

Iberulite fall and formation mechanism during a Sahara dust event in Switzerland in February 2021

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Abstract. Saharan dust events can produce micrometer-sized spherical mineral aggregates called iberulites, typically found on the Iberian Peninsula. To our knowledge, iberulites have not yet been reported from Central Europe. Two formation mechanisms—in-cloud and below-cloud scavenging—have been proposed. A strong iberulite fall (IF) was observed in Switzerland in February 2021, when large amounts of dust were transported from the Moroccan-Algerian border toward Central Europe during a Saharan Dust Event (SDE). Unlike previous IFs observed on the Iberian Peninsula, this IF occurred under low-temperature conditions near the freezing point, both in the cloud and at the surface. The relative humidity in the first 1000 meters below the clouds dropped from 70% to 40%. Comparison of the particle-size distribution (PSD) of the iberulite constituents with that in the free atmosphere revealed a deficit of particles in the 0.3–1.0 μm range relative to the dust in the free atmosphere. The meteorological conditions and the microstructure of the iberulites suggest that frozen hydrometeors act as scavengers, and that below-cloud scavenging is the most probable formation mechanism. Further discoveries of iberulites in Swiss SDE samples indicate that this phenomenon is widespread and not limited to regions near the dust source.

1 Introduction

Saharan Dust Events (SDEs) occur 10-40 times each year over the Alps, greatly increasing aerosol levels. They peak from February to June and from October to November Flentje et al. (2015); (Coen et al., 2004; Mauro et al., 2019). Since 2020, SDEs in February and March have become more frequent, particularly in the Western Euro-Mediterranean region (Cuevas-Agulló et al., 2024), with higher dust concentrations per event. This trend is driven by westerly wind blockages shifting poleward due to

31 blocking highs and the development of cutoff subtropical lows. During these events, aerosol levels often exceed European and
32 Swiss standards (24-hour limit: $50 \mu\text{g m}^{-3}$), with ground-level concentrations during SDEs exceeding the limits by a factor of 2.
33 Desert dust has been associated with adverse health effects, including increased cardiovascular mortality and respiratory issues,
34 although impacts vary by the dust's geographic source (Fussell and Kelly, 2021; Tobias et al., 2019). In Europe, dust accounts
35 for over 40,000 deaths annually (Wang et al., 2020). To date, 150 countries are affected, with more than 100 classified as non-
36 dust-source regions (Li et al., 2025). Due to the growing threat of sand and dust storms, international agencies have prioritized
37 mitigation efforts over the past twenty years (Unccd, 2022). Dust also influences climate and ecosystems (Buseck and Pósfai,
38 1999; Lian et al., 2025) by affecting Earth's radiative balance (Heald et al., 2014), degrading air quality (Proestakis et al., 2025),
39 supplying nutrients to marine and terrestrial ecosystems (Kellogg and Griffin, 2006), reducing snow and ice albedo and
40 accelerating melting (Francis et al., 2018), and adsorbing trace gases, thereby altering their physical properties (Tan et al., 2024).
41 Saharan dust notably impacts hail formation in Europe (Brennan and Wilhelm, 2025). Limited research exists on the economic
42 impact of desert dust, and no standardized data collection or analysis is in place (Middleton, 2017).

43 A distinct type of dust particles are spherule-shaped aggregates measuring between 50 and 300 μm in diameter, called iberulites,
44 a term originating from their first discovery in southern Spain (Díaz-Hernández and Párraga, 2008). Iberulites have also been
45 identified in Mallorca (Fiol et al., 2005; Fiol and Guijarro, 2000) and in Tenerife (Spain) (Cuadros et al., 2015), but no sightings
46 have been reported so far north of 42°N . Iberulites found in Spain display flow textures, including typical depressions (vortices)
47 and a shell structure with a coarse-grained core and a dense, fine-grained surface layer (Díaz-Hernández and Párraga, 2008;
48 Párraga et al., 2021). The iberulite constituents consist of minerals such as silicates, carbonates, sulfates, halides, oxides, and
49 phosphate-vanadate's, along with plant material, silica shells, plankton, and probably viruses. This mineral composition
50 suggests the Sahara and Sahel as primary mineral sources, while other minerals such as gypsum and alunite-jarosite are likely
51 formed in the atmosphere (neof ormation), probably via atmospheric pollutants during transport (Díaz-Hernández & Párraga,
52 2008; Párraga et al., 2021).

53 Two mechanisms were proposed for creating these aggregates (Díaz-Hernández & Sanchez-Navas, 2016; Párraga et al., 2021):
54 1. Coalescence of droplets within a cloud increases condensation nuclei within the resulting larger drop, known as in-cloud
55 scavenging (ICS) or 'nucleation scavenging,' and 2. an increase in concentration through collisions between dust particles and
56 drops falling from a cloud, called below-cloud scavenging (BCS) or impaction scavenging (Pruppacher and Klett, 2010). Both
57 processes are followed by the drying of the resulting drops.

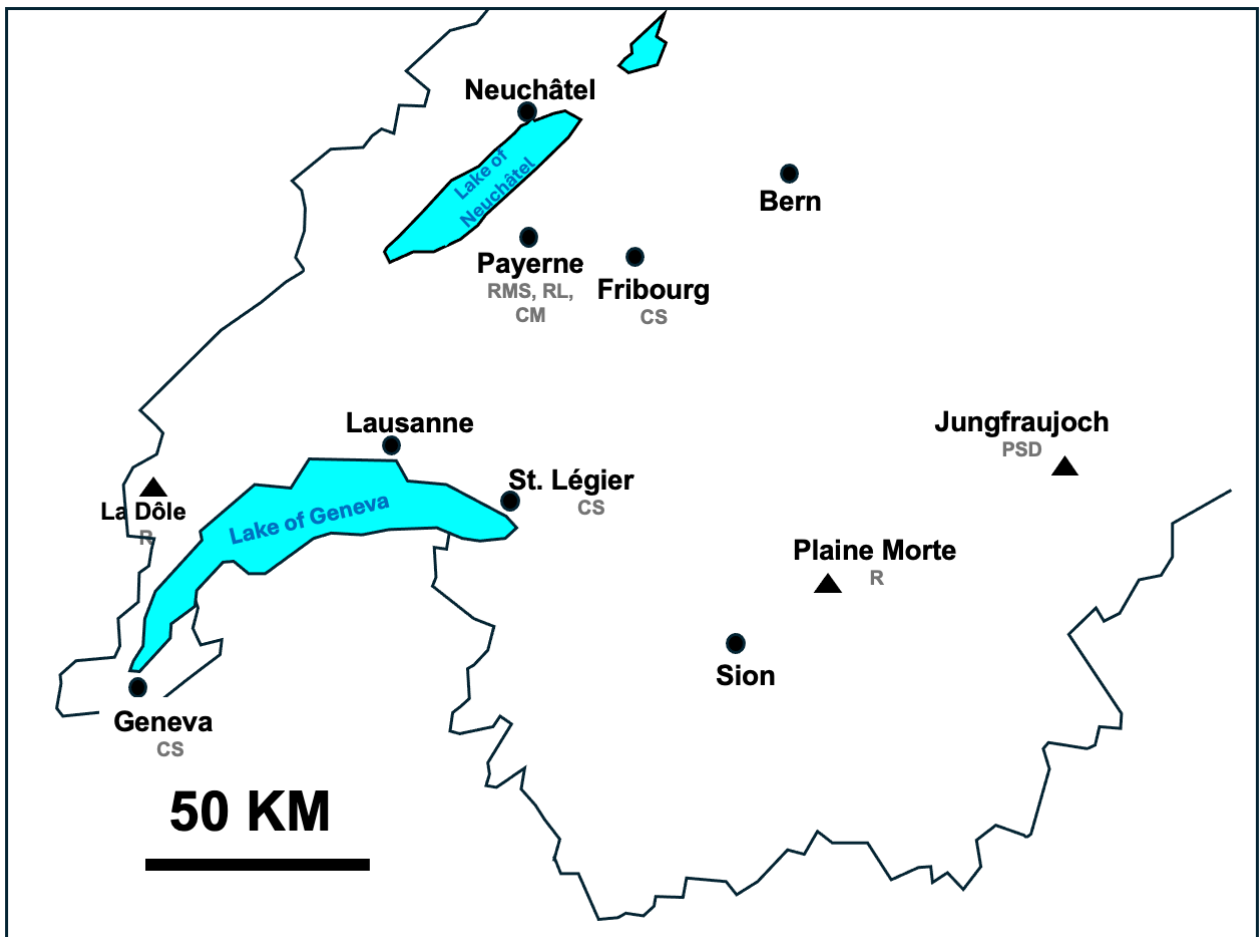
58 It is difficult to assess the importance of iberulites for climate, ecosystems, and health due to a lack of data. Most people who
59 collect bulk dust samples are probably unaware that iberulites may be present. Since the paper was submitted, analyses of
60 additional bulk aerosol samples collected in Switzerland have revealed that more than 50% contain iberulites, indicating that
61 they may be widespread. Important steps in deciding whether iberulites are important for models of dust deposition would be:
62 1) to check worldwide for the presence of iberulites and 2) to quantify the amount of dust deposited as iberulites.

63 A major SDE struck Switzerland on the afternoon of February 6th, 2021, making it one of the strongest since systematic
64 monitoring began in 2001 (Collaud Coen et al., 2025) at the Jungfraujoch laboratory (JFJ, inlet altitude 3585 m above sea
65 level). A record PM₁₀ concentration exceeding $700 \mu\text{g/m}^3$ was recorded (NABEL, BAFU/Empa). An iberulite fall (IF), he

56 first IF reported north of 42°N, was observed by the authors in Saint-Légier, Fribourg, and Geneva, all locations in western
57 Switzerland. In contrast to the Iberian IFs, the present fall happened under colder temperatures, and the meteorological
58 conditions were known along the iberulite trajectories. This study aims to characterize the collected iberulites, to determine
59 the possible formation mechanism, and to examine how the observed meteorological conditions influence the individual
70 steps in the formation of the iberulites in detail, leveraging the balloon radio-sonde data.

71 2 Methods

72 The fall of iberulites (IF) was observed (and heard!) standing outside Particle Vision's lab early in the afternoon on February
 73 6th in Fribourg(46.717° N, 7.083° E), but also in St. L  gier (46.471° N, 6.877° E), and Geneva (46.204° N, 6.143° E) (Fig.
 74 1, Table 1). Iberulites were collected at all three sites. The city of Fribourg (500-700 m ASL) is 50 km NNE of Lake Geneva
 75 on the Swiss Plateau, and the village of St. L  gier (576 m ASL) is on the NE shore of the same lake. The geology around
 76 both sites is dominated by calcareous mountains (Jura, Pr  alpes) (Matzenauer, 2011). In St. L  gier and Geneva, iberulites
 77 were collected from the ground with tweezers, and in Fribourg, an aluminum foil sheet was laid on the ground, from which
 78 iberulites were collected with a fine spatula. The iberulites were gently placed on a head specimen mount. Single dust
 79 particles sampled on the same foil were also collected and dispersed onto a particle filter (Millipore type AAWP, 47 mm,
 30 nitrocellulose, pore size 0.8   m). Collecting dust on aluminum foil is uncommon, but because the IF was unexpected,
 31 standard sampling equipment, such as a sampling head with an impactor and a pump, could not be set up in time. Aluminum
 32 foils are used to collect interstellar dust (Kearsley et al., 2010) and to identify laboratory sources of microplastic and
 33 nanoplastic contamination (Jones et al., 2024). Loss of particles during dispersal on the foil is negligible.



35 **Figure 1: Locations in western Switzerland with equipment mentioned in the text. Abbreviations: CS: collection site, R: weather**
36 **radar, RL: lidar radar, RMS: radio meteo sonde balloon launching site, PSD: particle size distribution determination site**

37 **Table 1** Sampling sites and analyses effectuated (iber.: iberulite; opt.: binocular imaging)

Site	Sampling	Dust PSD	Iber. PSD	Dust XRD	Dust SEM	Iber. Opt.	Meteo
Fribourg	x		x	x	x	x	
St. Léger	x		x	x	x	x	
Genève	x				x		
Jungfrauoch		x					x
Payerne							x

38

2.2 Observational data

Meteorological data were collected from radiosonde balloons launched by the Federal Office for Meteorology and Climatology of Switzerland (MeteoSwiss) station in Payerne (46.817° N, 6.933° E, Fig. 1) twice daily at 11:00 and 23:00 UTC. Payerne is located 17 km west and 47 km north of the collection sites in Fribourg and St. L  gier, respectively. The station is equipped with a Raman lidar, ceilometers, and a Doppler radar to measure precipitation, aerosol concentrations, and wind profiles. On February 6th, the ceilometer profile series from Payerne showed gaps in the morning and smaller gaps in the afternoon due to ground fog. The area is also covered by two weather radars: one on La Dole Mountain (46.425° N, 6.099° E), west of St. L  gier and near Geneva, and the other on Plaine Morte (46.383° N, 7.500° E, Fig. 1), in the central Bernese Alps southeast of the two collection sites.

The size distributions of fine- and coarse-mode aerosols in the dust cloud were measured at the high-altitude alpine station Jungfraujoch (JFJ, 7.98° N, 46.55° E; 3580 m ASL, Fig. 1) in the Bernese Alps. Its remote location, far from local human sources, makes it ideal for monitoring background atmospheric conditions. Jungfraujoch lab hosts a comprehensive long-term in situ monitoring program for aerosols and trace gases as part of the World Meteorological Organization (WMO) and is included in the Swiss National Air Pollution Monitoring Network (H  glin, 2021). Aerosol measurements were conducted using a custom-built Scanning Mobility Particle Size Spectrometer (SMPS) (Jur  nyi et al., 2011; Wiedensohler et al., 2012) and a white-light optical particle size spectrometer (FIDAS, Palas GmbH, Germany) (Fischer and H  glin, 2024; Brem et al., 2023).

2.3 Dust trajectory analysis

The dust RGB product, based on three infrared channels from the Spinning Enhanced Visible and Infrared Imager (SEVIRI) on EUMETSAT's second-generation MSG-11 satellite (Schmetz et al., 2002), was used to identify the source of dust arriving over western Switzerland in the afternoon of February 6th (Schepanski et al., 2009). The dust detection method is based on the decrease of the brightness temperature recorded by SEVIRI when dust is present in the atmosphere. Moderate dust loading can decrease the brightness temperature by 5–10 K (Sokolik, 2002). The plume's path was tracked using satellite images following the approach described by Ashpole and Washington (2013). The visually identified dust routes were confirmed using forward- and backward-trajectory calculations with the Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) model, developed by NOAA's Air Resources Laboratory (Draxler and Rolph, 2003). HYSPLIT uses archived three-dimensional meteorological data from ground and satellite observations, along with short-term forecasts. For this study, the Global Forecast System dataset from the National Centers for Environmental Prediction at 0.25° resolution was used to perform back-trajectory calculations. In all, 18 back-trajectory calculations have been performed: 3 arrival altitudes x 6 arrival times. The SEVIRI images and the backward trajectories trace the dust's origin for an arrival time over Fribourg at 2000 meters above ground level (AGL), between 12:00 and 18:00 UTC on February 6, roughly coinciding with the period of highest cloud density, to the Algerian-Moroccan border. The dust source meteorological data were obtained via the Real-time Environmental Applications and Display System (READY) from the archived Global Data Assimilation System (Rolph et al., 2017). The data were collected for the center of the main dust source observed in the satellite image, located at 33.8°N, 1.5°W.

23

24 **2.4 Scanning Electron Microscopy (SEM) and Optical Polarization Microscopy (OPM) analyses**

25 Both single particles and iberulites were isolated from samples collected on the ground in St. Legier and Geneva, as well as
 26 from the aluminum foil in Fribourg. A selection of iberulites was gently placed onto a ½-inch pin stub (Ted Pella, Redding,
 27 CA, USA) and examined using an Olympus SZX12 binocular microscope (Tokyo, Japan). All single particles collected on
 28 the foil in Fribourg were dispersed onto a carbon covered 1-inch pin stub and analyzed using automated SEM/EDX single-
 29 particle analysis (Computer-Controlled SEM, CCSEM). During CCSEM, particles are identified based on their Back-
 30 Scattered Electron (BSE) contrast, and images, along with Energy-Dispersive X-ray (EDX) spectra, are acquired. For this,
 31 a FEI Sirion FEG microscope equipped with an Oxford X-MAX EDX detector with a 150 mm² sensor size, a four-quadrant
 32 BSE detector, and AZtecFeature software (Oxford Instruments) was used at the Department of Geoscience of the University
 33 of Fribourg (Fribourg, Switzerland). Analyses were conducted at 20 kV. Particles were identified by simple gray-value (BSE
 34 intensity) thresholding. EDX analyses were performed for each identified particle. The AZtec software automatically applied
 35 a ZAF correction (PAP version of the f(pz) formalism, Pouchou and Pichoir (1985))—considering atomic number (Z),
 36 absorption (A), and fluorescence (F) effects—to the raw data for elemental concentrations. Quantification was based on
 37 factory-calibrated standards. The ZAF correction assumes homogeneous samples with smooth surfaces perpendicular to the
 38 beam and thicknesses greater than the excitation volume—conditions not met by these particles due to their rough surfaces
 39 and small sizes. However, studies (Kandler et al., 2018; Meier et al., 2018; Weinbruch et al., 1997) have shown that for
 40 particles larger than 1 µm, the relative standard deviation of the corrected major-element concentrations is approximately
 41 2%. For minor elements, the deviation is higher. Since only particles >1 µm in diameter were included in further data
 42 processing, ZAF-corrected concentrations were used without additional corrections. Several blank analyses of empty
 43 substrates have shown that, under the applied analytical conditions, the substrate surface contains, in an area of ca. 1 mm²,
 44 a maximum of five particles on unexposed substrates. These particles are organic in nature and most likely hair follicles.

45 **2.5 Particle size and microstructure analysis**

46 Images from both the optical microscope and the SEM were used to characterize the particle-size distribution of the
 47 individual constituents of the iberulites, as well as their microstructure, using Avizo software (Thermo Fisher). Selected
 48 iberulites were placed in a 1-inch round container and filled with impregnation resin. The resulting cylindrical samples were
 49 polished on the side containing the iberulites and then examined by electron microscopy. A BSE image of a polished iberulite
 50 section was also segmented using the Avizo software (Thermo Fisher), enabling the counting of individual particles and the
 51 determination of their equivalent spherical diameter d_{ve} , defined as the diameter of a sphere with the same volume as the
 52 particle with irregular shape (Kandler et al., 2018; Ott and Peters, 2008):

$$53 \quad d_{ve} = \frac{1}{P^2} \sqrt{64\pi B^3} \quad (1)$$

54 where B is the projected area, the particle covers on the sample substrate and P is the particle perimeter. Due to the resolution
 55 of the segmentation process, only particles larger than 0.05 µm can be detected and counted. To compare the SMPS with the

SEM-derived PSDs, the mobility diameter d_m obtained by SMPS must first be converted to the volume-equivalent diameter d_{ve} (Decarlo et al., 2004):

$$d_{ve} = \frac{C_c(d_{ve})}{C_c(d_m)\chi} d_m = CF d_m \quad (2)$$

with $C_c(d_m)/C_c(d_{ve})$, the Cunningham slip correction factor ratio between the mobility and volume equivalent diameters, CF the conversion factor, and the dynamic shape factor χ :

$$\chi = \frac{F_P}{F_{ve}} \quad (3)$$

which is defined as the ratio of the drag force on the nonspherical particle F_P to the drag force on its volume equivalent sphere F_{ve} when both move at the same relative velocity with respect to the gas (Hinds, 1999). The electrical mobility is the velocity v_e of a sphere ($c = 1$) with the same velocity as the measured particle:

$$v_e = \frac{neC_c(d_m)}{3\pi\eta d_m} = \frac{neC_c(d_{ve})}{3\pi\eta\chi d_{ve}} \quad (4)$$

From equation (4), the conversion equation (2) follows. $C_c(d_{ve})$ is unknown. The ratio $C_c(d_m)/C_c(d_{ve})$ has, to our knowledge, been measured only for salt particles smaller than 600 nm. The latter (extrapolated for particles > 600 nm) was used to convert the volume-equivalent diameter into the mobility diameter. It is always less than one, increases with size, and decreases with axis ratio. The axis ratios b/a (AR , the ratio of the longest dimension b to the orthogonal width a) and shape factors appear to depend on size, flow regime (Alexander et al., 2016), and orientation (Cheng et al., 1988), ranging from near unity for smaller particles to a value of 3 for the largest particles. The average AR inside the iberulites is 1.46, close to 1.4 reported for dust particles in the 2 to 10 μm range (Ansmann et al., 2011; Rocha-Lima et al., 2018; Jeong, 2024). Overall, CF is mostly smaller than 1.0. The converted volume-equivalent diameters are smaller than the mobility diameters.

2.6 X-ray diffraction

A loose, single dust particle sample, lacking iberulites but collected simultaneously with the iberulites that fell in Fribourg, was analyzed using a Rigaku Ultima IV diffractometer equipped with a position-sensitive detector. Diffractograms were recorded from 5° to 70° 2θ in step-scan mode ($0.02^\circ/\text{step}$ and $1 \text{ min}/^\circ$). The copper X-ray tube was operated at 40 kV and 40 mA. The Cu $K\alpha$ radiation was filtered through a nickel foil. Quantitative phase analysis via Rietveld refinement (Bish and Post, 1993) was performed on the recorded patterns using the PDXL2 software package (Rigaku). The refined instrumental parameters included the zero shift, sample displacement, and peak shape, modeled with a Pseudo-Voigt profile. Structure parameters for the phases in the sample were sourced from the International Center for Diffraction Data database; only the scale factor, unit-cell parameters, isotropic overall temperature factor, and, in some cases, the preferred orientation parameter (phyllosilicates, calcite) were the structural parameters refined using the March-Dollase formalism (Dollase, 1986). Phase fractions were determined through the scale factors. Atomic positions and site occupancies remained fixed during refinement. Multiple iterations were performed until the process converged.

3 Iberulite formation mechanisms

Proposed iberulite formation mechanisms are either in- (ICS) or below-cloud scavenging (BCS). Parameters influencing the number of particles in a final hydrometeor formed through ICS are critical supersaturation for activation, particle size, and temperature dependencies of the CCN/ICN activation (Lemaitre et al., 2024). The number of particles inside a hydrometeor in a cloud can also increase by collecting (scavenging) inactive particles through continued growth or collision (impaction) with them (Prodi, 1976; Prodi et al., 2018; Song and Lamb, 1994; Yang et al., 2023). Aggregation and coalescence efficiency of ice particles and droplets, respectively (Connolly et al., 2012), as well as the mineralogy (Atkinson et al., 2013; Zolles et al., 2015; Lüönd et al., 2010; Peckhaus et al., 2016; Harrison et al., 2019) of the particles, are further influencing parameters. In the global and annual mean, the aerosol mass scavenged in stratiform clouds was found to be primarily attributed to impaction processes. The aerosol number scavenged is primarily (>90%) attributed to impaction. 99% of this impaction scavenging occurs in clouds with temperatures less than 273 K. The in-cloud scavenging efficiency by impaction is for both ice columns and platelets almost constant for aerosol particles 0.1 -5µm in diameter (Croft et al., 2010).

For BCS, the capture efficiency Λ , i.e., the number of particles captured by a drop or ice particle per unit time along the trajectory during the fall, depends on particle size and the capture mechanism. If a single drop with radius R falls vertically, i.e., without deviation by wind, the maximum number of possible collisions N with particles with radius r along a trajectory l within the dust layer, mass density M_{air} , and particle density d can be calculated:

$$N = \frac{3 \Lambda M_{air} R^2 l}{4 \pi r^3 d} \quad (5)$$

The approximate number N_{total} in an iberulite with radius R was calculated using the segmentation by AVIZO of a BSE image of a polished section of an iberulite to extract the number of particles in the central section N_{sec} and using Eq.1 for calculating the average equivalent diameter d_{eq} of the particles. To calculate N_{total} , the iberulite sphere was cut into i flat discs with radius r_i , d_{ve} thick, and multiplied with the same particle number density N_{sec} as the central section:

$$N_{total} = \sum_1^{R/d} \frac{r_i^2}{R^2} N_{sec} \quad (6)$$

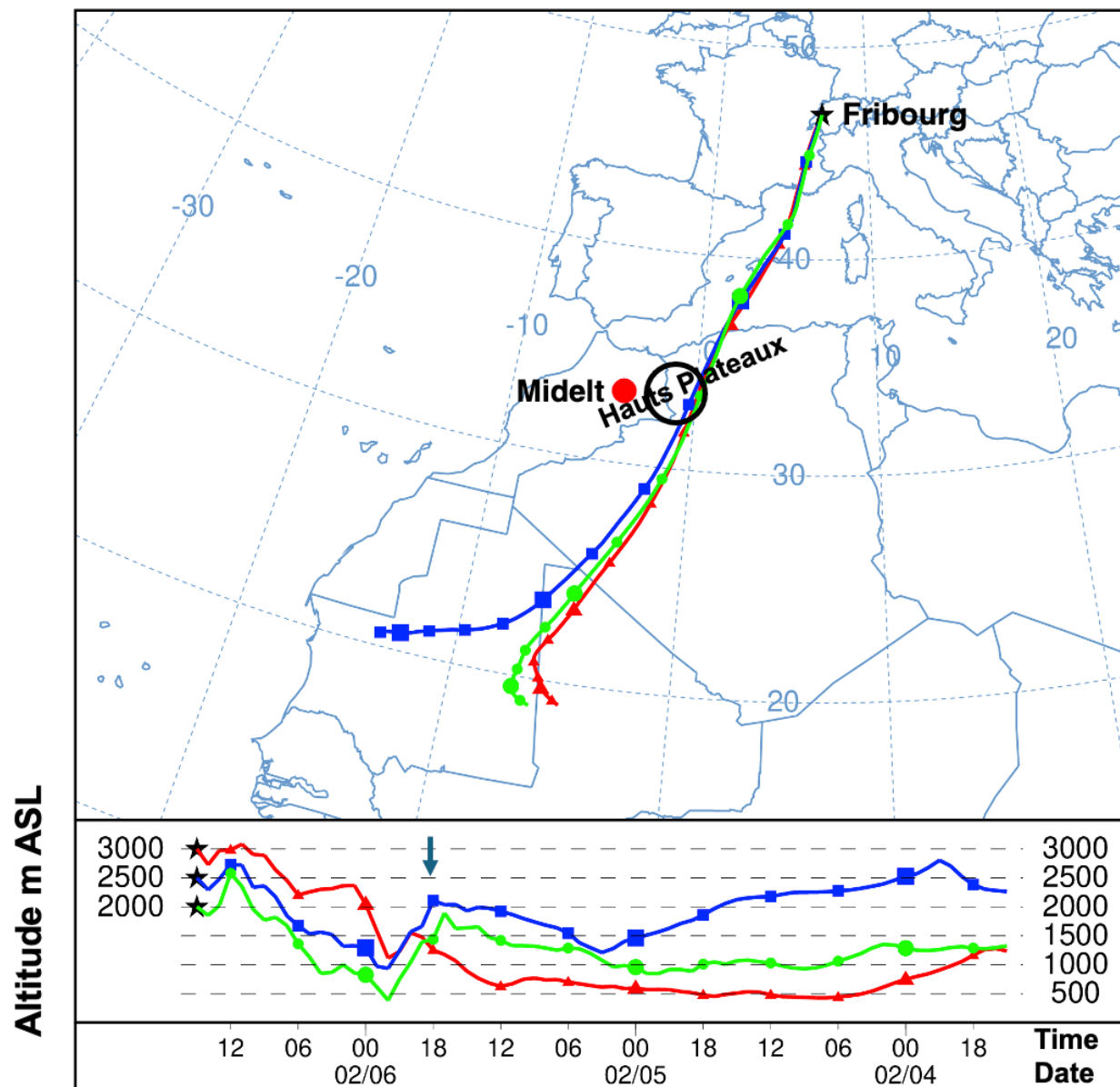
For particles larger than 1 µm in diameter colliding head-on with a droplet, bouncing is unlikely, and the sticking efficiency is assumed to be 100%. However, for particles with radii between 0.3 and 1.0 µm, Λ is less than 1. This occurs because of the airflow around the falling droplet. Particles in this size range, although on collision trajectories, tend to follow the flow lines and are diverted around the droplet. Larger particles, due to their inertia, cross the flow lines and impact the droplet directly, resulting in a Λ of 1 (capture by impaction). Small particles (< 0.3 µm) that follow the flow lines may be pushed by Brownian motion into the droplet's wake and deposited on its rear end (Brownian capture), with Λ also close to 1. Capture can also occur through thermophoretic and diffusiophoretic forces (Santachiara et al., 2013). Thermophoretic forces push aerosol particles toward a frozen drop during evaporation and away from it during vapor deposition. During diffusive crystal growth, a flow known as Stefan's flow occurs near the hydrometeor surface, exerting a force on nearby aerosol particles that pulls them toward the surface (Santachiara et al., 2023). For some particles following the flow lines, the latter approach the droplet's surface more closely than the particles' radii do. The particle then contacts the droplet's surface and sticks to it (capture by interception). This

process has a relatively low capture efficiency. For particles with radii between 0.3 and 1.0 μm , none of the preferred capture processes are effective, and the Λ decreases by two orders of magnitude. This reduction in Λ was first described by Greenfield (1957), and the absence of particles in the 0.3–1.0 μm range is called the Greenfield gap. Both formation mechanisms should lead to a depletion of particles in the 0.3–1.0 μm range in the iberulites relative to the dust cloud concentration, but the decrease is much more pronounced for BCS. The particle size distribution (PSD) inside an iberulite, compared to the PSD in the free atmosphere, can therefore indicate the formation mechanism of the iberulite.

4 Results

4.1 Dust source and trajectory

For air masses arriving at 2000, 2500, and 3000 m ASL over Fribourg in the early afternoon of February 6th, HYSPLIT back-trajectory calculations were performed for hourly arrival times from 9:00 to 15:00 h.

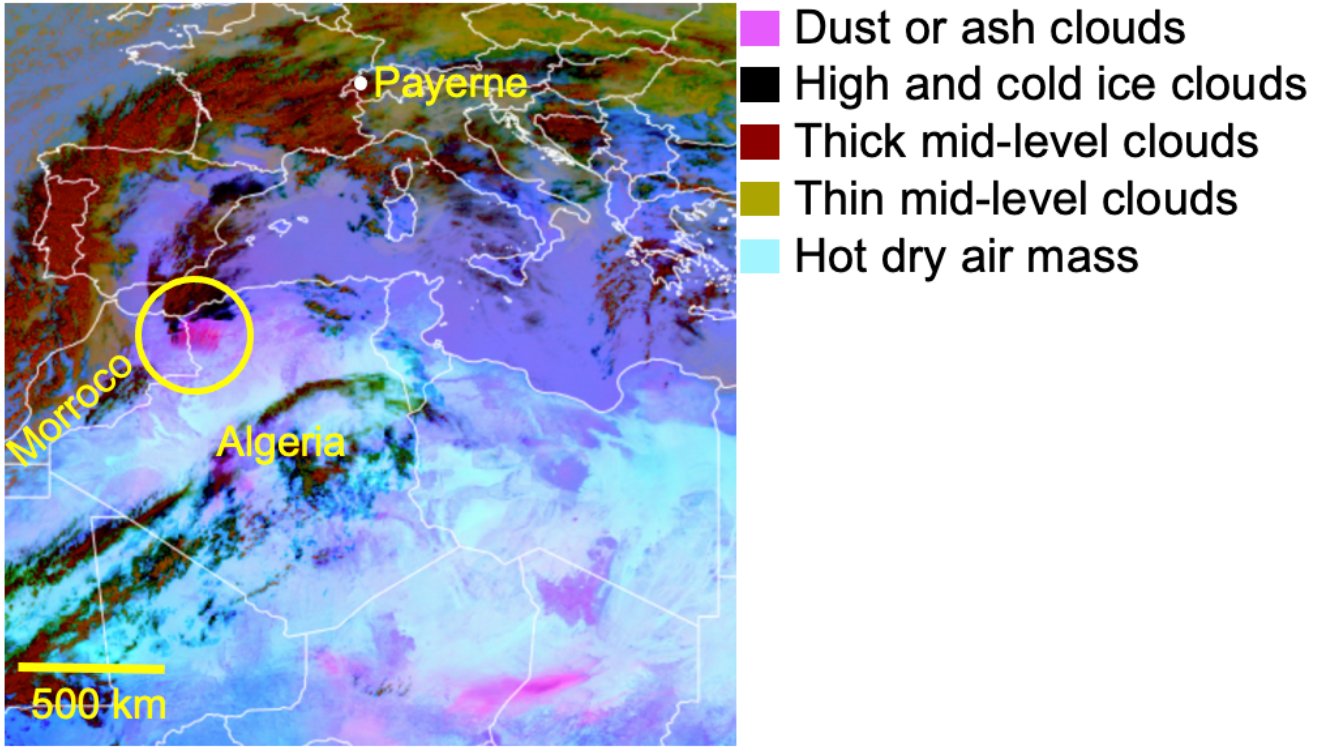


30

31 **Figure 2: Back trajectories for air masses arriving over Payerne at 15:00 UTC, February 6th. The black circle shows the location**
 32 **of the first appearance of the dust plume entrained across the Mediterranean Sea towards central Europe. Midelt is the location**
 33 **where meteorological conditions during the pickup of the dust are known. The black arrow points to the location markers inside**
 34 **the black circle e.g. close to the dust plume visible in Figure 2 (yellow circle).**

35 The dust originated in northwest Africa (Fig. 2), between the southern Moroccan border and northwestern Algeria. Satellite
 36 images showed a dust plume forming in a region called "Hauts Plateaux" (also known as "Hautes Plaines") between the Pre-
 37 Sahara and the Tell Atlas (Fig.3). The air masses arriving over the Hauts Plateaux during the early afternoon of February 5th
 38 were at altitudes ranging from 1500 m to 3500 m ASL, which is within the Planetary Boundary Layer (PBL; calculated

39 with the READY software (Rolph et al., 2017)), spanning from 2600 to 4500 m. Surface average wind velocity (at 10 m
40 above ground) increased from 40 km/h to approximately 70 km/h.



41
42 **Figure 3: Satellite RGB image (5.2.2021. 22:00 UTC) from the SEVIRI dust product. The dust plume transported to central**
43 **Europe is seen east of the Moroccan – Algerian border (center of yellow circle).**

44 Weather stations in the Atlas Mountains recorded hurricane-force gusts (e.g., 117 km/h on February 5th, 2021 in Midelt,
45 Morocco, Fig. 2) (Hoshyaripour 2021). This exceeds the threshold wind speed for dust emissions defined for western Algeria
46 (Marticorena et al., 1997). The dust plume was transported northeastward toward the Mediterranean Sea, where a low-
47 pressure system developed. The associated warm conveyor belt lifted the dust to higher altitudes and carried it with the polar
48 jet across the Mediterranean in an "atmospheric river" towards Central Europe (Francis et al., 2022; Rautela et al., 2024).
49 Source sensitivities (the probability that an air mass originates from a given source along a certain path) for Saharan dust (fine
50 mode) indicate that the most common pathway is from the Sahara, over the Mediterranean, directly to the Alps (Collaud
51 Coen et al., 2026).

52 4.2 Local meteorology

53 On the afternoon of February 6th, 2021, the Swiss Plateau—including Fribourg—experienced high-altitude mixed-phase clouds
54 (altocumulus) that eventually gave way to sunny clearing periods. Ground temperature measurements and T/RH profiles from
55 the radiosonde released at 11:00 UTC in Payerne showed relatively high temperatures for February 6th in Western Switzerland

56 (4°C at the ground and a 0°C level at 2800 m ASL), along with a typical inversion situation (Fig. 4). Compared to conditions
57 during previous IFs over the Iberian Peninsula (Diaz-Hernandez and Sanchez-Navas, 2016), the temperature was, nevertheless,
58 much lower. The vertical profile of relative humidity (RH) was complex: RH decreased from 97 % at ground level to a low 48
59 % at 1000 m, then increased to 61 % at 3200 m ASL (Fig. 4). RH remained between 48 and 55% up to an altitude of 4300 m
60 ASL, where it began to increase to 97 % resulting in a 100 m thick layer between 5000 and 5100 m ASL with temperatures
61 below the frost point. At this altitude, altocumulus clouds were reported.

62 The Payerne ceilometer time profile (Fig. 5) shows a cloud base at 5000 m ASL at midnight, with reflections above about 2500
63 m (the 0°C level), which may be due to aerosol (SDE) or ice particles. Reflections from ice particles and dust are generally
64 difficult to distinguish. Early in the morning, the cloud base lowered to 4000 m, and the reflections extended down to 2500 m.
65 Below 2500 m ASL, the temperature is above 0°C and rises rapidly; therefore, reflections below 2500 m ASL in the afternoon
66 are very unlikely to be from ice particles. The below-cloud reflections appear streaky in the ceilogram. After a large gap in
67 ceilometer recordings between 9:00 and 11:00 UTC due to ground fog blocking the laser, a reflection wall extending to the
68 ground appeared. These strong reflections from dust particles persisted throughout the day. The meteorological cloud base
69 lowered from 4000 m to 2000 m during the evening. Around 22:30, three short periods with very strong reflections were
70 observed; during these times, precipitation was recorded on the ground. After 23:00, the reflections became less intense.

71 Lidar measurements were taken between 15:00 and 17:00, during the signal's peak values and the average was used to derive
72 a mass-concentration profile (Fig. 6) across the dust plume. The altitude range of 4000-5000 meters served as the "no aerosol"
73 reference. The profile was created using a lidar ratio of $0.57 \text{ m}^2/\text{sr} \pm 9 \text{ sr}$ at 532 nm (Haarig et al., 2022), an Ångström
74 coefficient of 0.5 for long-range transported dust in northern Europe (Ansmann et al., 2001), and a mass extinction ratio at
75 532 nm of $0.47 \pm 0.04 \text{ m}^2 \text{ g}^{-1}$ (Nemuc et al., 2013). The profile showed three peaks. The highest concentration of $81 \text{ } \mu\text{g}/\text{m}^3$
76 was observed at an altitude of 1675 meters. The second-largest concentration was recorded just above the ground, and a
77 third small peak was located just above the cloud base at approximately 3500 meters.

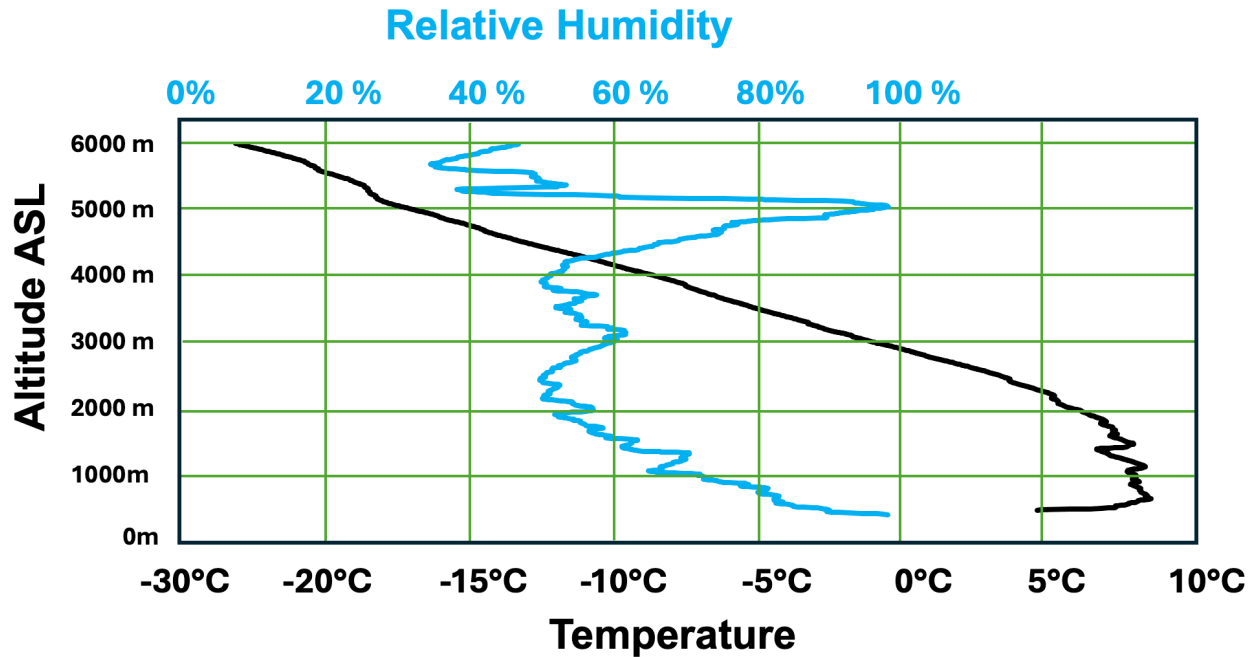


Figure 4: Temperature (black line) and relative humidity (blue line) as a function of altitude from the radio sounding. The balloon was released in Payerne at 11:00 h UTC. The balloon flew 30 km to the NW during the above recording (MeteoSwiss).

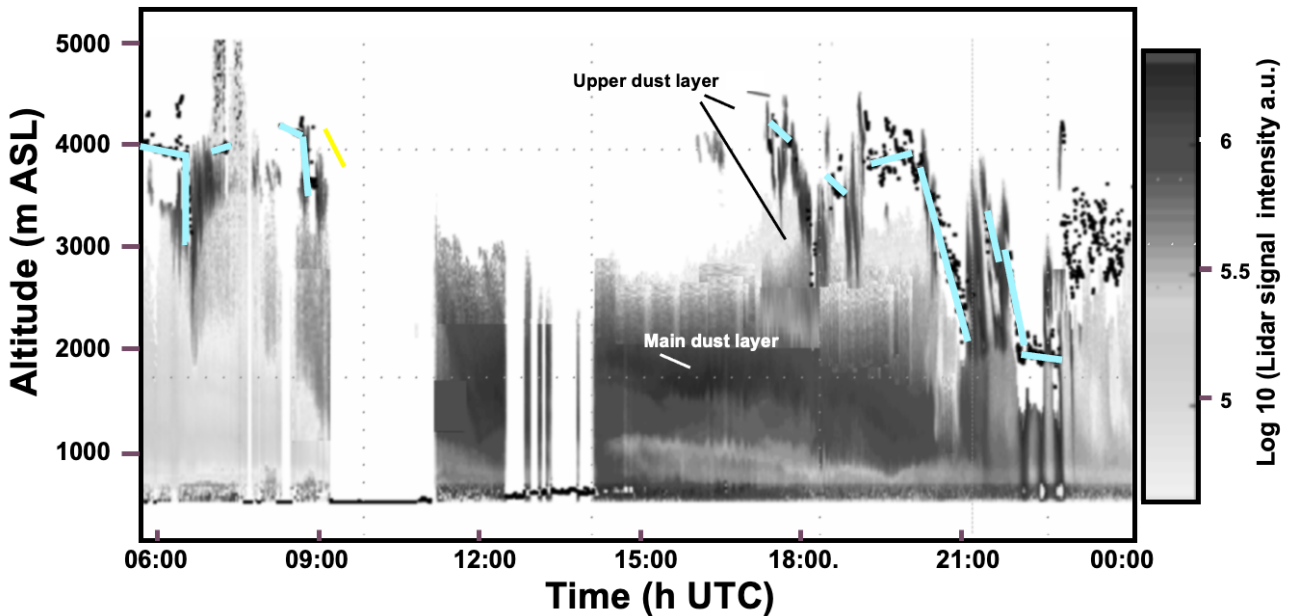


Figure 5: Ceilometer profile for February 6th (MeteoSwiss). The small black dots and blue lines above the reflections indicate the meteorological cloud base. The white strips are gaps due to ground fog.

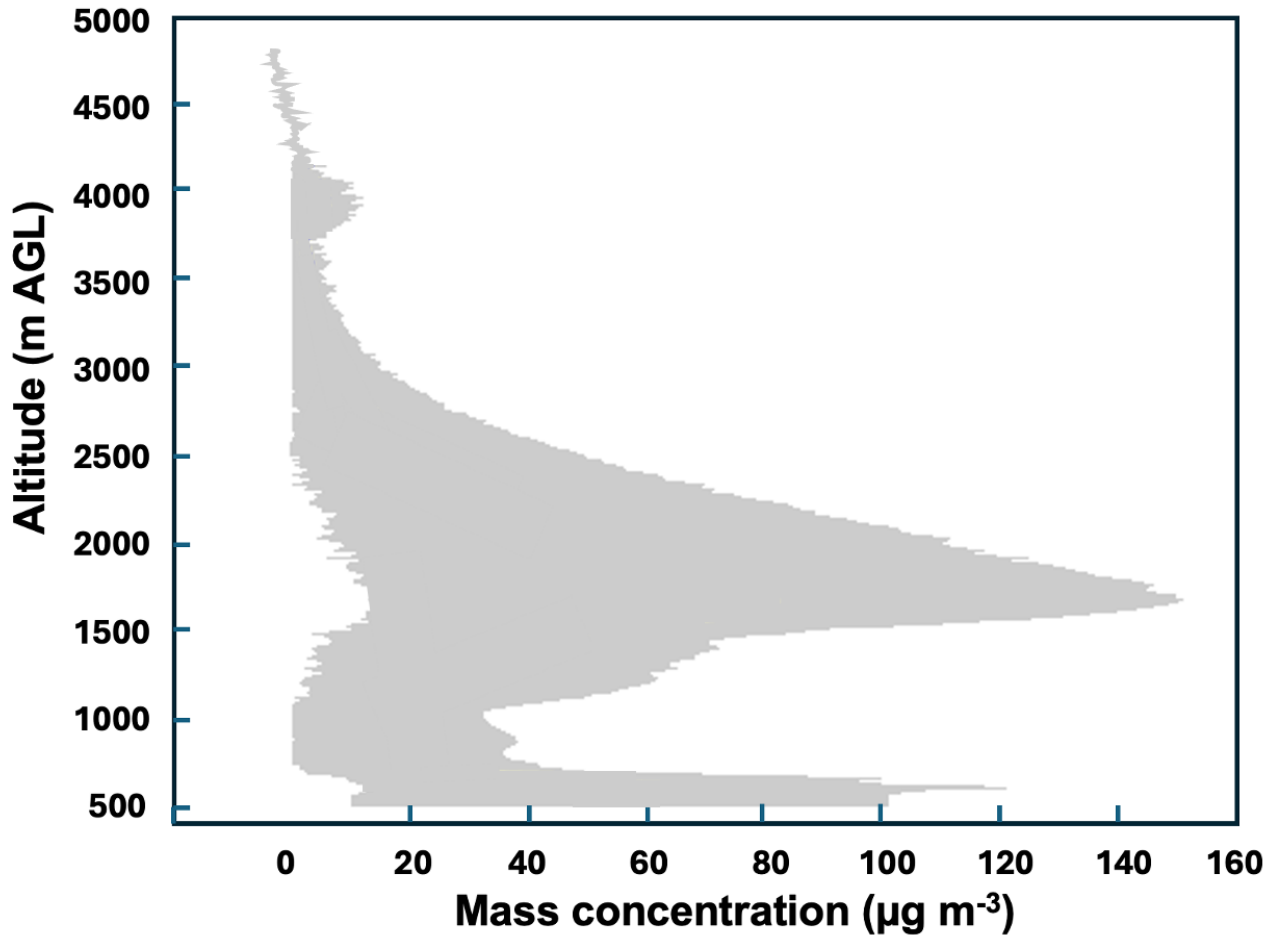


Figure 6: Aerosol mass concentration profile across the dust cloud calculated from the average of Lidar measurements taken between 15:00 and 16:59 h in Payerne (MeteoSwiss). The grey color indicates the bandwidth of the individual measurements

4.3 Morphology and microstructure of iberulites

Giant particles were previously observed following long-range transport of Saharan dust. Rapid transport, high turbulence, and electrical forces are thought to counteract the particles' weight. In our case, the large "particles" were not individual particles but multimineral particle aggregates with spherical to slightly ellipsoidal shapes, known as iberulites after their first recording on the Iberian Peninsula (Fig. 7) (Díaz-Hernández and Párraga, 2008). The cores of the iberulites consist of coarser mineral grains (Fig. 8 e,f, and Fig. 9) that have accumulated without a visible cement matrix. Apart from a thin, dense surface layer, the iberulite cores show no internal structure, such as layering or preferred grain orientation. The iberulites collected in St L  gier and Fribourg ($n = 425$) are larger than those from the Iberian Peninsula, with an average diameter of $113 \mu\text{m}$ (range: $43\text{--}365 \mu\text{m}$).

Most iberulites have smooth surfaces with few large surface pores (Fig. 8 a-d), most likely due to particles ejected from the aggregates. Spherical surface depressions, formed due to vortices in the air around the rear end of falling water drops (Ren

et al., 2020), as observed on Iberian iberulites (Díaz-Hernández and Párraga, 2008), are very rare among the sampled iberulites. Some iberulites were still wet.



Figure 7: Light microscopy image of iberulites collected in St Léger.

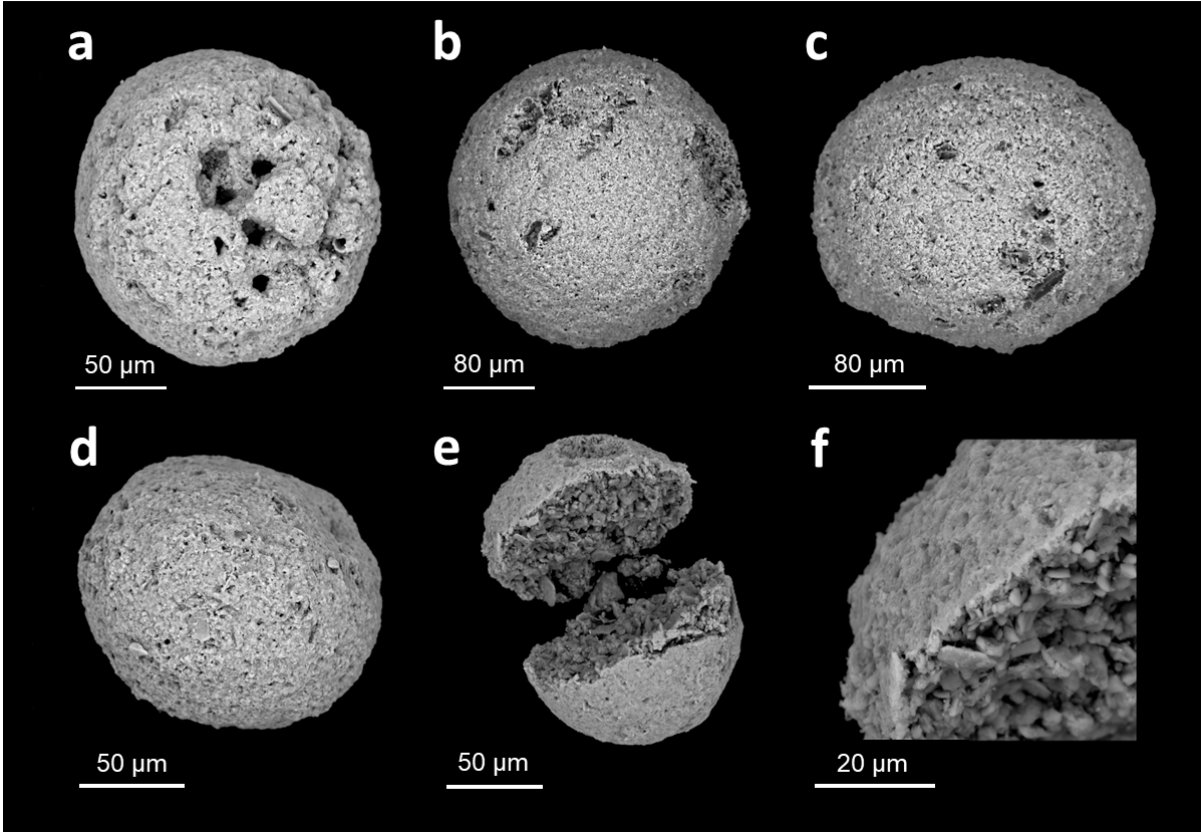


Figure 8: SEM images of some iberulites collected in St Léger on February 6th, 2021. Figure (f) shows the fine-grained rim and the coarse core of an iberulite.

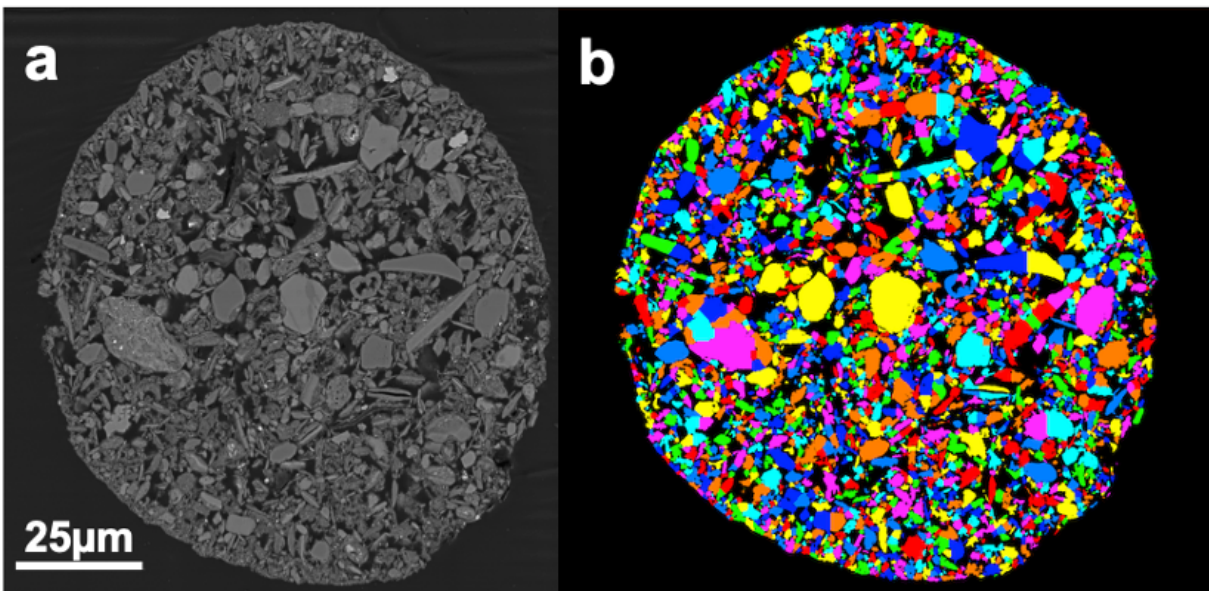


Figure 9: (a) BSE image of a polished section through an iberulite, and (b) the corresponding segmentation obtained from AVIZO. The different colors are used to better distinguish individual particles, but do not represent any specific elemental composition.

4.4 PSDs at JFJ and inside the iberulites

During periods when JFJ is cloud-free, above the PBL, and free of Saharan dust, the number PSD has two modes: a stronger one near $0.05\ \mu\text{m}$ and a weaker one between 0.15 and $0.4\ \mu\text{m}$ (Fig. 10, green bars). The PSD measured at 11:00 UTC on February 6th, when JFJ was within the Saharan dust plume, is monomodal, with a maximum near $0.150\ \mu\text{m}$ (Fig. 10, yellow bars). It shows, similar to PSDs observed during previous Saharan dust episodes at JFJ (Chou et al., 2011), much higher concentrations than measurements without SDEs. The PSD of particles within the iberulites is monomodal, with a maximum at $2.5\ \mu\text{m}$ (Fig. 10, blue bars), much larger than that for Saharan dust sampled at JFJ (e.g., $0.2\ \mu\text{m}$). For these dimensions, the PSD of the particles inside the iberulites shows rather a minimum.

The number of particles N_{sec} (598) in the iberulite section shown in Fig. 9a was determined after segmentation of the image, shown in Fig.9b, using AVIZO. The average equivalent diameter d_{eq} calculated with Eq.1 is $2.0\ \mu\text{m}$, and the total number of particles in the iberulite shown in Figure 9a is, using Eq. 8, 22'000! As a comparison, the number of collisions of a ice particle with a diameter of 1mm along a 1000m trajectory in air charged with $60\mu\text{gcm}^{-3}$ of dust with particles, $2.0\ \mu\text{m}$ in diameter, and an average density of $2.0\ \text{g cm}^{-3}$, assuming an average CE of 0.4, is calculated using Eq. 9: 26'000. Although both values are only an approximation due to the high variability of the variables entering the equations, it is nevertheless remarkable that they are of the same order of magnitude.

Coarse-mode particles arriving at JFJ were well documented in previous studies of particle volume distributions in dust plumes, with a mode in number PSDs at $1.4\ \mu\text{m}$ (Schwikowski et al., 1995) and in particle surface area distributions at $40\ \mu\text{m}$ (Brunner et al., 2021). Saharan dust clouds transported over long distances, such as to Barbados during the SALTRACE campaign (Weinzierl et al., 2017), to Aeronet sites in the eastern Caribbean (Velasco-Merino et al., 2018), over the Mediterranean Sea during the ChArMEx/ADRIED campaign (Denjean et al., 2016), and over Cape Verde during the SAMUN2 campaign (Weinzierl et al., 2011) have similar PSDs. The largest particles observed at the Caribbean sites are well above the maximum diameter expected for settling under Stokes' law during transport. The modes observed at JFJ and in the campaigns cited above are like the PSDs observed in the source regions (Panta et al., 2023; Bencherif et al., 2022; Dubovik et al., 2006), but are lower in concentration. The presence of particles with diameters $> 10\mu\text{m}$ inside the iberulites is thus not exceptional.

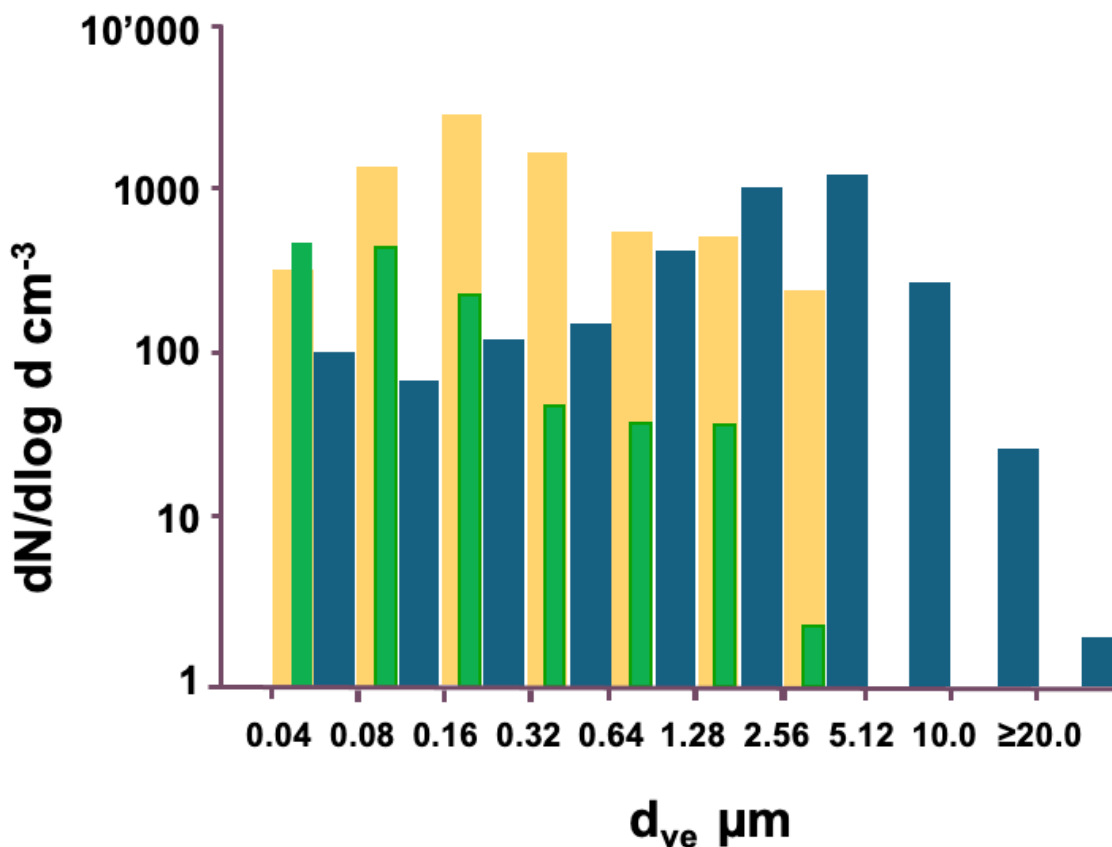
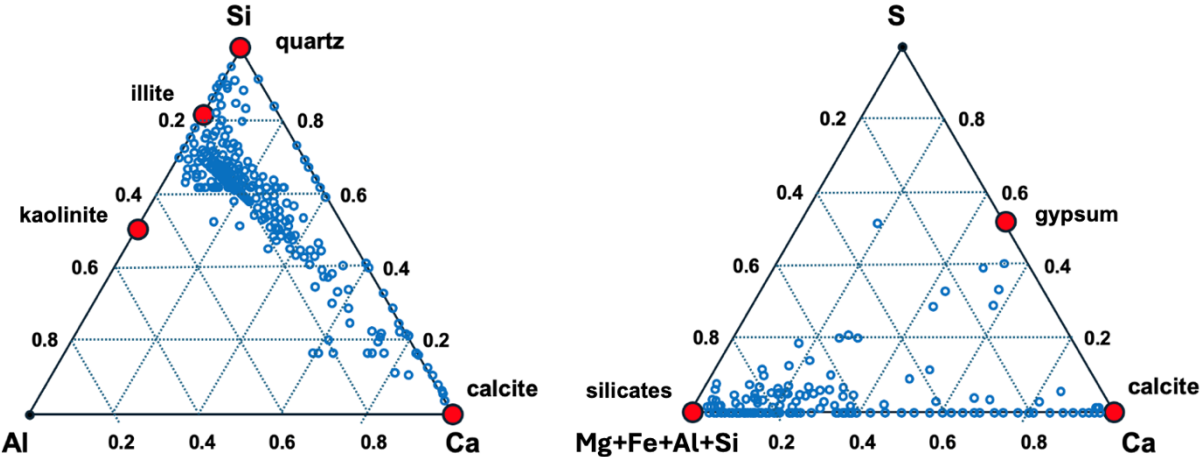


Figure 10: Jungfrauoch PSD measured in April 2004, when JFJ was above the PBL, cloud-free, and no SDE (green bars); during the SDE of February 6th at 12:00 (yellow bars, shifted 1 bin to the left relative to the original data, see text); and for an iberulite sample collected in St. L gier (blue bars).

4.5 XRD -results (Rietveld analysis) and EDS - analyses

X-ray diffraction from individual particle samples collected on the foil in Fribourg showed that the dust contained phyllosilicates (illite, muscovite, and kaolinite, totaling 35.1 wt%), followed by carbonates (28.0 wt% calcite, 4.0 wt% dolomite) and quartz (23.3 wt%) and K-feldspar 1-2wt%). The concentrations of gypsum and halite were below the detection limit. Analysis of individual particles within an iberulite confirms the XRD results. EDS analyses on the polished iberulite section have a resolution of 1–2 μm , depending on section thickness and local composition. Compositions of particles $<1 \mu m$ should therefore be interpreted with caution. Trends from the ideal muscovite composition toward the calcite composition, as well as toward the quartz compositions, in the (Mg+Fe+Si+K–Al–Ca) ternary diagram indicate that the particles are mixtures of the three phases (Fig. 11 a). The ternary diagram (Mg+Fe+Si+K–S–Ca), with Mg+Fe+Si+K as the "silicate" corner and Ca as the "calcite" corner (Fig. 11 b), shows measurements mainly along the joint between the silicate and calcite corners, indicating that the analyses are from particle aggregates containing both minerals, and some measurements scattered in the triangle silicates-calcite-gypsum,

48 indicating mixtures among these phases or the presence of sulfur on the particle surfaces. This lack of sulfur in the rims of the
 49 iberulites sampled in this campaign, differentiate them from the iberulites of the Iberian Peninsula (Díaz-Hernández and
 50 Párraga, 2008; Párraga et al., 2021).
 51



52
 53 **Figure 11: Ternary diagrams: (a) (Mg+Fe+Si+K–Al–Ca), (b) (Mg+Fe+Si+K–S–Ca).**

54 **5. Discussion**

55 **5.1 Dust composition**

56 Dust from potential source areas has been characterized by bulk chemistry (Kandler et al., 2007) and mineralogy (Scheuven et
 57 al., 2013a). Mineralogical fingerprints include phyllosilicates, such as the ratio of illite to kaolinite concentrations (I/K)
 58 (Caquineau et al., 1998; Chester and Johnson, 1971; Scheuven et al., 2013b). In Northern Africa, the general trend is an increase
 59 in I/K from S to N and from E to W. The value for the present single particles collected in Fribourg is approximately 2.9,
 60 consistent with values reported from the Western Sahara (e.g., >2 in NW Algeria and Morocco), which is the source area for the
 61 SDE determined from back-trajectory calculations and satellite data. High carbonate contents (>10 wt.%, calcite > dolomite)
 62 are also a fingerprint of dust from northwestern Africa (Avila et al., 1997). Similar to illite, carbonate content decreases
 63 continuously toward the south (Paquet, 1984). The high carbonate content is characteristic for dust emitted from regions within
 64 and adjacent to the Atlas Mountains (Grousset et al., 1992), the source for the present SDE determined by backtrajectory
 65 calculations and satellite imagery.

66

5.2 Iberulite formation and insights from internal structure

The layer at 5000 m altitude with supersaturated atmospheric conditions (0.11SS) in a mixed-phase cloud is most likely the source of the ice/snowflake precipitation that led to the scavenging process by frozen hydrometeors. With an RH of 97%, the highest value measured by the balloon sonde, and assuming a liquid water content $>0.01 \text{ g m}^{-3}$, the fraction of water bound in ice particles relative to the total water content is about 80% (Korolev et al., 2017). The conditions are ideal for the crystallization of plate-like ice particles (Fukuta and Takahashi, 1999; Hallett and Bailey, 2009; Libbrecht, 2005). The temperature at the upper boundary of the dust layer was below 0°C and became positive around 2800 m. However, melting of the frozen hydrometeors most likely began at a much lower level due to low relative humidity, in which the water vapor density of the surrounding air is lower than the equilibrium water vapor density of frozen hydrometeors at the freezing point; therefore, sublimation of water vapor occurs from the surface of the latter. Because of this sublimation process, latent cooling can push the surface temperature of the hydrometeor below 0°C (Matsuo and Sasyo, 1981; Heymsfield et al., 2015). The width of the non-melting layer in the atmosphere increases nearly linearly with decreasing relative humidity, from 120 m at 90% to 700 m at 50%. The relative humidity at the top level of the main Sahara dust layer is around 50% and remains constant down to an altitude of 1700 m ASL. It is therefore probable that the melting level is displaced far into the massive Saharan dust layer and that the encounter between dust and hydrometeors occurred when the latter were frozen. Although difficult to distinguish from aerosol particles, the reflections just below the cloud base may originate from frozen hydrometeors. Both solid ice crystals and snow crystals can capture particles. Several studies on BSC (Kyrö et al., 2009; Paramonov et al., 2011) indicate that frozen hydrometeors are more effective scavengers of aerosol particles than rain on an equivalent water-content basis. The fluttering, zigzagging fall style of frozen hydrometeors (Hashino et al., 2016; Kajikawa, 1992; Tagliavini et al., 2021; McCorquodale and Westbrook, 2021; Stringham et al., 1969) prolongs their trajectories, thereby increasing the probability of collisions with aerosol particles and, consequently, CE. CE also increases with temperature and relative humidity, although there is a discrepancy between experimental/natural data and theory (Santachiara et al., 2023). The highest CE has been recorded for mixed precipitation types at temperatures slightly below and above 0°C (Miller and Wang, 1989). The temperatures between the likely origin of the frozen hydrometeors and 2000 m ASL are in this range.

At these temperatures, frozen hydrometeors develop a stickier liquid layer on their surfaces. The increase in CE at higher RH is attributable to increased collector capacitance (Paramonov et al., 2011). Thus, the conditions for effective scavenging were met. The Greenfield gap for frozen hydrometeor fall situations was smaller, spanning $0.09\text{--}0.3 \text{ }\mu\text{m}$, compared with the gap observed for rain. This range is also observed in the iberulites collected in the present fieldwork, supporting the role of frozen hydrometeors as dust scavengers and that BSC was at work. The increased CE of frozen hydrometeors is probably also responsible for the larger diameters observed in Swiss iberulites than in Iberian iberulites. When frozen hydrometeors exit the non-melting zone, they melt or evaporate, exhibiting a wide range of melting behaviors (Hallett and Oraltay, 2005). All frozen hydrometeors form thin liquid films on the surface and become stickier (Knight, 1979)(Leinonen and Lerber, 2018)(Hallett and Oraltay, 2005). Large frozen hydrometeors increase in size through coalescence in the melting layer (McFarquhar et al., 2007). The behavior of scavenged aerosol particles within melting hydrometeors has not been analyzed to date. Most likely, they are incorporated into the newly formed raindrops. Within and below the melting layer, due to the very low relative humidity (45% at 1000 m, at the bottom of the massive dust layer), raindrops will dry very rapidly. A technological analog of this process is the (spray) drying of a powder suspension, which is used, e.g., in the ceramic, food, or pharmaceutical industry (Santos et al., 2018). The difference

between the drying of a droplet in a spray dryer and that of a rain droplet falling out of a cloud is the temperature difference between the droplet and the drying airstream, which is around 200°C in a spray dryer. Droplets in a spray dryer relax into spherical shapes, analogous to raindrops. The latent heat of water evaporation induces a strong heat flux from the droplet surface into the gas stream. This, in turn, drives thermophoretic particle displacement toward the droplet surface and a specific microcirculation in a layer near the surface, driven by the surface-tension gradient (Iskandar et al., 2003). This microcirculation entrains particles toward the droplet surface, leading to size- and density-based segregation. Smaller and/or less dense particles move faster toward the droplet surface, forming a surface layer. Coarser particles are concentrated in the core. Such a shell-core structure was observed in the present iberulite and has also been reported in the Iberian examples (Díaz-Hernández and Párraga, 2008; Párraga et al., 2021), with fine-grained phyllosilicates concentrated in the shell and coarse-grained quartz grains in the core. The sulfur-rich rims observed in Iberian iberulites, but lacking in our iberulite, were explained by the condensation of gaseous sulfuric acid on the iberulite rind and the absorption of atmospheric SO₂ attributed to industrial emissions in North Africa. If the present iberulites are formed through BCS above Western Switzerland with few heavy industries, such rims are not expected. Even if ICS before the dust cloud arrives over Switzerland should be the iberulite formation mechanism, the trajectory of the air masses arriving over Fribourg crosses only over Marseille (France) a region with heavy industries. The air masses crossed the Algerian sea border between the industrial hotspots of Algier and Arzew- Mostaganem (Etkind et al., 2026) and at 2000m ASL over Marseille, with a boundary level height which in wintertime is around 1000m (Riandet et al., 2023), e.g. below the dust cloud level.

6 Conclusion

To our knowledge, this analysis is the first detailed study of iberulite fall in Switzerland. In contrast to the Iberian cases, the IF fall occurred under much colder conditions, e.g., temperatures around 0° and decreasing relative humidity along the fall trajectory. Iberulites form through the accumulation of particles inside raindrops. A comparison of the PSD inside the iberulites with the PSD measured in the dust cloud at the Jungfraujoch, which encountered the same dust cloud, revealed a Greenfield gap in the iberulite PSD, i.e., a lack of particles in the range between 0.1 and 0.3 µm. The presence of a Greenfield gap within the iberulites supports below-cloud scavenging, i.e., particle capture via collisions with droplets below a cloud, as the mechanism for particle accumulation. The gap arises from particle trajectories in the flow around the falling droplet, which depend on particle size. The freezing temperature, low relative humidity, and ceilogram signals suggest that particle collisions occurred not with raindrops but with snowflakes or ice crystals. These are known to be better scavengers than raindrops. Since the paper was submitted the check of bulk aerosol samples revealed more than 50% contained iberulites.

Author contribution

RJ, FP and JD initiated the research. RJ, FP, JD and NC carried out the measurements. GB, RJ, FP, JVD and NC analyzed the data. SM provided data from Jungfraujoch. GB, wrote the paper. All co-authors proofread and commented on the paper. GB prepared the manuscript with contributions from all co-authors.

Competing interests

The authors declare that they have no conflict of interest.

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44

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