-3859-2020, 2020.

AMAP: AMAP Assessment 2021: Impacts of Short-lived Climate Forcers on Arctic Climate, Air Quality, and Human Health, 2021.

Backman, J., Schmeisser, L., and Asmi, E.: Asian Emissions Explain Much of the Arctic Black Carbon Events, Geophys. Res. Lett., 48, e2020GL091913, https://doi.org/10.1029/2020GL091913, 2021.

Daoudi, M., Schiffmann, P., Faccinetto, A., Frobert, A., and Desgroux, P.: Comprehensive characterization of particulate matter emissions produced by a liquid-fueled miniCAST burner, Aerosol Sci. Technol., 57, 872–889, https://doi.org/10.1080/02786826.2023.2228372, 2023.

Dulitz, K., Amedro, D., Dillon, T. J., Pozzer, A., and Crowley, J. N.: Temperature-(208–318 K) and pressure-(18–696 Torr) dependent rate coefficients for the reaction between OH and HNO₃, Atmospheric Chem. Phys., 18, 2381–2394, https://doi.org/10.5194/acp-18-2381-2018, 2018.

Ess, M. N. and and Vasilatou, K.: Characterization of a new miniCAST with diffusion flame and premixed flame options: Generation of particles with high EC content in the size range 30 nm to 200 nm, Aerosol Sci. Technol., 53, 29–44, https://doi.org/10.1080/02786826.2018.1536818, 2019.

Finlayson-Pitts, B. J. and Pitts, J.: Atmospheric chemistry. Fundamentals and experimental techniques, 1986.

Gao, R. S., Schwarz, J. P., Kelly, K. K., Fahey, D. W., Watts, L. A., Thompson, T. L., Spackman, J. R., Slowik, J. G., Cross, E. S., Han, J.-H., Davidovits, P., Onasch, T. B., and Worsnop, D. R.: A Novel Method for Estimating Light-Scattering Properties of Soot Aerosols Using a Modified Single-Particle Soot Photometer, Aerosol Sci. Technol., 41, 125–135, https://doi.org/10.1080/02786820601118398, 2007.

Goutelle, S., Maurin, M., Rougier, F., Barbaut, X., Bourguignon, L., Ducher, M., and Maire, P.: The Hill equation: a review of its capabilities in pharmacological modelling, Fundam. Clin. Pharmacol., 22, 633–648, https://doi.org/10.1111/j.1472-8206.2008.00633.x, 2008.

Henning, S., Ziese, M., Kiselev, A., Saathoff, H., Möhler, O., Mentel, T. F., Buchholz, A., Spindler, C., Michaud, V., Monier, M., Sellegri, K., and Stratmann, F.: Hygroscopic growth and droplet activation of soot particles: uncoated, succinic or sulfuric acid coated, Atmospheric Chem. Phys., 12, 4525–4537, https://doi.org/10.5194/acp-12-4525-2012, 2012.

Jathar, S. H., Gordon, T. D., Hennigan, C. J., Pye, H. O. T., Pouliot, G., Adams, P. J., Donahue, N. M., and Robinson, A. L.: Unspeciated organic emissions from combustion sources and their influence on the secondary organic aerosol budget in the United States, Proc. Natl. Acad. Sci., 111, 10473–10478, https://doi.org/10.1073/pnas.1323740111, 2014.

Khalizov, A. F., Xue, H., Wang, L., Zheng, J., and Zhang, R.: Enhanced Light Absorption and Scattering by Carbon Soot Aerosol Internally Mixed with Sulfuric Acid, J. Phys. Chem. A, 113, 1066–1074, https://doi.org/10.1021/jp807531n, 2009.

Laborde, M., Schnaiter, M., Linke, C., Saathoff, H., Naumann, K.-H., Möhler, O., Berlenz, S., Wagner, U., Taylor, J. W., Liu, D., Flynn, M., Allan, J. D., Coe, H., Heimerl, K., Dahlkötter, F., Weinzierl, B., Wollny, A. G., Zanatta, M., Cozic, J., Laj, P., Hitzenberger, R., Schwarz, J. P., and Gysel, M.: Single Particle Soot Photometer intercomparison at the AIDA chamber, Atmospheric Meas. Tech., 5, 3077–3097, https://doi.org/10.5194/amt-5-3077-2012, 2012.

Lefevre, G.: Impact of Organic Coating on Soot Angular and Spectral Scattering Properties, Env. Sci Technol, 9, 2019.

Levy II, H., Moxim, W. J., Klonecki, A. A., and Kasibhatla, P. S.: Simulated tropospheric NO : Its evaluation, global distribution and individual source contributions, J. Geophys. Res. Atmospheres, 104, 26279–26306, https://doi.org/10.1029/1999JD900442, 1999.

Li, W., Wang, Y., Yi, Z., Guo, B., Chen, W., Che, H., and Zhang, X.: Evaluation of MERRA-2 and CAMS reanalysis for black carbon aerosol in China, Environ. Pollut., 343, 123182, https://doi.org/10.1016/j.envpol.2023.123182, 2024.

Lim, C. Y., Hagan, D. H., Coggon, M. M., Koss, A. R., Sekimoto, K., de Gouw, J., Warneke, C., Cappa, C. D., and Kroll, J. H.: Secondary organic aerosol formation from the laboratory oxidation of biomass burning emissions, Atmospheric Chem. Phys., 19, 12797–12809, https://doi.org/10.5194/acp-19-12797-2019, 2019.

Mamakos, A., Khalek, I., Giannelli, R., and Spears, M.: Characterization of Combustion Aerosol Produced by a Mini-CAST and Treated in a Catalytic Stripper, Aerosol Sci. Technol., 47, 927–936, https://doi.org/10.1080/02786826.2013.802762, 2013.

Maricq, M. M.: Coagulation dynamics of fractal-like soot aggregates, J. Aerosol Sci., 38, 141–156, https://doi.org/10.1016/j.jaerosci.2006.11.004, 2007.

Mason, Y. C., Schoonraad, G.-L., Orasche, J., Bisig, C., Jakobi, G., Zimmermann, R., and Forbes, P. B. C.: Comparative sampling of gas phase volatile and semi-volatile organic fuel emissions from a combustion aerosol standard system, Environ. Technol. Innov., 19, 100945, https://doi.org/10.1016/j.eti.2020.100945, 2020.

Massling, A., Nielsen, I. E., Kristensen, D., Christensen, J. H., Sørensen, L. L., Jensen, B., Nguyen, Q. T., Nøjgaard, J. K., Glasius, M., and Skov, H.: Atmospheric black carbon and sulfate concentrations in Northeast Greenland, Atmos Chem Phys, 15, 9681–9692, https://doi.org/10.5194/acp-15-9681-2015, 2015.

Matsumi, Y., Comes, F. J., Hancock, G., Hofzumahaus, A., Hynes, A. J., Kawasaki, M., and Ravishankara, A. R.: Quantum yields for production of O(1D) in the ultraviolet photolysis of ozone: Recommendation based on evaluation of laboratory data, J. Geophys. Res. Atmospheres, 107, ACH 1-1-ACH 1-12, https://doi.org/10.1029/2001JD000510, 2002.

McCarty, J. L., Aalto, J., Paunu, V.-V., Arnold, S. R., Eckhardt, S., Klimont, Z., Fain, J. J., Evangeliou, N., Venäläinen, A., Tchebakova, N. M., Parfenova, E. I., Kupiainen, K., Soja, A. J., Huang, L., and Wilson, S.: Reviews and syntheses: Arctic fire regimes and emissions in the 21st century, Biogeosciences, 18, 5053–5083, https://doi.org/10.5194/bg-18-5053-2021, 2021.

Möhler, O., Büttner, S., Linke, C., Schnaiter, M., Saathoff, H., Stetzer, O., Wagner, R., Krämer, M., Mangold, A., Ebert, V., and Schurath, U.: Effect of sulfuric acid coating on heterogeneous ice nucleation by soot aerosol particles, J. Geophys. Res. Atmospheres, 110, https://doi.org/10.1029/2004JD005169, 2005.

Moore, R. H., Ziemba, L. D., Dutcher, D., Beyersdorf, A. J., Chan, K., Crumeyrolle, S., Raymond, T. M., Thornhill, K. L., Winstead, E. L., and Anderson, B. E.: Mapping the Operation of the Miniature Combustion Aerosol Standard (Mini-CAST) Soot Generator, Aerosol Sci. Technol., 48, 467–479, https://doi.org/10.1080/02786826.2014.890694, 2014.

Moschos, V., Schmale, J., Aas, W., Becagli, S., Calzolai, G., Eleftheriadis, K., Moffett, C. E., Schnelle-Kreis, J., Severi, M., Sharma, S., Skov, H., Vestenius, M., Zhang, W., Hakola, H., Hellén, H., Huang, L., Jaffrezo, J.-L., Massling, A., Nøjgaard, J. K., Petäjä, T., Popovicheva, O., Sheesley, R. J., Traversi, R., Yttri, K. E., Prévôt, A. S. H., Baltensperger, U., and Haddad, I. E.: Elucidating the present-day chemical composition, seasonality and source regions of climate-relevant aerosols across the Arctic land surface, Environ. Res. Lett., 17, 034032, https://doi.org/10.1088/1748-9326/ac444b, 2022.

Naumann, K.-H.: COSIMA—a computer program simulating the dynamics of fractal aerosols, J. Aerosol Sci., 34, 1371–1397, https://doi.org/10.1016/S0021-8502(03)00367-7, 2003.

Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N. M., Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y., Jimenez, J. L.,

Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J., McLaren, R., Mutzel, A., Osthoff, H. D., Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R. H., Shiraiwa, M., Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J., and Zaveri, R. A.: Nitrate radicals and biogenic volatile organic compounds: oxidation, mechanisms, and organic aerosol, Atmospheric Chem. Phys., 17, 2103–2162, https://doi.org/10.5194/acp-17-2103-2017, 2017.

Park, K., Kittelson, D. B., Zachariah, M. R., and McMurry, P. H.: Measurement of inherent material density of nanoparticle agglomerates, J. Nanoparticle Res., 6, 267–272, https://doi.org/10.1023/b:nano.0000034657.71309.e6, 2004.

Przybylak, R.: The Climate of the Arctic, Springer International Publishing, Cham, https://doi.org/10.1007/978-3-319-21696-6, 2016.

Saathoff, H., Naumann, K.-H., Möhler, O., Jonsson, Å. M., Hallquist, M., Kiendler-Scharr, A., Mentel, T. F., Tillmann, R., and Schurath, U.: Temperature dependence of yields of secondary organic aerosols from the ozonolysis of α -pinene and limonene, Atmospheric Chem. Phys., 9, 1551–1577, https://doi.org/10.5194/acp-9-1551-2009, 2009.

Schaap, M., Van Loon, M., Ten Brink, H. M., Dentener, F. J., and Builtjes, P. J. H.: Secondary inorganic aerosol simulations for Europe with special attention to nitrate, Atmospheric Chem. Phys., 4, 857–874, https://doi.org/10.5194/acp-4-857-2004, 2004.

Schnaiter, M., Horvath, H., Möhler, O., Naumann, K.-H., Saathoff, H., and Schöck, O. W.: UV-VIS-NIR spectral optical properties of soot and soot-containing aerosols, J. Aerosol Sci., 34, 1421–1444, https://doi.org/10.1016/S0021-8502(03)00361-6, 2003.

Srivastava, D., Vu, T. V., Tong, S., Shi, Z., and Harrison, R. M.: Formation of secondary organic aerosols from anthropogenic precursors in laboratory studies, Npj Clim. Atmospheric Sci., 5, 22, https://doi.org/10.1038/s41612-022-00238-6, 2022.

Tillmann, R., Saathoff, H., Brauers, T., Kiendler-Scharr, A., and Mentel, T. F.: Temperature dependence of the rate coefficient for the α -pinene reaction with ozone in the range between 243 K and 303 K, Phys. Chem. Chem. Phys., 11, 2323, https://doi.org/10.1039/b813407c, 2009.

Wagner, J., Ubele, A. A., Schenzinger, V., and Kreuter, A.: Extended Aerosol Optical Depth (AOD) time series analysis in an Alpine Valley: A Comparative Study from 2007 to 2023, Aerosol Measurements & amp; Instrumentation (AMI), https://doi.org/10.5194/ar-2023-20, 2024.

Weiss, J. N.: The Hill equation revisited: uses and misuses, FASEB J., 11, 835–841, https://doi.org/10.1096/fasebj.11.11.9285481, 1997.

Wittbom, C., Eriksson, A. C., Rissler, J., Carlsson, J. E., Roldin, P., Nordin, E. Z., Nilsson, P. T., Swietlicki, E., Pagels, J. H., and Svenningsson, B.: Cloud droplet activity changes of soot aerosol upon smog chamber ageing, Atmospheric Chem. Phys., 14, 9831–9854, https://doi.org/10.5194/acp-14-9831-2014, 2014.

Yuan, C., Zheng, J., Ma, Y., Jiang, Y., Li, Y., and Wang, Z.: Significant restructuring and light absorption enhancement of black carbon particles by ammonium nitrate coating, Environ. Pollut., 262, 114172, https://doi.org/10.1016/j.envpol.2020.114172, 2020.

Zanatta, M., Laj, P., Gysel, M., Baltensperger, U., Vratolis, S., Eleftheriadis, K., Kondo, Y., Dubuisson, P., Winiarek, V., Kazadzis, S., Tunved, P., and Jacobi, H.-W.: Effects of mixing state on optical and radiative properties of black carbon in the European Arctic, Atmospheric Chem. Phys., 18, 14037–14057, https://doi.org/10.5194/acp-18-14037-2018, 2018.