



1 A Cross-Correlation Based Method for Determining 2 Size-Resolved Particle Growth Rates

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11

12 **Abstract.** The particle growth rate (GR) is a key parameter in aerosol dynamics and plays a crucial
13 role in understanding atmospheric new particle formation and its effects. A fast, robust and
14 reproducible calculation of GRs from aerosol number-size distribution data remains a challenge. In this
15 study, we introduce a new method that we call the maximum correlation method for calculating particle
16 and ion GRs from number-size distributions. We employed this novel method to calculate GRs from
17 Hyytiälä, Finland using 14 years of ion and total particle size distribution data and compared our results
18 against previous studies that used conventional methods for calculating the GRs. We found that our
19 method compares well against the published data and reproduces the seasonal variability and size-
20 dependent trends in the GRs. The maximum correlation method enables fast and repeatable GR
21 calculations from large aerosol datasets, which facilitates the systematic incorporation of GR analysis
22 into new particle formation studies.

23

24 1 Introduction

25 New particle formation (NPF) is the process by which gas-phase molecules cluster and grow to form
26 stable aerosol particles in the atmosphere (e.g. Kulmala et al., 2013). NPF plays a critical role in the
27 climate system, as it serves as a major source of cloud condensation nuclei (CCN), which influence
28 cloud properties and radiative forcing in the atmosphere (Gordon et al., 2017; Merikanto et al., 2009;
29 Yu and Luo, 2009; Zhao et al., 2024).

30 Particle growth rate (GR), defined as the rate of change of particle diameter, is a key quantity
31 characterizing NPF (Kulmala et al., 2012). GR is an important parameter when determining the
32 probability that the freshly formed particles reach CCN sizes, especially when the particles are only a
33 few nanometers in diameter and easily scavenged by the pre-existing aerosol (Cai et al., 2022; e.g.
34 Kerminen et al., 2018; Stolzenburg et al., 2018; Yli-Juuti et al., 2011).



35 In ambient observations the GR is often determined from particle or ion concentration
36 measurements that are resolved by both size and time. The methods to calculate GR can be roughly
37 divided into two categories: number size distribution based methods and size channel based methods.
38 In both methods one determines a so-called apparent GR, which represents the observed growth of the
39 particle population at the measurement coordinates over the duration of the NPF event. The apparent
40 GR does not separate out any processes responsible for the observed particle population growth (e.g.
41 condensational growth, coagulation or size-dependent loss processes) and it does not determine GR of
42 any single particle. A spatial homogeneity assumption is often made, which states that the apparent GR
43 is equal to the average GR of the aerosol particle population.

44 In number size distribution based methods the GR is determined by analyzing the number size
45 distribution across time. In the so-called mode fitting method (Dal Maso et al., 2005; Kulmala et al.,
46 2012) a log-normal distribution is fitted to the growing particle mode at successive time steps and the
47 rate of change of the peak values is used at calculating the GR. Usually a line is fitted through the peaks
48 in a selected size range and the corresponding GR is reported as the slope of the fit. This method works
49 best when the growing particle mode is fully visible in the data, instead of continuing to sizes smaller
50 (or larger) than measured with the applied instrument. Paasonen et al. (2018) developed an automatic
51 method based on mode fitting that identifies growing particle modes and calculates their GRs.
52 However, the method is less reliable for determining GRs in the smallest particle sizes at the onset of
53 NPF events.

54 In size channel based methods, the GR is determined by analysing the concentration time
55 series across different size channels. A conventional approach is to fit a function to the increasing
56 concentrations associated with the growing particle mode and track specific features from the fitted
57 curves. In the maximum concentration method, a Gaussian function is used, and the peak positions
58 determine the GR (Hirsikko et al., 2005). In the appearance time method, a sigmoid function is fitted
59 to the leading edges of the rising concentrations, and the midpoint values are used to calculate the GR
60 (Lehtipalo et al., 2014). This method is suitable also for cases when NPF is sustained over longer time
61 periods, e.g. during chamber experiments (Dada et al., 2020), and therefore Gaussian function is not
62 suitable for locating the maximum concentration.

63 A related approach is to estimate the time lag between the rising concentrations in two separate
64 size channels (Riccobono et al., 2012; Sihto et al., 2006). The GR is calculated by dividing the size
65 difference with the time lag. The size channel based methods are suitable for narrow size range
66 instruments and when the particle mode is not fully visible in the data, which is usually the case when
67 the growing particles are sub-5 nm in size.

68 In both number size distribution and size channel based methods automating the fitting
69 procedure for NPF events is challenging. The range of data around the NPF event used for the fitting
70 is usually manually selected by the researcher. This makes GR calculation labor-intensive and
71 subjective. Additionally, fitting-based methods can be very sensitive to the chosen data range, leading
72 to variability in GR estimates and reduced reproducibility. In part due to these limitations, despite the



73 abundance of aerosol number-size distribution data, comprehensive datasets that report GRs in
74 different size ranges are scarce.

75 In this study we introduce the maximum correlation method, which is an automatic time lag
76 based method for GR calculation. Our objective is to apply this method to a large aerosol particle
77 dataset collected from Hyytiälä, Finland, and compare the results with previously published size-
78 resolved GR datasets from the same location.

79 2 Methods

80 2.1 Maximum correlation method

81 When no other processes significantly influence the particle size distribution, particle growth during or
82 after an NPF event leads to an increase in particle number concentration that is observed earlier in the
83 smaller size channels and later in the larger size channels. The task is to find a way to calculate the
84 time displacement between the concentration rise for the different size channels and use it in the
85 calculation of the GR.

86 We assume that a good condition for finding the optimal time displacement is when the
87 concentrations in two size channels are maximally correlated. Next we will outline how this idea is
88 used to calculate particle GRs from number-size distribution data.

- 89
- 90 1. Let us choose particle diameters d_1 and d_2 ($d_1 < d_2$) from the number size distribution
 - 91 2. From the number-size distribution interpolate the concentration time series $N_1(t)$ and $N_2(t)$
92 that correspond to the size channels represented by the chosen diameters.
 - 93 a. It is possible to select an arbitrary time window for the time t . Here we chose one
94 day from midnight to midnight local time since NPF tends to follow a diurnal cycle
95 in most environments.
 - 96 b. Here we required that no more than 5% of the concentration values in the size range
97 of interest were missing, otherwise the day was categorized as bad data. Days when
98 the instrument was not measuring were also categorized as bad data.
 - 99 3. Calculate normalized cross-correlation for $N_1(t)$ and $N_2(t)$:

100
$$R_{N_1 N_2}(\tau) = \frac{1}{M(\tau)} \sum_t \frac{(N_1(t - \tau) - \underline{N_1})(N_2(t) - \underline{N_2})}{\sigma_1 \sigma_2}$$

- 101 a. $N_1(t)$ and $N_2(t)$ are normalized by subtracting their means $\underline{N_1}$ and $\underline{N_2}$ and dividing
102 by their standard deviations σ_1 and σ_2 . This makes the method less sensitive to
103 baseline concentration levels or the differences in concentration amplitude.
- 104 b. τ is the time displacement. Unless otherwise stated we varied τ from -23 h to 23 h at
105 increments of 1s. We refer to the limiting displacement value as τ_{lim} .



- 106 c. $M(\tau)$ is the number of overlapping data points between time series $N_1(t)$ and $N_2(t)$
107 for a given τ . We divide the sum by $M(\tau)$ in order to reduce the effect of overlap on
108 the cross-correlation since more overlapping points lead to a higher cross-correlation.
109 d. In this study the channel concentrations $N_1(t)$ and $N_2(t)$ were smoothed using a 3
110 hour rolling mean. The concentration increases caused by regional NPF are expected
111 to last several hours and the smoothing window width is optimized for preserving
112 these peaks while removing higher frequency fluctuations that could cause erroneous
113 values for the cross correlation. The chosen smoothing window width depends on
114 the noise level of the data and time scale of the particle growth process under study.
115 4. Find the time displacement at maximum correlation $\tau_{max} = \arg \max(R_{N_1 N_2}(\tau))$
116 a. Return a missing value for results where $\tau_{max} \leq 0$ s and for the limiting case $\tau_{max} =$
117 τ_{lim}
118 5. Calculate the growth rate as $GR_{d_1-d_2} = \frac{d_2-d_1}{\tau_{max}}$

119
120 Figure 1 illustrates how the maximum correlation method is used to calculate particle GR
121 from the number size distribution. In this case we calculated the GR of negative ions from 2 nm to 3
122 nm. The maximum concentration method was also applied to the case for comparison.

123 If the diameters d_1 and d_2 are too widely spaced, the correlation between the size channels is
124 influenced by unrelated atmospheric processes. This is why the size range may have to be divided into
125 smaller size increments and then the maximum correlation method can be applied to each smaller size
126 range separately. In order to calculate GR for the whole size range one should add the maximum
127 correlation time displacements $\tau_{max,i}$ for each smaller size range numbered by i

128
$$\tau_{max} = \sum_{i=1}^n \tau_{max,i}$$

129 and use the resulting τ_{max} in the final GR calculation. Here we used a condition that if for any of the
130 smaller size ranges $\tau_{max,i} \leq 0$ or $\tau_{max,i} = \tau_{lim}$ then the GR for the whole size range would also be a
131 missing value.

132 2.2 Hyytiälä dataset

133 The maximum correlation method can be used on individual days when a growing particle mode is
134 present or to determine the distribution of GRs across multiple days in an automated fashion, as
135 demonstrated in this study.

136 We tested the method on an ion and total particle number size distribution (INSD/PNSD)
137 dataset from the SMEAR II station. SMEAR II station is located in Hyytiälä, Finland in a rural boreal
138 forest environment (24.30E, 61.85N, 180m; Hari and Kulmala, 2005). The dataset is approximately 14
139 years long (Feb 2010-Dec 2024). No NPF event classification was done on the days in the dataset prior
140 to the GR analysis. The INSDs and PNSDs were measured by a Neutral cluster and Air Ion



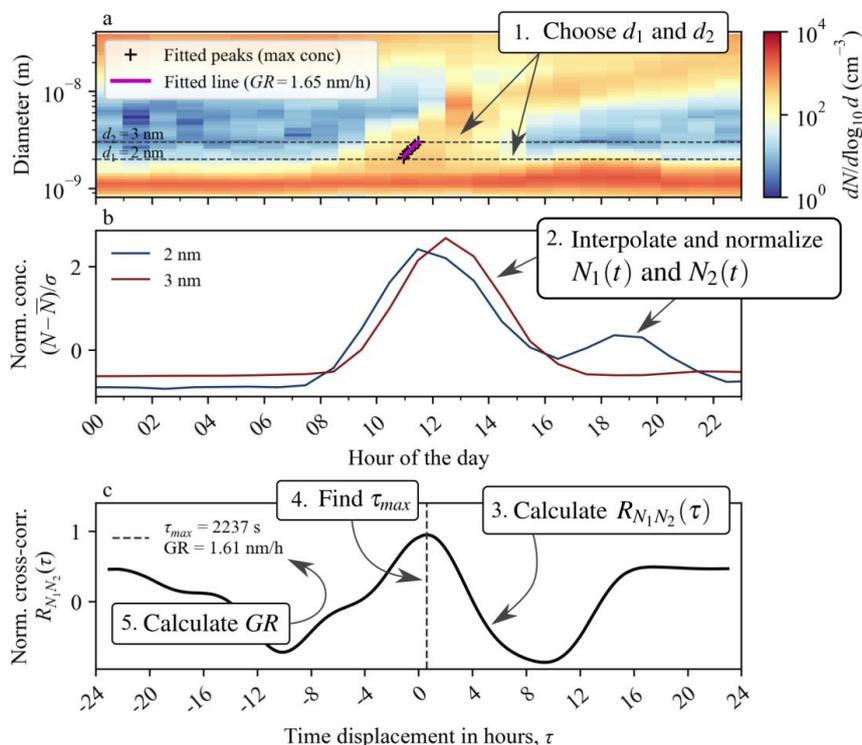
141 Spectrometer (NAIS, Ariel Ltd.; Mirme and Mirme, 2013) and Differential Mobility Particle Sizer
 142 (DMPS; Aalto et al., 2001).

143 The NAIS measured the number size distribution of air ions and total particles in the mobility
 144 equivalent diameter range of approximately 0.8–40 nm and 2.5–40 nm respectively. The DMPS system
 145 measured the number size distribution of total particles in the mobility equivalent diameter range of 3–
 146 1000 nm. The number size distributions were averaged to 1 hour time resolution.

147 GRs were calculated in three size ranges: 2–3 nm, 3–7 nm and 7–20 nm. To ensure each size
 148 range was separately subdivided into smaller, equally spaced logarithmic bins: two bins for 2–3 nm,
 149 three for 3–7 nm, and four for 7–20 nm. The concentrations at each diameter were interpolated from
 150 the number size distribution (number concentrations normalized by logarithm of bin width,
 151 $dN/d \log_{10} d$) measured by the NAIS or the DMPS.

152 We chose the above instruments and size ranges since comparable GR datasets already exist
 153 from Hyytiälä (Gonzalez Carracedo et al., 2022; Hirsikko et al., 2005; Manninen et al., 2009, 2010;
 154 Yli-Juuti et al., 2011).

155



156

157 **Figure 1:** Example case from 15 March 2010 illustrating the calculation of GR for 2–3 nm negative ions
 158 using the maximum correlation method. (a) Negative INSD measured by the NAIS with peak diameters
 159 fitted using the maximum concentration method and the GR found by fitting a line to them, (b)



160 **normalized particle concentrations in the two size channels (rolling mean not applied yet), and (c) cross-**
161 **correlation between size channels, with the maximum correlation and corresponding growth rate shown.**
162 **For the purpose of illustrating the method the size range was not divided into smaller increments.**

163 **3 Results and discussion**

164 The GR distributions obtained from the 14-year dataset in Hyytiälä using the maximum correlation
165 method are shown in Figure 2. The distributions appear to consist of two distinct parts that we call the
166 background and the signal.

167 We separated the signal distribution from the background by identifying a local minimum, β ,
168 between them. The likely reason for the background distribution at very low GRs is that the small
169 number of overlapping points at large τ can produce relatively high cross-correlations, especially with
170 the normalization. In the absence of particle growth the maximum cross correlation tends to occur when
171 the early morning of the smaller size channel is aligned with the late evening of the larger size channel,
172 leading to a distribution of low GR values.

173 To further illustrate this we prepared Figures A1-A6, which show examples of negative INSDs
174 on days that are from the signal distribution and on days that are from the background distribution for
175 each analyzed size class. The majority of the signal days, especially looking at the larger size ranges,
176 show features typical of NPF days, while the background days contain mostly days that would be
177 classified as non-event or undefined days (Dal Maso et al., 2005). This supports excluding them from
178 the GR analysis.

179 The number of large outliers is relatively small in all GR distributions, which suggests that
180 most local particle plumes that might cause very high GRs are filtered away by the requirement that
181 $\tau_{max} > 0$ s. We chose to ignore the large outliers due to their small number. In other environments the
182 number of outliers may be higher due to increased local particle emissions.

183 As a result of separating out the background distribution, the remaining signal distribution
184 contains the GR values from the days with true particle growth. If no signal distribution is visible it is
185 likely because particle growth is not happening or not detected by the measurement. However, further
186 testing needs to be done by using data from environments that have extremely slow (e.g. Arctic sites)
187 or fast (e.g. some coastal sites) particle growth to see if a similar separation into background and growth
188 distributions occurs.

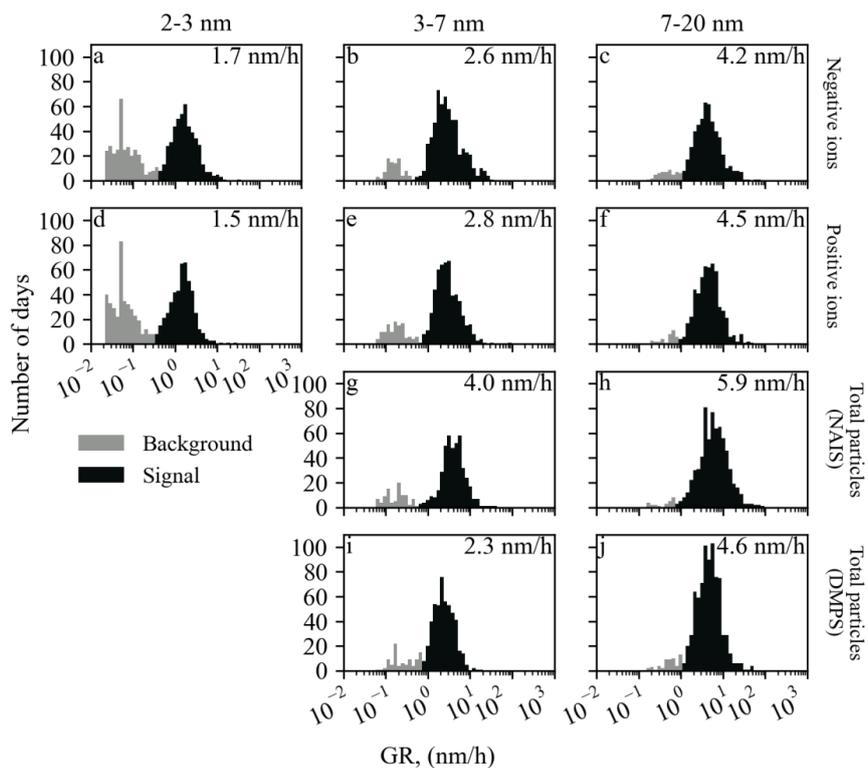
189 The effect of dividing the 3-7 nm size range equally into n smaller increments of logarithmic
190 width $\Delta \log_{10} d$ is illustrated in Figure 3a. Dividing the size range into two smaller increments
191 already dramatically improved the signal-to-noise ratio in the GR distribution. Further increasing the
192 number of increments removed high GR days from the distribution. This is because if there are more
193 smaller size ranges the probability of getting $\tau_{max,i} = 0$ s, especially on high GR days, is increased and
194 in this case a missing value is returned for the whole GR. In order to maximize the signal-to-noise ratio
195 and minimize the number of discarded high GR days, we aimed for a log difference of approximately
196 0.1 for the size increments when dividing the size range.



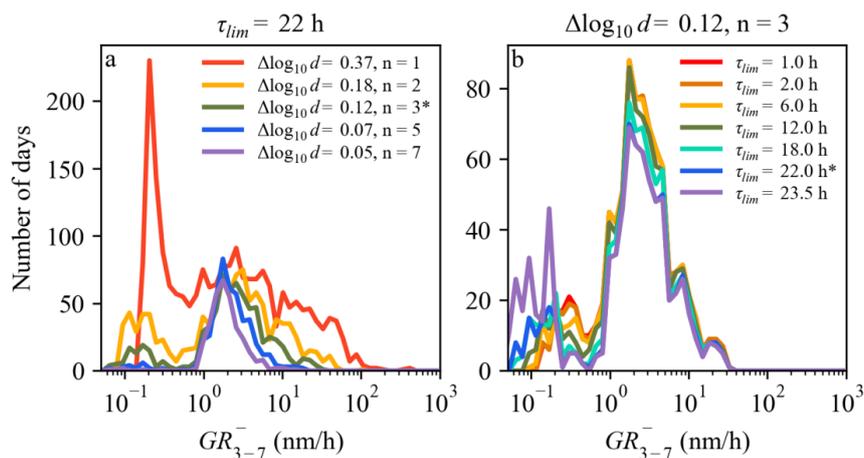
197 The effect of using different time displacement ranges is shown in Figure 3b. The choice of
198 τ_{lim} does not significantly influence the shape of the signal distribution, however using larger τ_{lim}
199 slightly reduces the number of days in the signal distribution. This is likely because for larger τ range
200 there are more chances to find a τ_{max} that results in a missing value. The background distribution is
201 pushed towards the smaller GRs as τ_{lim} is increased, which is in line with the idea that when there is
202 no particle growth the τ_{max} tends to occur with less overlapping data points. The shape of the
203 background distribution does not significantly change until $\tau_{lim} = 23.8$ h at which point the number
204 of days in the background distribution is increased. This is likely because higher cross correlations
205 occur when the number of overlapping data points becomes very small. The local minimum between
206 the background and the signal is not much affected by the choice of $\Delta \log_{10} d$ or τ_{lim} .

207 In addition there were days when GR could not be calculated due to the data quality not being
208 good enough (bad data) or due to $\tau_{max} \leq 0$ s (negative or infinite GR) or $\tau_{max} = \tau_{lim}$ (limiting case).
209 The number of days in each category is shown in Figure 4. The relatively large number of bad data is
210 explained by our rather strict criteria for usable data (<5% missing data in the size range of interest).
211 Additionally, days when the instrument was not measuring were categorized as bad data. By
212 comparison Yli-Juuti et al. (2011) was able to calculate $GR_{1.5-3}$ on 5%, GR_{3-7} on 4% and GR_{7-20} on
213 2% of all the days using the conventional methods, which are clearly lower percentages compared to
214 our method.

215 We investigated how well the GRs calculated using the maximum correlation method compare
216 with GRs calculated using the maximum concentration method. We randomly sampled 25 days from
217 the 3-7 nm negative ion GRs and calculated the GR in the corresponding size range using the maximum
218 concentration method. Figure 5 shows strong positive correlation between the two sets of GRs ($\rho =$
219 0.9). The number size distributions on these days are shown in Figure A3.
220



221
222 **Figure 2: The GR distributions are separated into background and signal parts. The median GR from the**
223 **signal distribution is shown in the upper right corner.**
224



225

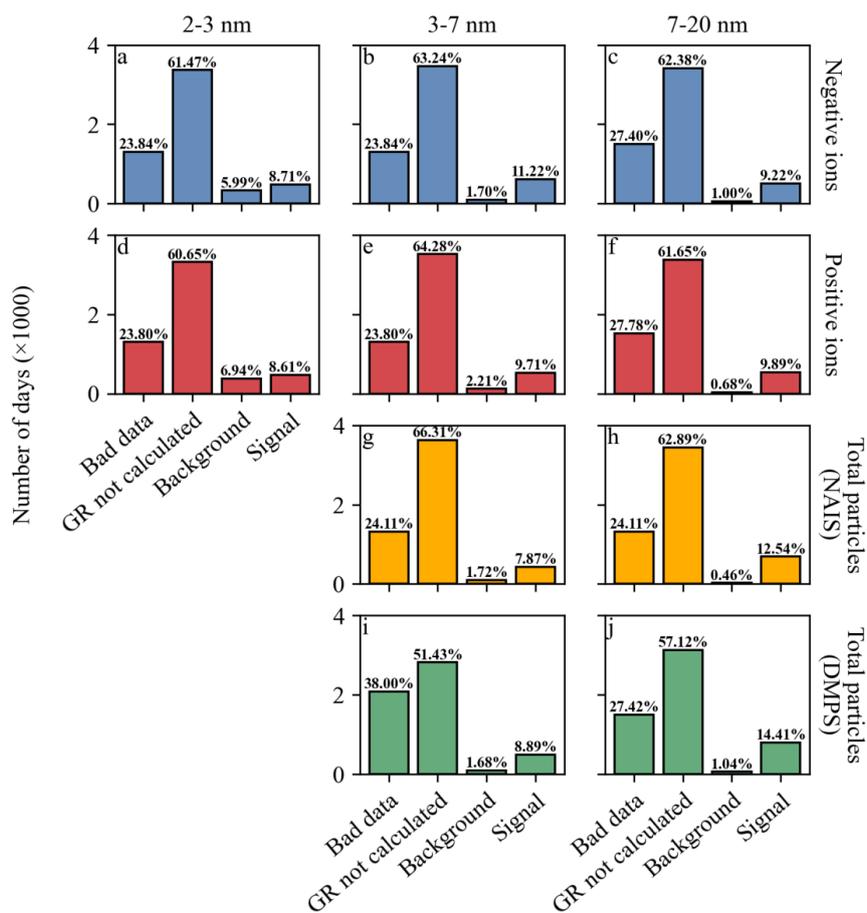
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Figure 3: The effect of (a) dividing the size range into n smaller size increments of logarithmic width $\Delta \log_{10} d$ and (b) using different sized time displacement windows $\tau \in [-\tau_{lim}, \tau_{lim}]$. The GR distributions were calculated from negative ions. The values used in this study are marked with an asterisk.



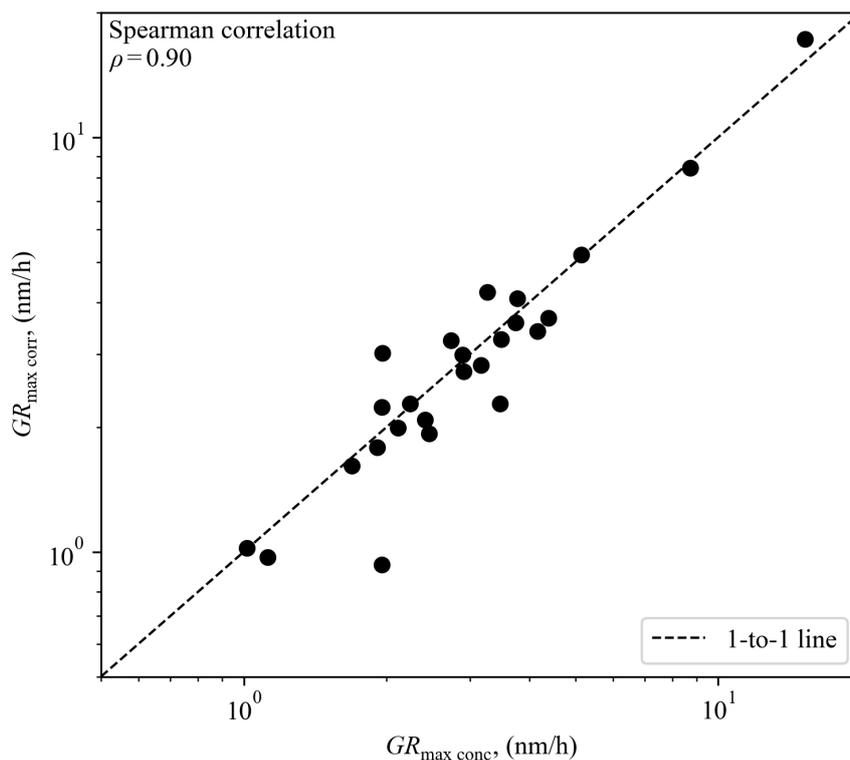
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231 **Figure 4: The number of days categorized as bad data, GR not calculated, background and**

232 **signal. Percentage of total days is shown on top of each bar.**

233

234



235
236 **Figure 5: Comparison of GRs calculated using the maximum concentration and maximum correlation**
237 **methods. The days were randomly selected from the 3-7 nm negative ion GRs. The number size**
238 **distributions on these days are presented in Figure A3.**

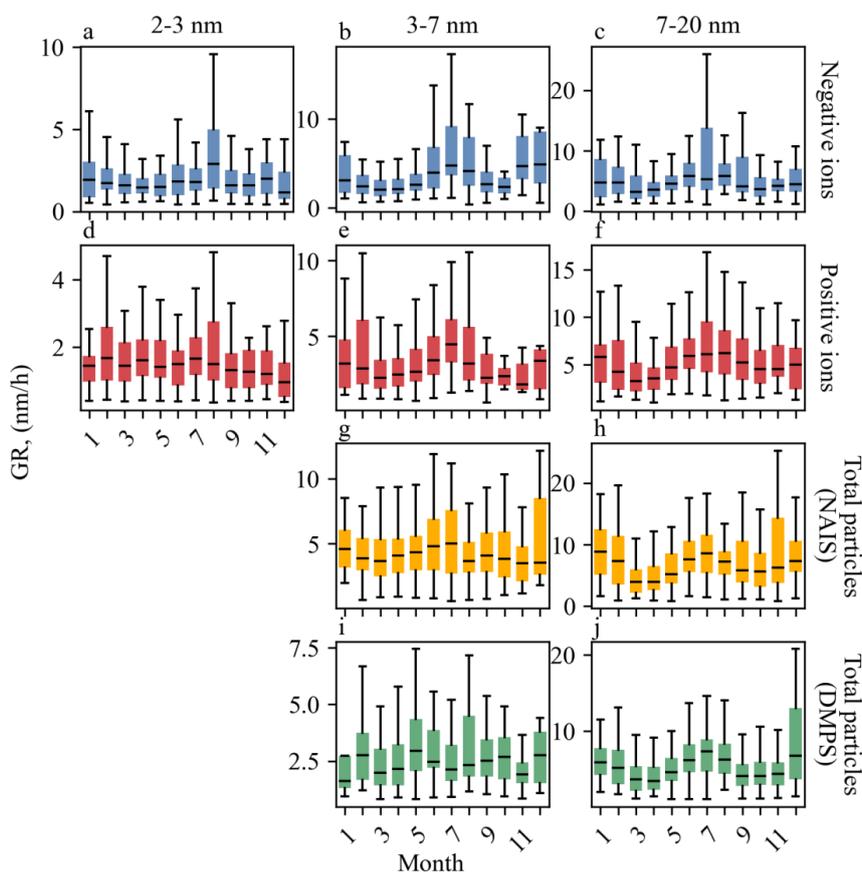
239
240 The median GRs showed an increasing trend with particle size, a pattern commonly observed across
241 various environments (Kerminen et al., 2018). The median GRs found for the positive and negative
242 ions are similar. This is also supported by previous observations, although in the smallest size range
243 variation exists across studies (Gonzalez Carracedo et al., 2022; Hirsikko et al., 2005; Manninen et al.,
244 2009; Yli-Juuti et al., 2011).

245 The median GRs calculated from the total particles measured by the NAIS are higher than the
246 ones calculated from the ions, while the median GRs from DMPS agree better with the ion GRs. The
247 other studies from Hyytiälä calculated the total particle GRs from DMPS or DMA-train data and found
248 that they were overall similar to the ion GRs (Hirsikko et al., 2005; Manninen et al., 2009; Yli-Juuti et
249 al., 2011), except in the sub-3 nm size range (Gonzalez Carracedo et al., 2022). This suggests that the
250 GRs may be overestimated when calculated from the NAIS total particle mode.

251 Figure 6 illustrates the seasonal variation in the GRs. In 3-7 nm and 7-20 nm we observed a
252 maximum in GRs during the summer months. Between 2-3 nm the ion GRs stay more constant



253 throughout the year, although a summer maximum is seen in negative ions. This seasonal behavior of
254 GRs is a common characteristic observed in Hyytiälä (Nieminen et al., 2018; Paasonen et al., 2018;
255 Yli-Juuti et al., 2011)
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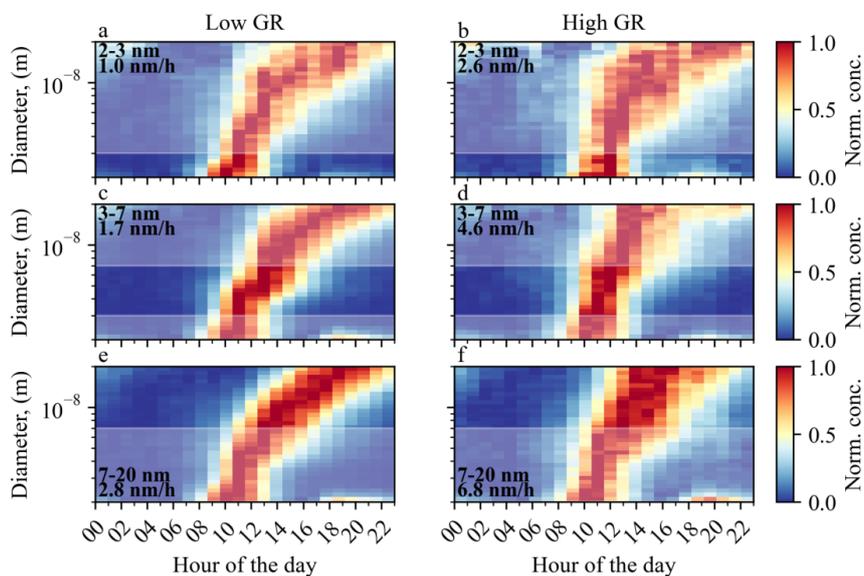


257
258 **Figure 6: Seasonal variation of the GRs.**
259

260 Next we divided the GR data in each size range into two parts: days where the GR was below the
261 median (low GRs) and days where the GR was above or equal to the median (high GRs). Figure 7
262 shows the median diurnal cycles of the normalized negative INSDs for low and high GR days. Each
263 diurnal cycle shows a growing negative ion mode that is first detected in the smallest sizes around
264 midday. This is a defining feature of NPF. The slope of the growing ion mode in the size range of
265 interest increases when going from low to high GRs.



266 This shows that the GRs detected by our method are related to particle growth during NPF.
267 Furthermore the data indicates that on low GR days the ions indeed grow slower compared to high GR
268 days.
269



270
271 **Figure 7: Median diurnal normalized negative INSDs in the different size ranges for low and high GRs.**
272 **The size range where the GR was calculated is shown in each subplot and also indicated by the fully**
273 **transparent area on the plot. The median GR is shown below the size ranges.**
274

275 Finally, we compared our results with published GR data from Hyytiälä. Yli-Juuti et al. (2011)
276 employed different instruments and methods to comprehensively study ion and total particle GRs
277 during NPF events in Hyytiälä between 2003-2009. The instruments included an Air Ion Spectrometer
278 (AIS, Mirme et al., 2007), which is an earlier version of the NAIS used in this study, a DMPS and a
279 Balanced Scanning Mobility Analyzer (BSMA). The GRs were calculated from all the instruments
280 using the maximum concentration method, while the mode fitting method was only used on the DMPS
281 data. It should be noted that the correlation between the methods was found to be rather good ($R=0.72$),
282 but the maximum concentration method typically showed slightly higher values (median difference 1.1
283 nm/h). For the smallest size range (1.5-3 nm) only ion data was used and for the larger size ranges (3-
284 7 nm and 7-20 nm) both ion and total particle data were used. The reported final median GRs in each
285 size class were averages of all the GRs calculated from the different instruments and methods (median
286 GR was calculated from each method/instrument pair and mean was taken over the resulting median
287 GRs).

288 Hirsikko et al. (2005) used data from the same instruments as Yli-Juuti et al. (2011) between
289 2003-2004, however only the maximum concentration method was used to calculate the GRs and the

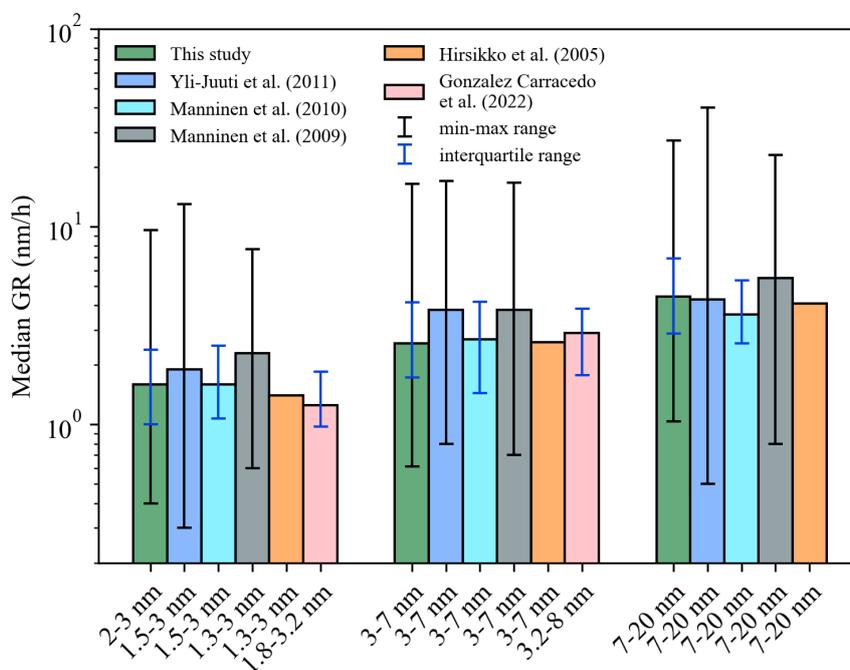


290 smallest size range was 1.3-3 nm. Otherwise the median GR was calculated similarly to Yli-Juuti et al.
291 (2011). Manninen et al. (2009) and (2010) used the same methodology as Hirsikko et al. (2005) for
292 March-June 2007 and March 2008-May 2009 respectively.

293 Gonzalez Carracedo et al. (2022) used the appearance time method and the maximum
294 concentration method to calculate GRs in two different size ranges from ions measured by NAIS and
295 total particles measured by DMA train. The authors found good agreement between the two methods.
296 The data was measured in Hyytiälä between March-September 2020. We chose only the ion GRs to
297 calculate the median GR in the smallest size range (1.8-3.2 nm) and both ion and total particle GRs to
298 calculate the median GR in the larger size range (3.2-8 nm).

299 For comparing our results we calculated the median 2-3 nm GR from ions and the median 3-
300 7 nm and 7-20 nm GRs using ion and DMPS data. We used the 0 and 99 percentile values as our min
301 and max values. The results of the comparison are shown in Figure 8. Overall the median GRs in all
302 the size ranges compare well across all studies. Also the min-max ranges and the interquartile ranges
303 of the GRs (if reported) compare well across the studies.

304



305
306 **Figure 8: Comparison of GRs obtained in this study against published data from Hyytiälä. The bar height**
307 **is the median GR in each size range and the error bars show different ranges of GR values reported in the**
308 **studies.**



309 4 Conclusions

310 We presented a new method, called the maximum correlation method, for determining size-resolved
311 particle GRs from aerosol number-size distribution data. The method was tested and validated in the
312 sub-20 nm size range, showing its suitability for NPF studies. The proposed method is a time lag based
313 method, which uses cross-correlation to determine the optimum time lag used for the GR calculation.
314 We applied the method to approximately 14 years of ion and total particle data from the SMEAR II
315 station in Hyytiälä, Finland and calculated GRs in three different size ranges: 2-3 nm, 3-7 nm and 7-
316 20 nm.

317 The obtained median GRs, along with the variability measured by the interquartile range and
318 the minimum–maximum range, were consistent with findings from previous studies. Our results
319 reproduced the previously observed GR maximum during summer and an increasing GR as a function
320 of particle size. On a subset of days, the GRs estimated by our method showed a positive correlation
321 with those obtained using the maximum concentration method. Median diurnal negative INSDs on GR
322 days and negative INSDs on example days showed that the growth of negative ions was predominantly
323 due to NPF events.

324 One should keep in mind that the cross-correlation method gives the apparent GR of the
325 particle population, which is roughly equal to the condensational GR under non-polluted environments
326 with roughly homogenous sources of condensable vapors in surrounding areas. In polluted
327 environments the effect of coagulation on the GR should be taken into account (Cai et al., 2021) and
328 the heterogeneities in condensable vapor concentrations upwind the observation site alter strongly the
329 apparent GR (Hakala et al., 2023; Kivekäs et al., 2016).

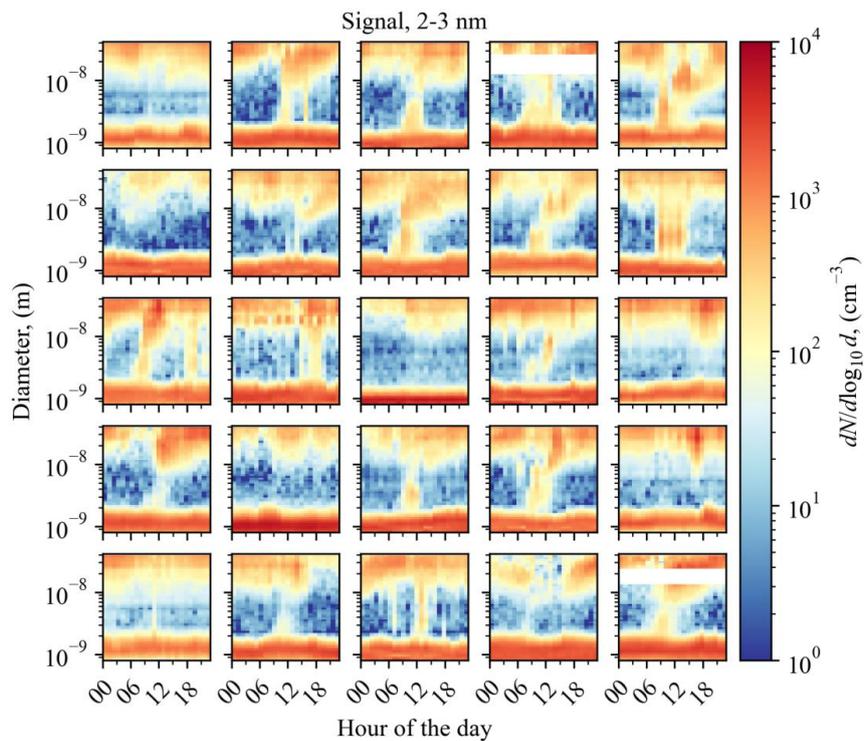
330 In the future it is important to test our method using aerosol data from other types of
331 environments. For example in places with local emission sources there may be multiple particle growth
332 events during a single day. In this case running the maximum correlation method using different time
333 windows may be necessary in order to separate out the different growth processes. On the other hand,
334 the suitability for detecting very low or very high GRs; as well as weak NPF should be further
335 investigated.

336 The maximum correlation method allows one to efficiently and systematically calculate GRs
337 from ion and total particle number size distributions. The proposed method is readily applied to large
338 collections of data, which facilitates the GR analysis from new and existing aerosol datasets. When
339 combined with statistical NPF classification methods, such as the nanoparticle ranking method (Aliaga
340 et al., 2023), it could replace the conventional labor-intensive NPF analysis, especially when NPF is
341 the dominant source of particles in the size range.

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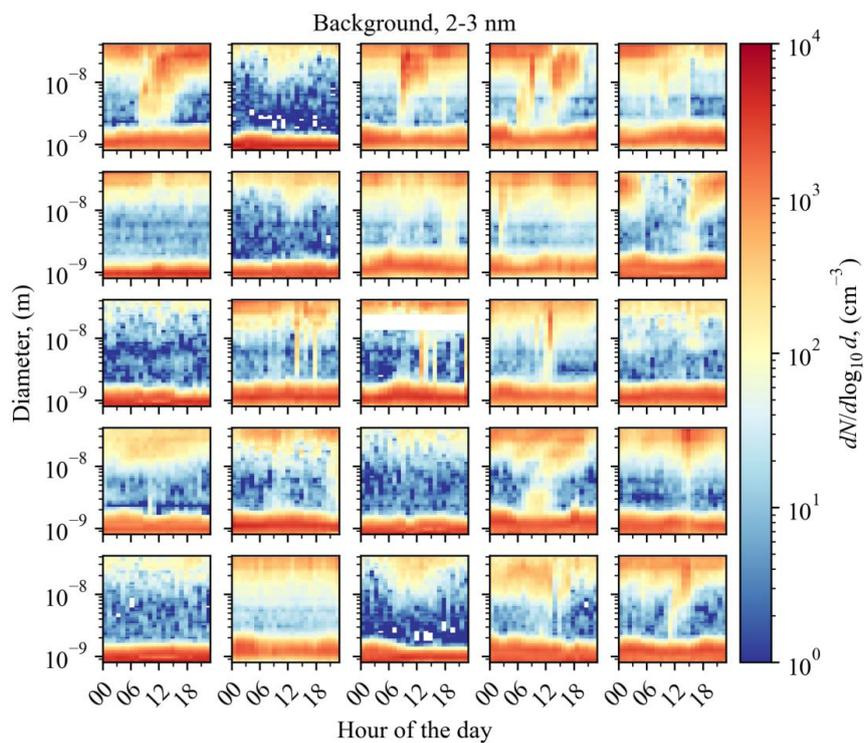
343 Appendix A



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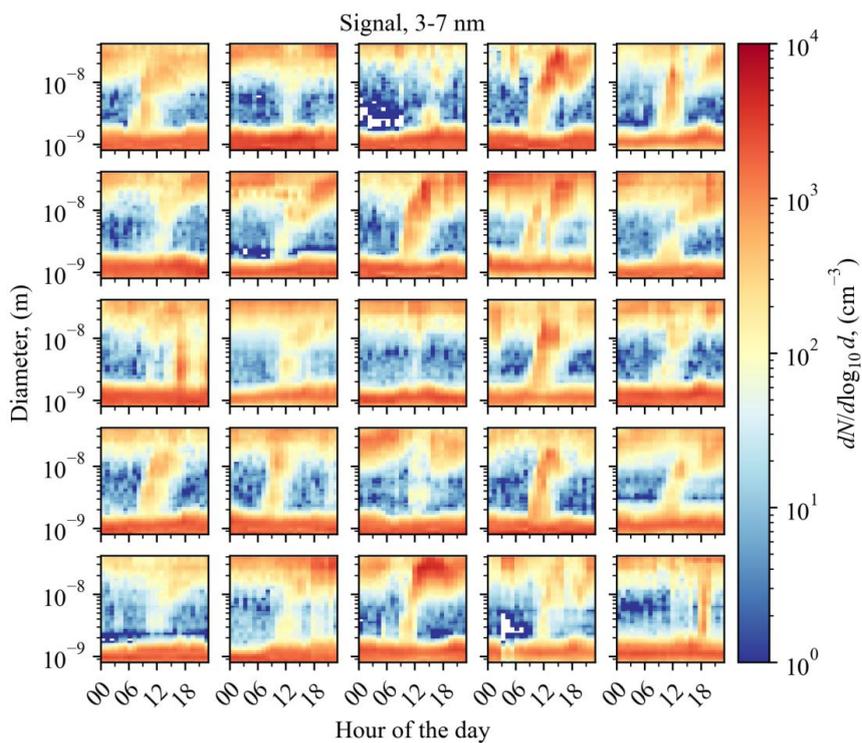
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Figure A1: Example negative INSDs on days sampled from the GR_{2-3}^- distribution's signal part.



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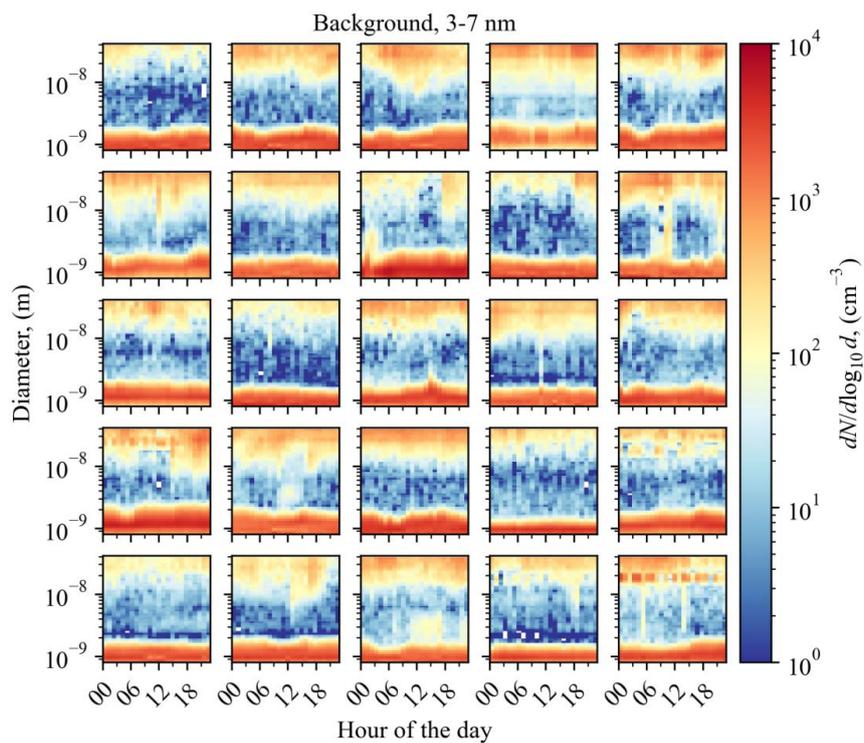
Figure A2: Example negative INSDs on days sampled from the GR_{2-3} distribution's background part.



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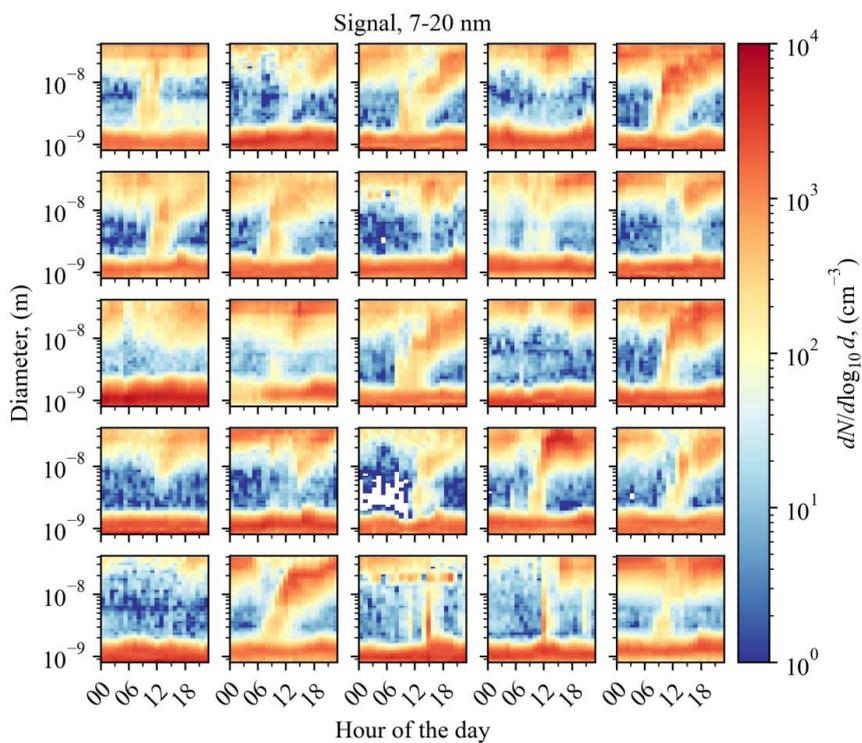
Figure A3: Example negative INSDs on days sampled from the GR_{3-7}^- distribution's signal part.



350

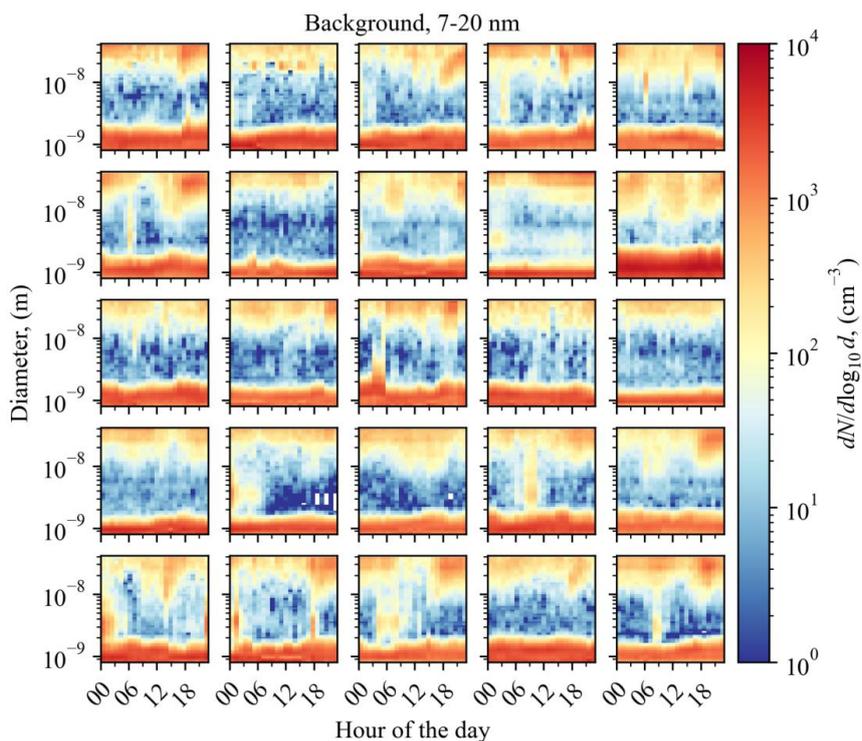
351

Figure A4: Example negative INSDs on days sampled from the GR_{3-7}^- distribution's background part.



352
353

Figure A5: Example negative INSDs on days sampled from the GR_{7-20} distribution's signal part.



354

355 **Figure A6:** Example negative INSDs on days sampled from the GR_{7-20} distribution's background part.

356

357 **Data availability.** The NAIS dataset is available at <https://doi.org/10.5281/zenodo.15648699>
358 (Lampilahti et al., 2025). The DMPS dataset can be accessed through the SmartSMEAR data portal at
359 <https://smear.avaa.csc.fi/>.

360

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362 contributed to refining the methodology. JL wrote the manuscript with input from all co-authors.

363

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366

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376 References

- 377 Aalto, P., Hämeri, K., Becker, E., Weber, R., Salm, J., Mäkelä, J. M., Hoell, C., O'Dowd, C. D.,
378 Hansson, H.-C., Väkevä, M., Koponen, I. K., Buzorius, G., and Kulmala, M.: Physical
379 characterization of aerosol particles during nucleation events, *Tellus B*, 53, 344–358,
380 <https://doi.org/10.3402/tellusb.v53i4.17127>, 2001.
- 381 Aliaga, D., Tuovinen, S., Zhang, T., Lampilahti, J., Li, X., Ahonen, L., Kokkonen, T., Nieminen, T.,
382 Hakala, S., Paasonen, P., Bianchi, F., Worsnop, D., Kerminen, V.-M., and Kulmala, M.: Nanoparticle
383 ranking analysis: determining new particle formation (NPF) event occurrence and intensity based on
384 the concentration spectrum of formed (sub-5 nm) particles, *Aerosol Research*, 1, 81–92,
385 <https://doi.org/10.5194/ar-1-81-2023>, 2023.
- 386 Cai, R., Li, C., He, X.-C., Deng, C., Lu, Y., Yin, R., Yan, C., Wang, L., Jiang, J., Kulmala, M., and
387 Kangasluoma, J.: Impacts of coagulation on the appearance time method for new particle growth rate
388 evaluation and their corrections, *Atmospheric Chemistry and Physics*, 21, 2287–2304,
389 <https://doi.org/10.5194/acp-21-2287-2021>, 2021.
- 390 Cai, R., Deng, C., Stolzenburg, D., Li, C., Guo, J., Kerminen, V.-M., Jiang, J., Kulmala, M., and
391 Kangasluoma, J.: Survival probability of new atmospheric particles: closure between theory and
392 measurements from 1.4 to 100 nm, *Atmospheric Chemistry and Physics*, 22, 14571–14587,
393 <https://doi.org/10.5194/acp-22-14571-2022>, 2022.
- 394 Dada, L., Lehtipalo, K., Kontkanen, J., Nieminen, T., Baalbaki, R., Ahonen, L., Duplissy, J., Yan, C.,
395 Chu, B., Petäjä, T., Lehtinen, K., Kerminen, V.-M., Kulmala, M., and Kangasluoma, J.: Formation
396 and growth of sub-3-nm aerosol particles in experimental chambers, *Nat Protoc*, 15, 1013–1040,
397 <https://doi.org/10.1038/s41596-019-0274-z>, 2020.
- 398 Dal Maso, M., Kulmala, M., Riipinen, I., Wagner, R., Hussein, T., Aalto, P. P., and Lehtinen, K. E.:
399 Formation and growth of fresh atmospheric aerosols: eight years of aerosol size distribution data
400 from SMEAR II, Hyytiälä, Finland, *Boreal Environ. Res.*, 10, 323, 2005.
- 401 Gonzalez Carracedo, L., Lehtipalo, K., Ahonen, L. R., Sarnela, N., Holm, S., Kangasluoma, J.,
402 Kulmala, M., Winkler, P. M., and Stolzenburg, D.: On the relation between apparent ion and total
403 particle growth rates in the boreal forest and related chamber experiments, *Atmospheric Chemistry
404 and Physics*, 22, 13153–13166, <https://doi.org/10.5194/acp-22-13153-2022>, 2022.



- 405 Gordon, H., Kirkby, J., Baltensperger, U., Bianchi, F., Breitenlechner, M., Curtius, J., Dias, A.,
406 Dommen, J., Donahue, N. M., Dunne, E. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Frege, C.,
407 Fuchs, C., Hansel, A., Hoyle, C. R., Kulmala, M., Kürten, A., Lehtipalo, K., Makhmutov, V.,
408 Molteni, U., Rissanen, M. P., Stozkhov, Y., Tröstl, J., Tsagkogeorgas, G., Wagner, R., Williamson,
409 C., Wimmer, D., Winkler, P. M., Yan, C., and Carslaw, K. S.: Causes and importance of new particle
410 formation in the present-day and preindustrial atmospheres, *Journal of Geophysical Research:*
411 *Atmospheres*, 122, 8739–8760, <https://doi.org/10.1002/2017JD026844>, 2017.
- 412 Hakala, S., Vakkari, V., Lihavainen, H., Hyvärinen, A.-P., Neitola, K., Kontkanen, J., Kerminen, V.-
413 M., Kulmala, M., Petäjä, T., Hussein, T., Khoder, M. I., Alghamdi, M. A., and Paasonen, P.:
414 Explaining apparent particle shrinkage related to new particle formation events in western Saudi
415 Arabia does not require evaporation, *Atmospheric Chemistry and Physics*, 23, 9287–9321,
416 <https://doi.org/10.5194/acp-23-9287-2023>, 2023.
- 417 Hari, P. and Kulmala, M.: Station for measuring ecosystem-atmosphere relations (SMEAR II), *Boreal*
418 *Environ. Res.*, 10, 315–322, 2005.
- 419 Hirsikko, A., Laakso, L., Hörrak, U., Aalto, P. P., Kerminen, V.-M., and Kulmala, M.: Annual and
420 size dependent variation of growth rates and ion concentrations in boreal forest, *Boreal Env. Res.*, 10,
421 357–369, 2005.
- 422 Kerminen, V.-M., Chen, X., Vakkari, V., Petäjä, T., Kulmala, M., and Bianchi, F.: Atmospheric new
423 particle formation and growth: review of field observations, *Environ. Res. Lett.*, 13, 103003,
424 <https://doi.org/10.1088/1748-9326/aadf3c>, 2018.
- 425 Kivekäs, N., Carpman, Jimmie, Roldin, Pontus, Leppä, Johannes, O'Connor, Ewan,
426 Kristensson, Adam, and Asmi, E.: Coupling an aerosol box model with one-dimensional flow: a
427 tool for understanding observations of new particle formation events, *Tellus B: Chemical and*
428 *Physical Meteorology*, 68, 29706, <https://doi.org/10.3402/tellusb.v68.29706>, 2016.
- 429 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M.,
430 Aalto, P. P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and
431 Kerminen, V.-M.: Measurement of the nucleation of atmospheric aerosol particles, *Nat. Protocols*, 7,
432 1651–1667, <https://doi.org/10.1038/nprot.2012.091>, 2012.
- 433 Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petäjä, T.,
434 Sipilä, M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Järvinen, E., Äijälä, M.,
435 Kangasluoma, J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkilä, J., Vanhanen, J., Aalto, J., Hakola,
436 H., Makkonen, U., Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamäki, H., Back, J.,
437 Kortelainen, A., Riipinen, I., Kurten, T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F.,
438 Lehtinen, K. E. J., Laaksonen, A., Kerminen, V.-M., and Worsnop, D. R.: Direct observations of



- 439 atmospheric aerosol nucleation, *Science*, 339, 943–946, <https://doi.org/10.1126/science.1227385>,
440 2013.
- 441 Lampilahti, J., Paasonen, P., Tuovinen, S., Lehtipalo, K., Kerminen, V.-M., and Kulmala, M.: Dataset
442 for “A Cross-Correlation-Based Method for Determining Size-Resolved Particle Growth Rates,”
443 <https://doi.org/10.5281/zenodo.15648699>, 2025.
- 444 Lehtipalo, K., Leppä, J., Kontkanen, J., Kangasluoma, J., Franchin, A., Wimmer, D., Schobesberger,
445 S., Junninen, H., Petäjä, T., Sipilä, M., Mikkilä, J., Vanhanen, J., Worsnop, D. R., and Kulmala, M.:
446 Methods for determining particle size distribution and growth rates between 1 and 3 nm using the
447 Particle Size Magnifier, *Boreal Environ. Res.*, 19, 22, 2014.
- 448 Manninen, H. E., Nieminen, T., Riipinen, I., Yli-Juuti, T., Gagné, S., Asmi, E., Aalto, P. P., Petäjä,
449 T., Kerminen, V.-M., and Kulmala, M.: Charged and total particle formation and growth rates during
450 EUCAARI 2007 campaign in Hyytiälä, *Atmos. Chem. Phys.*, 9, 4077–4089,
451 <https://doi.org/10.5194/acp-9-4077-2009>, 2009.
- 452 Manninen, H. E., Nieminen, T., Asmi, E., Gagné, S., Häkkinen, S., Lehtipalo, K., Aalto, P., Vana,
453 M., Mirme, A., Mirme, S., Hörrak, U., Plass-Dülmer, C., Stange, G., Kiss, G., Hoffer, A., Törö, N.,
454 Moerman, M., Henzing, B., de Leeuw, G., Brinkenberg, M., Kouvarakis, G. N., Bougiatioti, A.,
455 Mihalopoulos, N., O’Dowd, C., Ceburnis, D., Arneth, A., Svenningsson, B., Swietlicki, E., Tarozzi,
456 L., Decesari, S., Facchini, M. C., Birmili, W., Sonntag, A., Wiedensohler, A., Boulon, J., Sellegri, K.,
457 Laj, P., Gysel, M., Bukowiecki, N., Weingartner, E., Wehrle, G., Laaksonen, A., Hamed, A.,
458 Joutsensaari, J., Petäjä, T., Kerminen, V.-M., and Kulmala, M.: EUCAARI ion spectrometer
459 measurements at 12 European sites – analysis of new particle formation events, *Atmos. Chem. Phys.*,
460 10, 7907–7927, <https://doi.org/10.5194/acp-10-7907-2010>, 2010.
- 461 Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J., and Carslaw, K. S.: Impact of
462 nucleation on global CCN, *Atmos. Chem. Phys.*, 9, 8601–8616, 2009.
- 463 Mirme, A., Tamm, E., Mordas, G., Vana, M., Uin, J., Mirme, S., Bernotas, T., Laakso, L., Hirsikko,
464 A., and Kulmala, M.: A wide-range multi-channel Air Ion Spectrometer, *Boreal Environ. Res.*, 12,
465 247–264, 2007.
- 466 Mirme, S. and Mirme, A.: The mathematical principles and design of the NAIS – a spectrometer for
467 the measurement of cluster ion and nanometer aerosol size distributions, *Atmos. Meas. Tech.*, 6,
468 1061–1071, <https://doi.org/10.5194/amt-6-1061-2013>, 2013.
- 469 Nieminen, T., Kerminen, V.-M., Petäjä, T., Aalto, P. P., Arshinov, M., Asmi, E., Baltensperger, U.,
470 Beddows, D. C. S., Beukes, J. P., Collins, D., Ding, A., Harrison, R. M., Henzing, B., Hooda, R., Hu,
471 M., Hörrak, U., Kivekäs, N., Komsaare, K., Krejci, R., Kristensson, A., Laakso, L., Laaksonen, A.,



- 472 Leaitch, W. R., Lihavainen, H., Mihalopoulos, N., Németh, Z., Nie, W., O'Dowd, C., Salma, I.,
473 Sellegrì, K., Svenningsson, B., Swietlicki, E., Tunved, P., Ulevicius, V., Vakkari, V., Vana, M.,
474 Wiedensohler, A., Wu, Z., Virtanen, A., and Kulmala, M.: Global analysis of continental boundary
475 layer new particle formation based on long-term measurements, *Atmospheric Chemistry and Physics*,
476 18, 14737–14756, <https://doi.org/10.5194/acp-18-14737-2018>, 2018.
- 477 Paasonen, P., Peltola, M., Kontkanen, J., Junninen, H., Kerminen, V.-M., and Kulmala, M.:
478 Comprehensive analysis of particle growth rates from nucleation mode to cloud condensation nuclei
479 in boreal forest, *Atmospheric Chemistry and Physics*, 18, 12085–12103, <https://doi.org/10.5194/acp-18-12085-2018>, 2018.
- 481 Riccobono, F., Rondo, L., Sipilä, M., Barmet, P., Curtius, J., Dommen, J., Ehn, M., Ehrhart, S.,
482 Kulmala, M., Kürten, A., Mikkilä, J., Paasonen, P., Petäjä, T., Weingartner, E., and Baltensperger,
483 U.: Contribution of sulfuric acid and oxidized organic compounds to particle formation and growth,
484 *Atmospheric Chemistry and Physics*, 12, 9427–9439, <https://doi.org/10.5194/acp-12-9427-2012>,
485 2012.
- 486 Sihto, S.-L., Kulmala, M., Kerminen, V.-M., Dal Maso, M., Petäjä, T., Riipinen, I., Korhonen, H.,
487 Arnold, F., Janson, R., Boy, M., Laaksonen, A., and Lehtinen, K. E. J.: Atmospheric sulphuric acid
488 and aerosol formation: implications from atmospheric measurements for nucleation and early growth
489 mechanisms, *Atmospheric Chemistry and Physics*, 6, 4079–4091, <https://doi.org/10.5194/acp-6-4079-2006>, 2006.
- 491 Stolzenburg, D., Fischer, L., Vogel, A. L., Heinritzi, M., Schervish, M., Simon, M., Wagner, A. C.,
492 Dada, L., Ahonen, L. R., Amorim, A., Baccarini, A., Bauer, P. S., Baumgartner, B., Bergen, A.,
493 Bianchi, F., Breitenlechner, M., Brilke, S., Mazon, S. B., Chen, D., Dias, A., Draper, D. C., Duplissy,
494 J., Haddad, I. E., Finkenzeller, H., Frege, C., Fuchs, C., Garmash, O., Gordon, H., He, X., Helm, J.,
495 Hofbauer, V., Hoyle, C. R., Kim, C., Kirkby, J., Kontkanen, J., Kürten, A., Lampilahti, J., Lawler,
496 M., Lehtipalo, K., Leiminger, M., Mai, H., Mathot, S., Mentler, B., Molteni, U., Nie, W., Nieminen,
497 T., Nowak, J. B., Ojdanic, A., Onnela, A., Passananti, M., Petäjä, T., Quéléver, L. L. J., Rissanen, M.
498 P., Sarnela, N., Schallhart, S., Tauber, C., Tomé, A., Wagner, R., Wang, M., Weitz, L., Wimmer, D.,
499 Xiao, M., Yan, C., Ye, P., Zha, Q., Baltensperger, U., Curtius, J., Dommen, J., Flagan, R. C.,
500 Kulmala, M., Smith, J. N., Worsnop, D. R., Hansel, A., Donahue, N. M., and Winkler, P. M.: Rapid
501 growth of organic aerosol nanoparticles over a wide tropospheric temperature range, *PNAS*, 115,
502 9122–9127, <https://doi.org/10.1073/pnas.1807604115>, 2018.
- 503 Yli-Juuti, T., Nieminen, T., Hirsikko, A., Aalto, P. P., Asmi, E., Hörrak, U., Manninen, H. E.,
504 Patokoski, J., Dal Maso, M., Petäjä, T., Rinne, J., Kulmala, M., and Riipinen, I.: Growth rates of
505 nucleation mode particles in Hyytiälä during 2003–2009: variation with particle size, season, data
506 analysis method and ambient conditions, *Atmos. Chem. Phys.*, 11, 12865–12886,
507 <https://doi.org/10.5194/acp-11-12865-2011>, 2011.



- 508 Yu, F. and Luo, G.: Simulation of particle size distribution with a global aerosol model: contribution
509 of nucleation to aerosol and CCN number concentrations, *Atmos. Chem. Phys.*, 9, 7691–7710, 2009.
- 510 Zhao, B., Donahue, N. M., Zhang, K., Mao, L., Shrivastava, M., Ma, P.-L., Shen, J., Wang, S., Sun,
511 J., Gordon, H., Tang, S., Fast, J., Wang, M., Gao, Y., Yan, C., Singh, B., Li, Z., Huang, L., Lou, S.,
512 Lin, G., Wang, H., Jiang, J., Ding, A., Nie, W., Qi, X., Chi, X., and Wang, L.: Global variability in
513 atmospheric new particle formation mechanisms, *Nature*, 631, 98–105,
514 <https://doi.org/10.1038/s41586-024-07547-1>, 2024.